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THE GEOCHEMICAL STRATIGRAPHY OF THE VOLCANIC ROCKS OF THE
WITWATERSRAND TRIAD IN THE KLERKSDORP AREA,
TRANSVAAL

by

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Thesis submitted in fulfilment of the requirements for the degree
of Master of Science.

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January 1984

This thesis represents the original work
of the author except where specific
acknowledgement is made to the work
of others.

A handwritten signature in black ink, appearing to read "J.B. Bower". The signature is written in a cursive style with a long horizontal flourish extending to the right.

January 1984

ABSTRACT

This study was initiated with the aim of identifying the existence of any geochemical criteria which may be used to distinguish between the various volcanic formations within the Witwatersrand triad.

The Witwatersrand triad comprises three sequences: the Dominion Group at the base, the Witwatersrand Supergroup in the middle, and the Ventersdorp Supergroup at the top. It is underlain by Archaean basement rocks, and covered by rocks of the Transvaal sequence. The Dominion Group consists of the sedimentary Rhenosterspruit Quartzite Formation at the base, overlain by a bimodal succession of lavas. Basaltic lavas are the major component of the Rhenosterhoek Formation, while the overlying Syferfontein Porphyry Formation consists primarily of dacitic porphyries. Intercalations of one lava type within the other are common, however, so each formation is not the exclusive domain of only one lava type.

The Witwatersrand Supergroup, a predominantly argillaceous and arenaceous sequence, contains two narrow volcanic horizons, one of which, the Jeppestown Amygdaloid (now Crown Formation), consisting of tholeiitic andesites, occurs in the study area. The overlying Ventersdorp Supergroup has, at its base, the basaltic Klipriviersberg Group, of which four out of six formations are present in the study area, namely, the Alberton, Crkney, Loraine and Edenville Formations. This group is succeeded unconformably by the Platberg Group, consisting of the sedimentary Kameeldoorns Formation, followed by the (informal) Goedgenoeg, Makwassie Quartz Porphyry and Rietgat Formations. The Goedgenoeg and Rietgat Formations are basaltic, while the Makwassie rocks range from basaltic to dacitic, the majority being tholeiitic andesites and andesites.

The Pniel sequence at the top of the Ventersdorp Supergroup consists of the sedimentary Bothaville Formation, and the Allanridge Formation, the lavas of which are basaltic with some andesitic tendencies.

A well-defined geochemical stratigraphy was found to exist. From the eleven volcanic formations examined, nine distinct geochemical units emerged, as the Loraine and Edenville Formations were found to have the

same geochemical characteristics, as did the Goedgenoeg and Rietgat Formations. Despite having undergone low-grade greenschist facies metamorphism, very clear variation patterns with height are displayed by the immobile elements Ti, P, Nb, Zr and Y, and the light rare earth elements La, Ce and Nd. In contrast, much scatter was observed in the variation patterns of Na, K, Mn, Ba and Rb.

Three techniques were employed to effect discrimination between formations - orthogonal discrimination, interelement and ratio vs ratio plots, and discriminant analysis. Confidence limits placed on normal probability plots served to isolate outlier samples for further examination by the various discrimination techniques.

A successful test of the efficacy of the discrimination techniques was afforded when fourteen samples from an unknown succession were positively identified as representative of the Klipriviersberg Group.

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1. PREAMBLE

The Witwatersrand triad is a sequence of mildly deformed sedimentary and volcanic rocks, now classified into the Dominion Group, Witwatersrand Supergroup and Ventersdorp Supergroup. These rocks represent a continuous supracrustal sequence deposited in intracratonic basins of the Kaapvaal Province granite-greenstone belt, approximately between 2.8 and 2.3 Gy (see Figure 1.1). The distribution of these three sequences is shown in Figure 1.2, while currently recognised outcrops are indicated in Figures 1.3, 1.4 and 1.5. In addition to their close association spatially, structurally and temporally, the presence of detrital gold and uranium mineralisation has long provoked widespread interest in these rocks. They therefore have an extensive history of investigation, which is summarised in a literature review in Appendix I.

The name "Ventersdorp Beds" was first proposed by Hatch (1903c) for a succession of lavas, pyroclastic rocks and boulder beds that he found unconformably overlying the "Lower Witwatersrand Beds" and granite near Ventersdorp. During the following year the name was modified to "Ventersdorp Series" (Hatch and Corstorphine, 1904a).

Molengraaff (1905) suggested the name "Dominion Reef Series" for a succession of reefs found in the Klerksdorp district, comprising sericite schists interbedded with gold-bearing conglomerates. These reefs are overlain by amygdaloidal basic lavas and porphyritic lavas, which were subsequently included in the Dominion Reef Series by Nel (1934a). Truter (1949) elevated the sequence to the status of system, which it retained until SACS (1980) changed it to the rank of group. Although numerous attempts have been made to correlate the Dominion Group with other volcanic and volcano-sedimentary units in South Africa, present evidence implies that its development was restricted to the Western Transvaal (see Figure 1.3), where it represents a precursor to the main Witwatersrand basin (Watchorn, 1980).

Stratigraphically, the group is divided into three formations (see Figure 1.6):

- (1) The Rhenosterspruit Quartzite Formation at the base composed mainly of coarse arkosic sediments, lenticular grits and conglomerates;

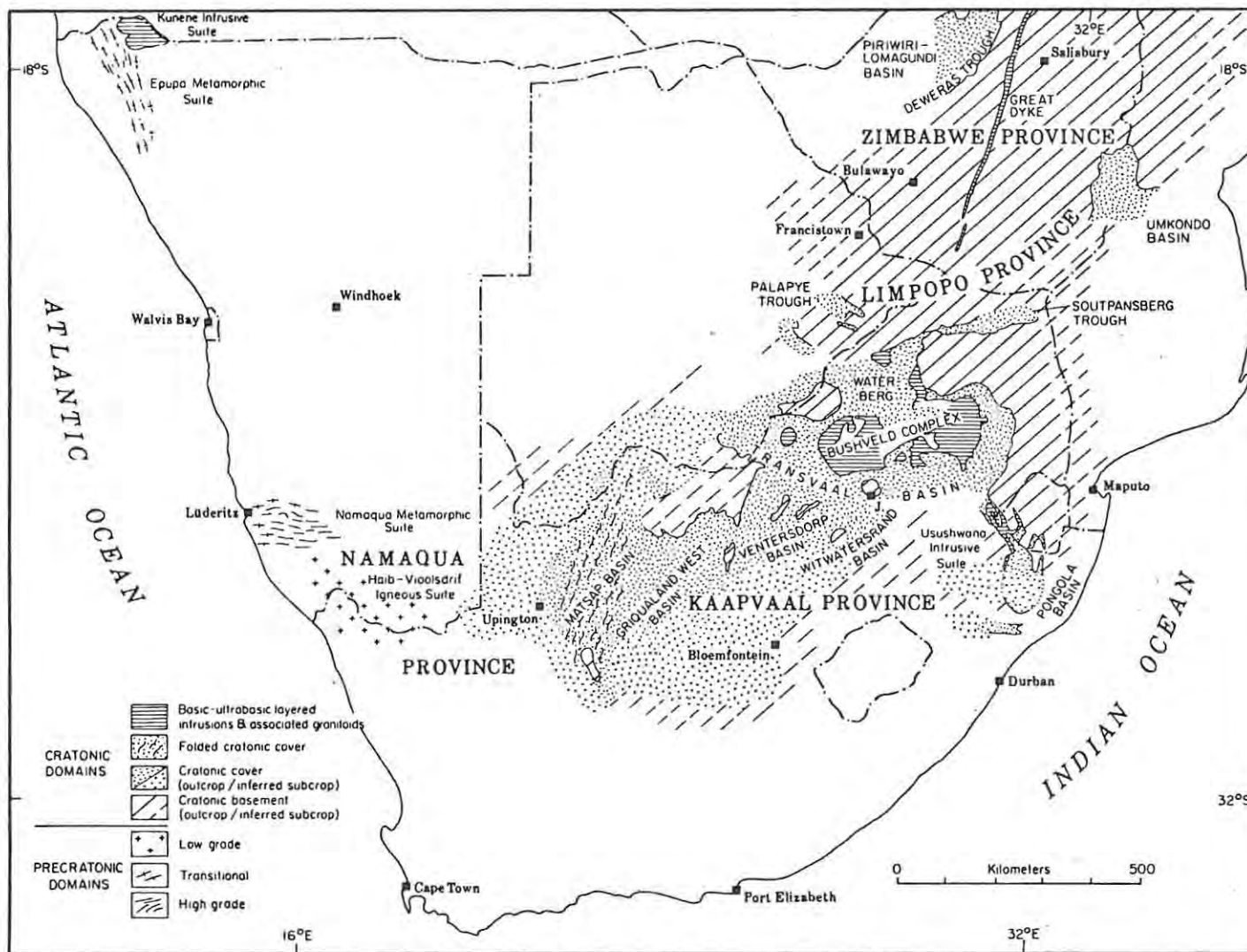


Figure 1.1 Nonpalinspastic map showing Late Archaean and Early Proterozoic supracrustal development from 2900 Ma to 1800 Ma (from Tankard et al., 1982)

Figure 1.2 Distribution of the Dominion, West Rand and Ventersdorp sequences (from Tankard et al., 1982)

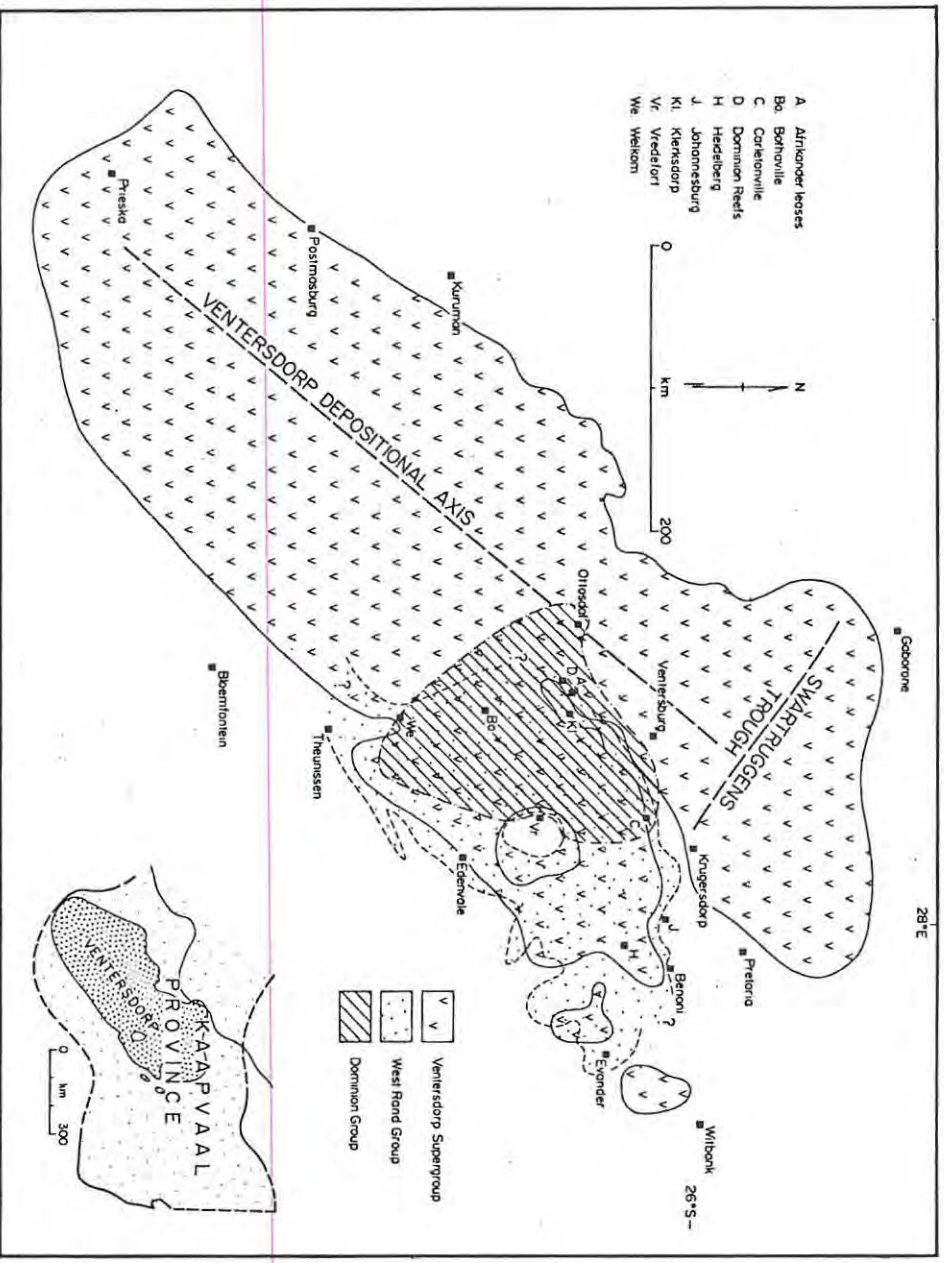


Figure 1.3 Distribution of outcrops of the Dominion Group (from SACS, 1980)

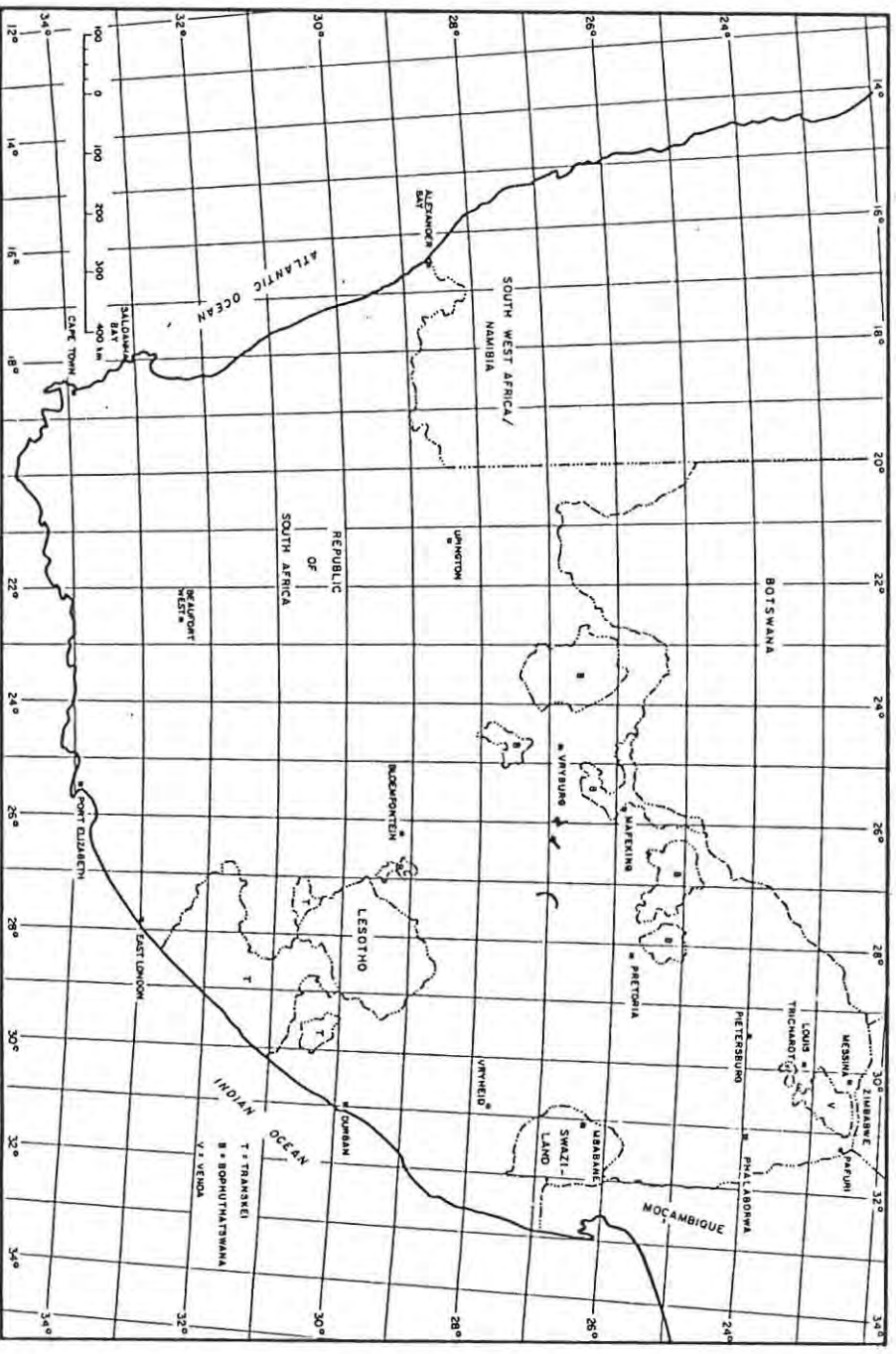


Figure 1.4 Distribution of outcrops of the Witwatersrand Supergroup (from SACS, 1980)

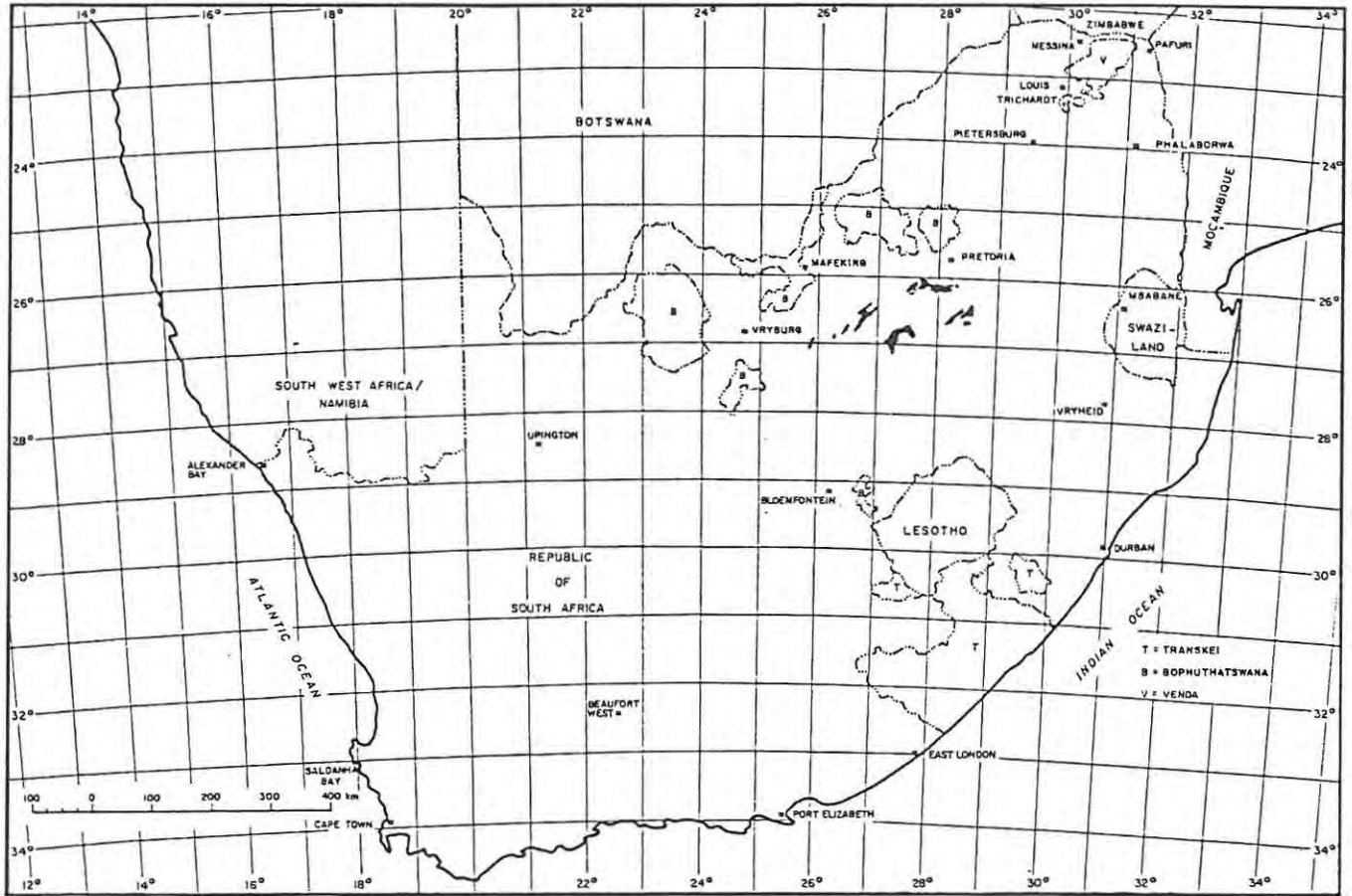
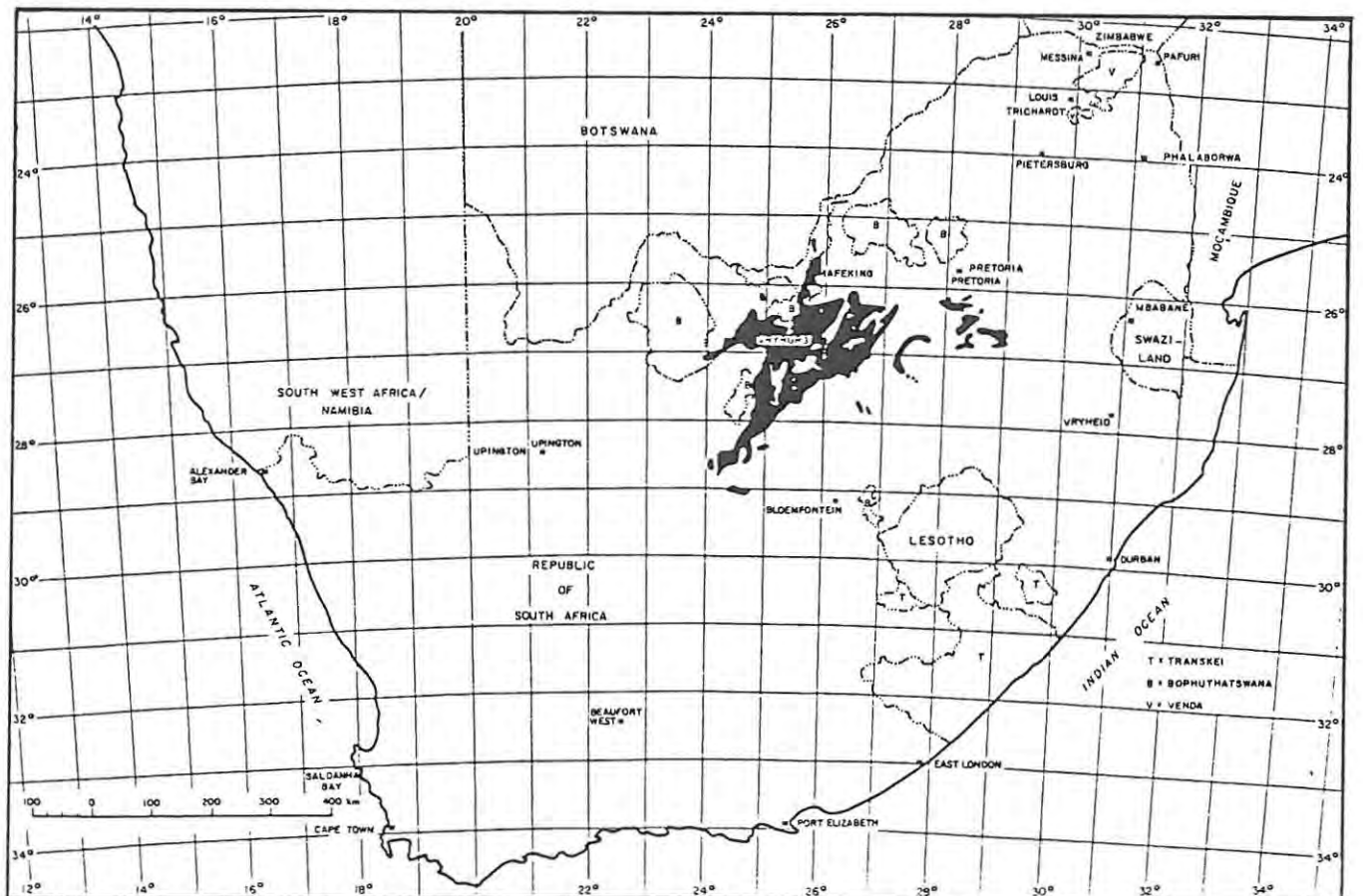


Figure 1.5 Distribution of outcrops of the Ventersdorp Supergroup (from SACS, 1980)



- (2) the Rhenosterhoek Andesite Formation consisting of intermediate lavas and tuffs; and
- (3) the Syferfontein Porphyry Formation, consisting of porphyritic felsic lavas, with interbedded tuffs and cherty amygdaloidal volcanics.

Palaeocurrent data indicate that the basal sediments prograded in a westerly direction and a braided fluvial regime is postulated for their origin (Watchorn, 1980), while the increased proportion of lenticular volcanic units towards the top of the Rhenosterspruit Formation indicates a period of coeval volcanism and sedimentation.

M P Bowen (in prep.) has shown that the basic to intermediate lavas of the Dominion Group as a whole, regardless of the formation in which they occur, represent a continuous differentiation sequence, while the felsic porphyries form a geochemically discrete group within which very little compositional variation occurs. For this reason, the Dominion lavas will be referred to in this study as members of either of two strictly informal groups, namely, the "Dominion Group basic lavas" and the "Dominion Group porphyries".

A major investigation of the structure and stratigraphy of the Ventersdorp Supergroup was undertaken by Winter (1965b). With the aid of an exhaustive array of literature and over two hundred boreholes, Winter devised a stratigraphic column based on three primary subdivisions incorporating six formations, as shown in Table 1.1. He selected the Bothaville area as his type areas as it includes all the type boreholes, maximum thicknesses of most of the formations, and most of the geographical areas from which the formation names are derived.

A portion of Winter's thesis was later updated (Winter, 1976) and a lithostratigraphic subdivision of the Ventersdorp Supergroup proposed. The most important change was the elevation of the Klipriviersberg Formation, formerly the Langgeleven Formation, to group status and its division into six formations. Figure 1.7 shows the stratigraphic nomenclature and correlation currently used.

The Venterspost Conglomerate Formation, which contains the gold-bearing Ventersdorp contact reef, is not included in either the underlying

Table 1.1 Stratigraphic column of the Ventersdorp System in the Bothaville type area (from Winter, 1965b).

Primary Subdivision	Formation	Representative Rocks
Upper	Allanridge (A)	Dark green amygdaloidal lavas
	Bothaville (B)	Quartzites, conglomerates, subordinate dark shales
Middle	Rietgat (R)	Green-grey amygdaloidal and porphyritic lavas; minor intercalated tuffaceous, calcareous and cherty shales; impure limestones; quartzites; conglomerates
	Makwassie (M)	
	New Kameeldoorns (N)	Quartz and felspar porphyries, green-grey amygdaloidal lavas, minor intercalated sediments
Lower	Langgeleven or Lower Volcanic (L)	Conglomerates; quartzites; tuffaceous, calcareous and cherty shales, impure calcareous rocks
		Green-grey pyroclastics and amygdaloidal lavas

Figure 1.7 Stratigraphy of the Ventersdorp Supergroup in northern Orange Free State and west-central Transvaal (from Tankard et al., 1982).

Witwatersrand Supergroup, or the overlying Ventersdorp Supergroup. It is overlain conformably by the Westonaria Formation, which is the lowermost of the six formations within the Klipriviersberg Group and which is developed only in the deepest parts of the depositional basin (SACS, 1980). The overlying formations lap on towards the basin margins, and are, from the base, the Alberton Porphyry, Orkney, Jeannette Agglomerate, Lorraine and Edenville Formations. Certain distinctive characteristics enable these formations to be differentiated from one another (Winter, 1976), as shown in Table 1.2. Winter (1976) classified the group as "almost exclusively andesitic".

Table 1.2 Subdivision of the Klipriviersberg Group (from Winter, 1976).

Formations	Marker beds	Distinctive Characteristics	Contacts in type borehole, Langgeleven No. 1	Thickness	Correlative Areas (List Incomplete)
Edenville		Green chalcedony and clusters of milky quartz amygdales. 1.	1 827	568	Allanridge, Klerksdorp, S. East Rand
Lorraine	AZ 1 AZ 2	Variolitic and spherulitic structures surrounded by whitish groundmass hence "altered zones"	2 395 AZ 1 flow 2 395-2 401 AZ 2 flows 2 492-2 535	217	O.F.S., Bothaville, Klerksdorp, East Rand
Jeannette Agglomerate		Prominent thick agglomerate at top. Many agglomerate beds in contrast with above. 2.	2 612	97	O.F.S., Bothaville East Rand
Orkney	PZ 1 PZ 2	Purple-topped ash-flow tuffs. 3. 4.	2 709 PZ 1 flow 2 709-2 750 PZ 2 flow 2 945-2 965	480	O.F.S., Klerksdorp, Bothaville, Edenville
Alberton Porphyry	Porphyritic Marker	Porphyritic lavas including "porphyritic marker". 6.	3 189 Porphyritic marker flow 3 352-3 360	171	South Vaal, Ventersburg (?), Edenville, West Rand, Klipriviersberg, East Rand, Vredefort, Evander, Greylingstad
Westonaria		Dark green "talcose tuffs"	3 360 Absent		South Vaal, Westonaria Alberton, S.E. Rand, Greylingstad

1. Some variolitic zones noted by Winter [1965] from Allanridge area.
2. Some porphyritic lavas south of Orkney
3. Pienaar distinguishes PZ 1 and PZ 2 in Klerksdorp and O.F.S. areas from other purple tops by their round chlorite amygdales.
4. Pienaar discovered a zone of pink chalcedony amygdales in O.F.S. goldfield in this stratigraphic position.
5. At base, Pienaar recognised a flow containing characteristic almond-shaped amygdales.
6. The "Porphyritic Marker" contains phenocrysts of feldspar longer than 50 mm and thicker than 7 mm and consists of 1-4 flows of lava.

In the Klerksdorp area the Westonaria Formation is absent, so the Alberton Formation forms the basal unit. The Orkney Formation is present, but the Jeannette Agglomerate has not been recognised. The Lorraine and Edenville Formations will be shown in this study to be indistinguishable geochemically, and will therefore be treated as one unit.

The Platberg Group overlies the Klipriviersberg Group with an extensive unconformity. The basal Kameeldoorns Formation consists of clastic

wedges of immature sediments which accumulated in fault troughs when the bounding faults were active, and its distribution bears no relation to the depositional basin of the underlying Klipriviersberg Group. The Makwassie Quartz Porphyry Formation overlies the Kameeldoorns Formation without a break and consists of quartz porphyries with quartz-free porphyritic and non-porphyritic volcanic rocks and minor bodies of sediments. The overlying Rietgat Formation contains predominantly green-grey amygdaloidal porphyritic and non-porphyritic lavas, but also contains subordinate sediments and pyroclastic rocks. The basal Garfield Member is regarded as being transitional between the two latter formations (Winter, 1976), and may form a part of either.

In the Klerksdorp area, the Makwassie Formation consists of two discrete units - the usual quartz-feldspar porphyry above, underlain by a sequence of greenish-grey, crystalline, porphyritic lavas characterised by small, dark green and scattered larger feldspar phenocrysts. P.M. Strydom (pers.comm., 1980) has named this unit the "Goedgenoeg formation", restricting the use of the name "Makwassie Formation" for the actual quartz-feldspar porphyry. This nomenclature will be followed in the current study.

An unconformity separates the Platberg Group from the overlying Pniel sequence, consisting of the Bothaville Formation and Allanridge Andesite Formation. The former is a sedimentary unit of subgreywacke composition, comprising conglomerates, quartzites and shales, while the latter comprises alternating porphyritic and non-porphyritic lava flows (Winter, 1976).

The Ventersdorp Supergroup is overlain by the Transvaal Sequence, the basal unit of which is the Black Reef Quartzite Formation, while beneath the Ventersdorp Supergroup lies the Witwatersrand Supergroup, a predominantly argillaceous and arenaceous sequence. Formerly divided into the "Upper" and "Lower" divisions, now called the Central Rand and West Rand Groups respectively, the succession contains only two relatively thin volcanic horizons. The Bird amygdaloid marker in the Central Rand Group does not occur in the Klerksdorp area, and will not be included in the current study. The Crown Formation, formerly the Jeppestown Amygdaloid, however, is more widespread and will be considered briefly.

2. INTRODUCTION

2.1 Purpose of Study

The aim of this study was to analyse a large number of volcanic rocks of the Ventersdorp and Witwatersrand Supergroups and the Dominion Group from the Klerksdorp area in order to ascertain whether any geochemical criteria exist which may serve to discriminate between the various volcanic formations constituting these groups.

The need for such a study is borne out by the decision, as far back as 1961, by the Consultative Committee of the Economic Geology Research Unit at the University of the Witwatersrand, to investigate the problem of correlation within the Ventersdorp System. Their motivation was that:

"Such a study would be of considerable interest to the mining companies engaged in exploration where it was necessary to drill through thick developments of this System before penetrating into the underlying Witwatersrand rocks. In many instances a very considerable footage had to be drilled in the Ventersdorp rocks before a recognisable marker horizon could be identified and an estimate given of the depth at which the Witwatersrand rocks might occur beneath the Ventersdorp. If it were possible to identify the lavas and sediments, then time and effort could be saved where it was indicated that an excessively thick succession of Ventersdorp rocks was likely to be present below the horizon recognised. In short, holes could be stopped sooner where it became evident that the cover above Witwatersrand rocks placed the latter at unfavourable depths." (Winter, 1961).

The logical extension of this was to include the lavas of the Dominion Group, which are, in many areas, stratigraphically intimately related to those of the Ventersdorp Supergroup, and are sometimes visually indistinguishable, especially where boundary outcrops are scarce. Of the two volcanic horizons in the Witwatersrand Supergroup, only one,

the Crown Formation, more commonly known as the Jeppestown Amygdaloid, is present in the Klerksdorp area and is included briefly in this study for completeness' sake.

The EGRU project was conducted by reseachers such as Winter (1961, 1962, 1963a, 1964, 1965a) and Tyler (1979a,b), followed later by Watchorn (1980) with work on the Dominion Group. Their conclusions, however, were reached mainly by studying stratigraphic field relations and petrography. Until Wyatt (1976) presented his M.Sc. thesis, in which he provided 49 chemical analyses of rocks from the Klipriviersberg Group, no other rocks of the Ventersdorp Supergroup had been the subject of a systematic geochemical study. Only five chemical analyses of Dominion Group lavas have been published and these - two rhyolites and three andesites - were presented by Von Backström (1952, 1962).

Winter (1965b) stated that "the lavas of the Langgeleven Formation (now the Klipriviersberg Group) may be confused with flows occurring in and locally at the base of the Makwassie Formation. They may also be confused with similar volcanics of pre-Witwatersrand age where the latter do not show signs of a more intense deformation. They may easily be confused with flows of the Rietgat Formation, and have been confused with the Allanridge lavas. However, the Rietgat and Allanridge lavas are not interbedded with an abundance of agglomerates and tuffs, nor do they have an abundance of flows of the peculiar structure of amygdaloidal tuffs."

Referring to the Makwassie Formation, he stated that "the isopach map establishes that the Formation containing quartz porphyries is extensive, thick and continuous. This is an important conclusion considering the possible confusion which may arise through correlation of this phase of the Ventersdorp volcanics with the acid rocks of the Dominion Reef System." "The latter, however, are generally somewhat schistose and associated with other schistose rocks, they have a peculiar cherty appearance, being metamorphosed to a higher degree, and their phenocrysts do not appear to be as prominent as those of the Makwassie Formation." Winter (1976) added that "it may be difficult to distinguish between the schistose structures and possible flow (rhyolitic) and compactive structures."

Winter (1965b) also said that "the green-grey amygdaloidal lavas (of the Makwassie Formation) closely resemble the Langgeleven Formation and can only be distinguished from them if porphyry or sediments are intercalated with them. Others resemble flows of the Rietgat Formation. Establishment of the upper boundary becomes doubtful in areas where green-grey porphyritic lavas occur extensively in both formations."

It is therefore clear that, although visual criteria can and do exist whereby formations may be distinguished from one another, very often these criteria are not evident or may even be absent. The need for an alternative method of identification is thus patently obvious and it is hoped that the geochemical methods employed in this thesis, which have yielded encouraging results, will be implemented by others in order to validate and elaborate on results in and further afield from the current study area.

The preliminary work on this project was executed jointly by T.B. and M.P. Bowen. This included the collection of core samples, subsequent processing and chemical analysis. All further interpretation and manipulation of data was performed individually unless otherwise acknowledged. Petrographic descriptions and petrogenetic interpretations of these data are included in a dissertation by M.P. Bowen (in prep.).

2.2 Sampling Procedure

Samples were obtained from borehole cores drilled in a triangular area of 1 500km² roughly defined by the towns Hartbeesfontein-Harrisburg-Vierfontein (see Figure 2.1). Borehole core was made available by Anglo American Prospecting Services (Pty) Ltd of Klerksdorp. Table 2.1 lists the boreholes which were sampled together with the number of samples used from each hole:

Table 2.1 Number of samples analysed from each borehole

<u>Ventersdorp Supergroup</u>		<u>Dominion Group</u>		<u>Crown Formation</u>	
JBF-1	31	DRS-6	22	R-1	3
JHA-1	9	DSF-7	49	JY-8	<u>3</u>
SH-1	8	DSF-10	16	<u>Total</u>	<u>6</u>
VHD-1	15	DHF-8	5		
WS-4	24	DRL-13	8		
WS-5	101	DRH-14	3		
PK-10	6	DRH-15	<u>8</u>		
W-20	<u>15</u>	<u>Total</u>	<u>111</u>		
<u>Total</u>	<u>209</u>				

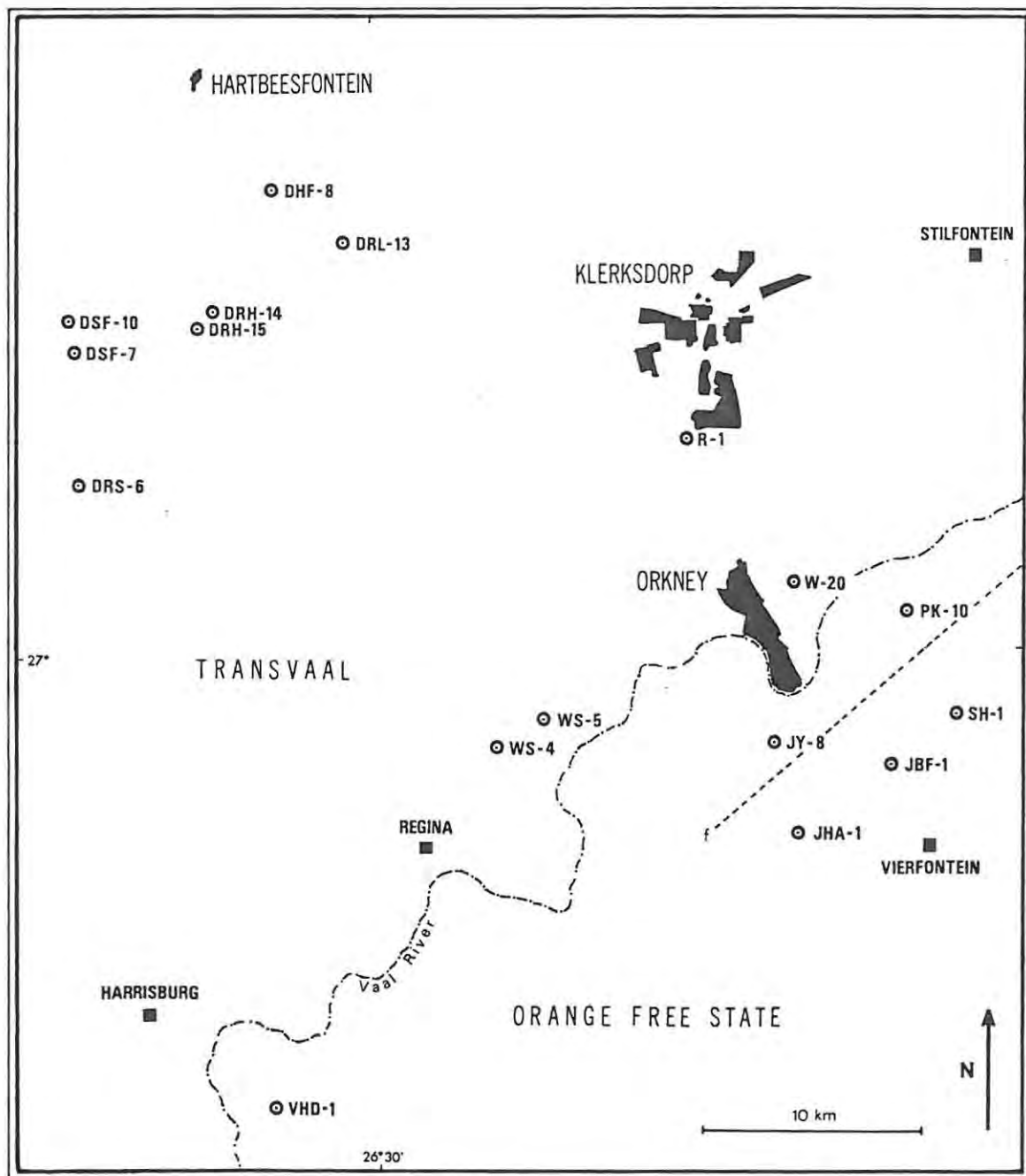
The localities of these boreholes are shown in Figure 2.1.

Those headed "Ventersdorp Supergroup" and "Crown Formation" are all situated to the south of Klerksdorp, four in the Transvaal and six across the Vaal River in the Orange Free State. WS-4, WS-5 and VHD-1 were drilled directly into Ventersdorp lavas. JBF-1, SH-1, PK-10 and W-20 had a cover of Transvaal Supergroup rocks ranging between 181 and 733m thick before the Ventersdorp lavas were reached, while the lavas in JHA-1 were capped by 165m of Karoo rocks. In the "Crown Formation" holes, the Jeppestown Amygdaloid or Crown Formation is found between 1688m and 1731.3m from the surface in R-1, and between 1754.1m and 1808.9m in JY-8. All the "Dominion Group" holes were drilled directly into Dominion Group lavas, with the exception of DRS-6, in which the lavas are capped by 86.26m of the Rietgat Formation of the Ventersdorp Supergroup.

Pemberton (1978) made a survey of major and trace element variations within individual Karoo basalt lava flows in order to determine the best position in which to sample a flow for chemical analysis. He concluded that flows should be sampled as close to their base as possible, because :

1. The major element abundances for the basal and topmost samples of each flow tend to be very similar but within the interior the composition varies erratically.
2. The base of a flow is easy to recognise and locate in the field.

Figure 2.1 Borehole locality map



The latter reason is especially relevant to flows in the Klipriviersberg Group, many of which have "swirled tops", (flow or brecciated tops, consisting of 0,3 to 6 metres of small, sub-rounded blocks of highly amygdaloidal lava (Winter, 1965b)) making the exact boundary difficult to determine in core samples. With a brecciated lava top the possibility of infiltration of the next flow down between the lava blocks of the underlying flow must be considered as well as alteration due to penetration of moisture into brecciated flow tops. A further phenomenon to avoid is the entrainment of clasts into the sole of an advancing lava flow as it moves across a rubble-strewn surface.

An alternative method is to take a number of samples throughout a flow and average the results to obtain a representative analysis. This was not suitable for the current study, however, as the average distance between samples in the boreholes in which the most detailed sampling was done, was about 22m. Within-flow sampling would have increased the number of samples to unmanageable numbers for a project, the aim of which was to determine inter- and intraformational variations. Consequently the samples for the current project were taken as close to the base of flows as possible, but above the basal amygdaloidal zone.

Pemberton (1978) found that the removal of amygdales from samples leads to a slight difference in element content from those with amygdales included. Although he concluded that the overall shift in concentration due to the presence of amygdales is a relatively minor one in comparison with the within-flow variations, highly amygdaloidal zones were nevertheless carefully avoided during sampling in the current study.

Boreholes WS-5 and DSF-7 were selected by P.M. Strydom as holes containing the most complete and representative sequences of the Ventersdorp Supergroup and the Dominion Group respectively in the study area. These two boreholes were therefore sampled in detail at intervals of about 22m, while the other boreholes were sampled at much wider but roughly equal intervals, in order to obtain information on possible lateral variations within different formations, and to facilitate correlation on a broad scale across the study area. WS-5, however, contains only 22.4m of the Makwassie Formation, so WS-4 which contains a thickness of 367m of this Formation, was sampled in detail.

Within WS-5 and DSF-7 samples were usually taken from between two and several flows apart, with very few being taken from adjacent flows. The thickness of flows can vary between a few metres and several tens of metres or even more, but is usually from about 5 to 25 metres.

A total of 504 samples was collected from the Anglo American core yards at Vaal Reefs, Vaal Reefs South, and Klerksdorp. Of these, 326 were eventually processed for chemical analysis. The remaining 174 were omitted due to time and manpower constraints. The positions of the samples within each borehole are indicated in Figure 2.2, which shows condensed lithostratigraphic logs adapted from those provided by Mr P.M. Strydom of Anglo American Prospecting Services, Orkney. A breakdown of the number of samples analysed from each formation is set out in Table 2.2.

Table 2.2 Number of samples analysed from each formation

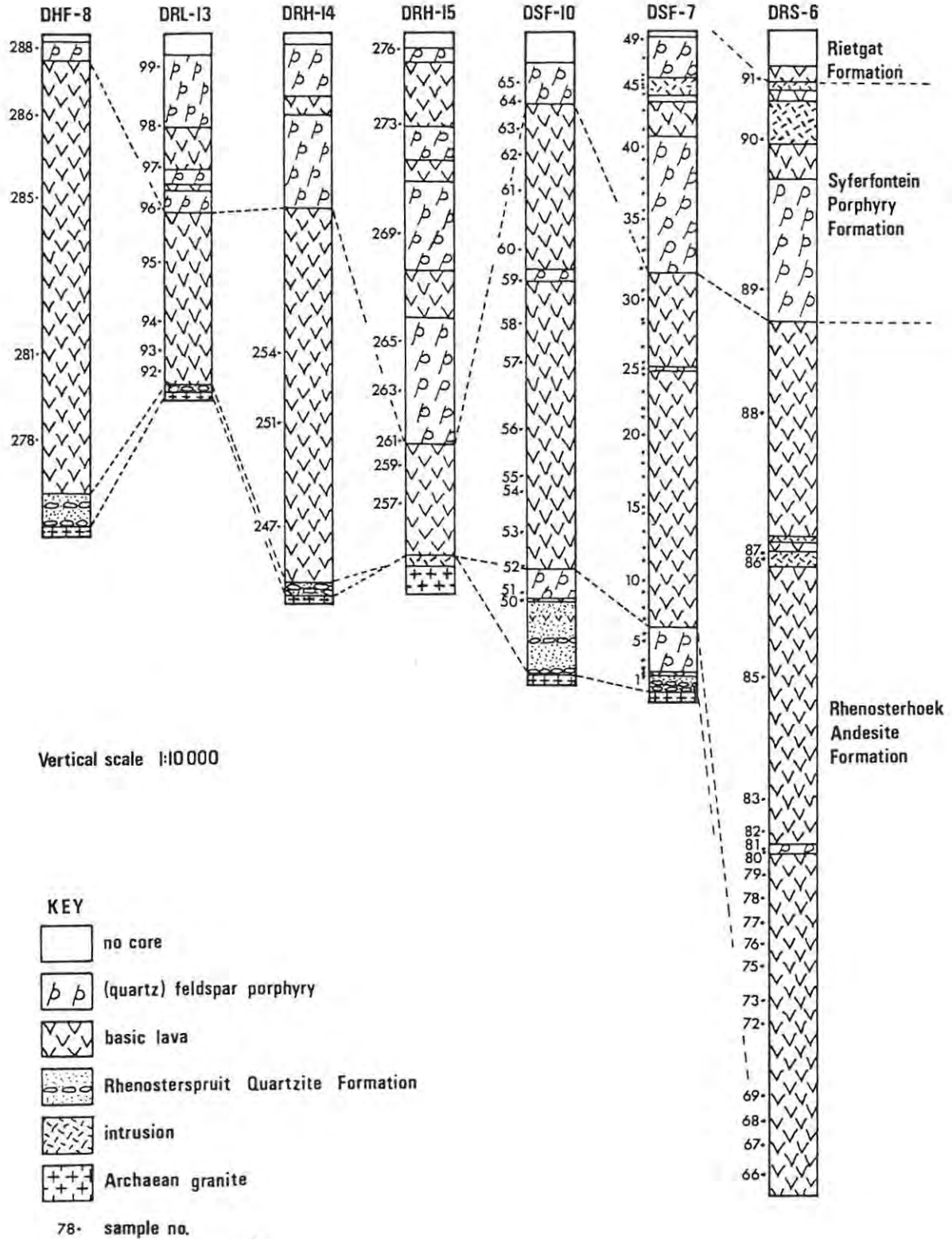
<u>FORMATION</u>	<u>TOTAL</u>
Allanridge	35
Rietgat	34
Makwassie	17
Goedgenoeg	35
Lorraine-Edenville	46
Orkney	32
Alberton	19
Crown	6
Dominion Group acid porphyries	32
Dominion Group basic lavas	<u>70</u>
	<u>326</u>

2.3 Sample preparation and chemical analysis

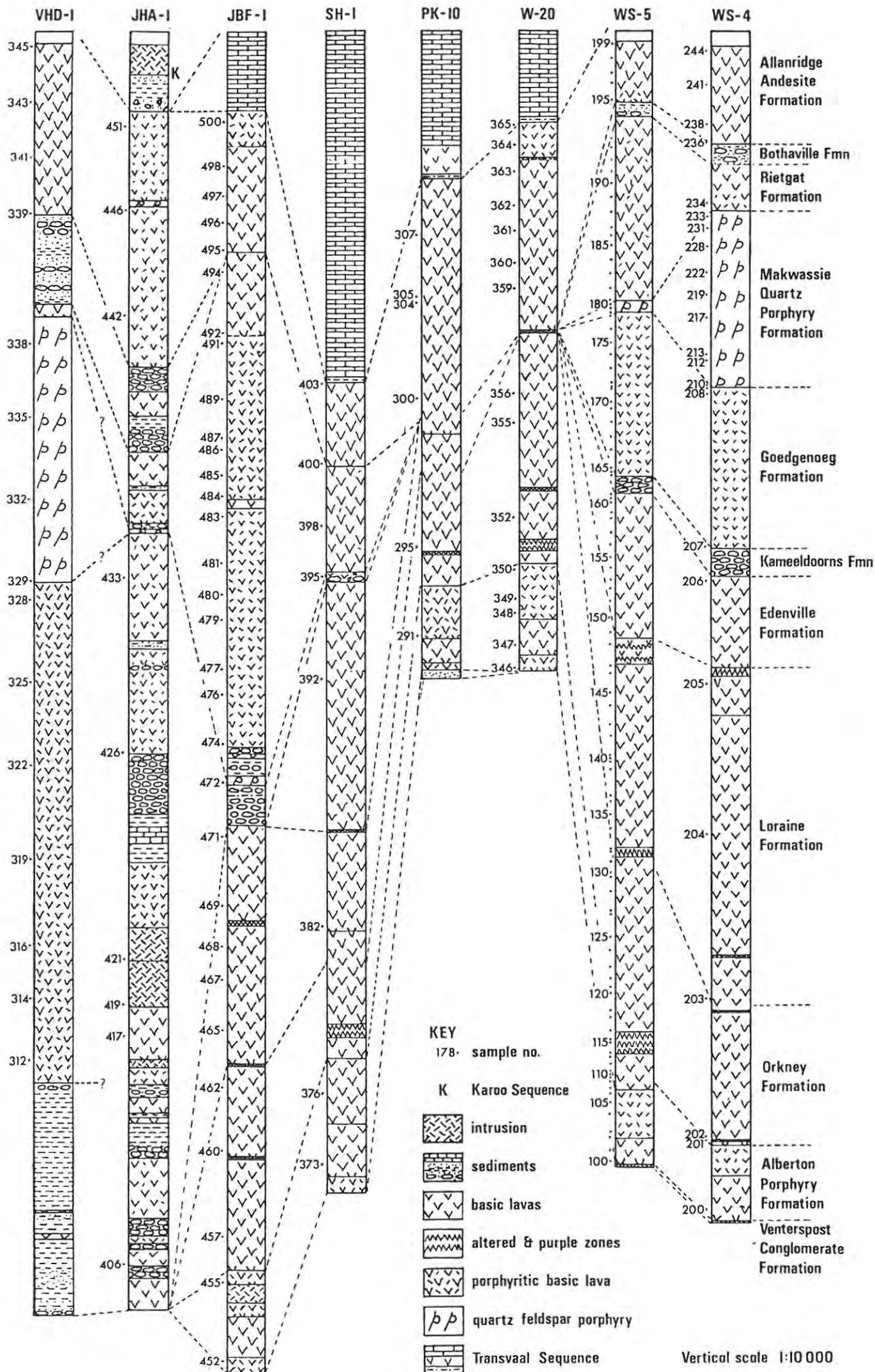
Each sample was prepared for chemical analysis by breaking up the sample in a rock splitter and removing the outer surface of the core by grinding, as this surface was frequently weathered and may have been

Figure 2.2 Condensed lithostratigraphic logs with sample positions

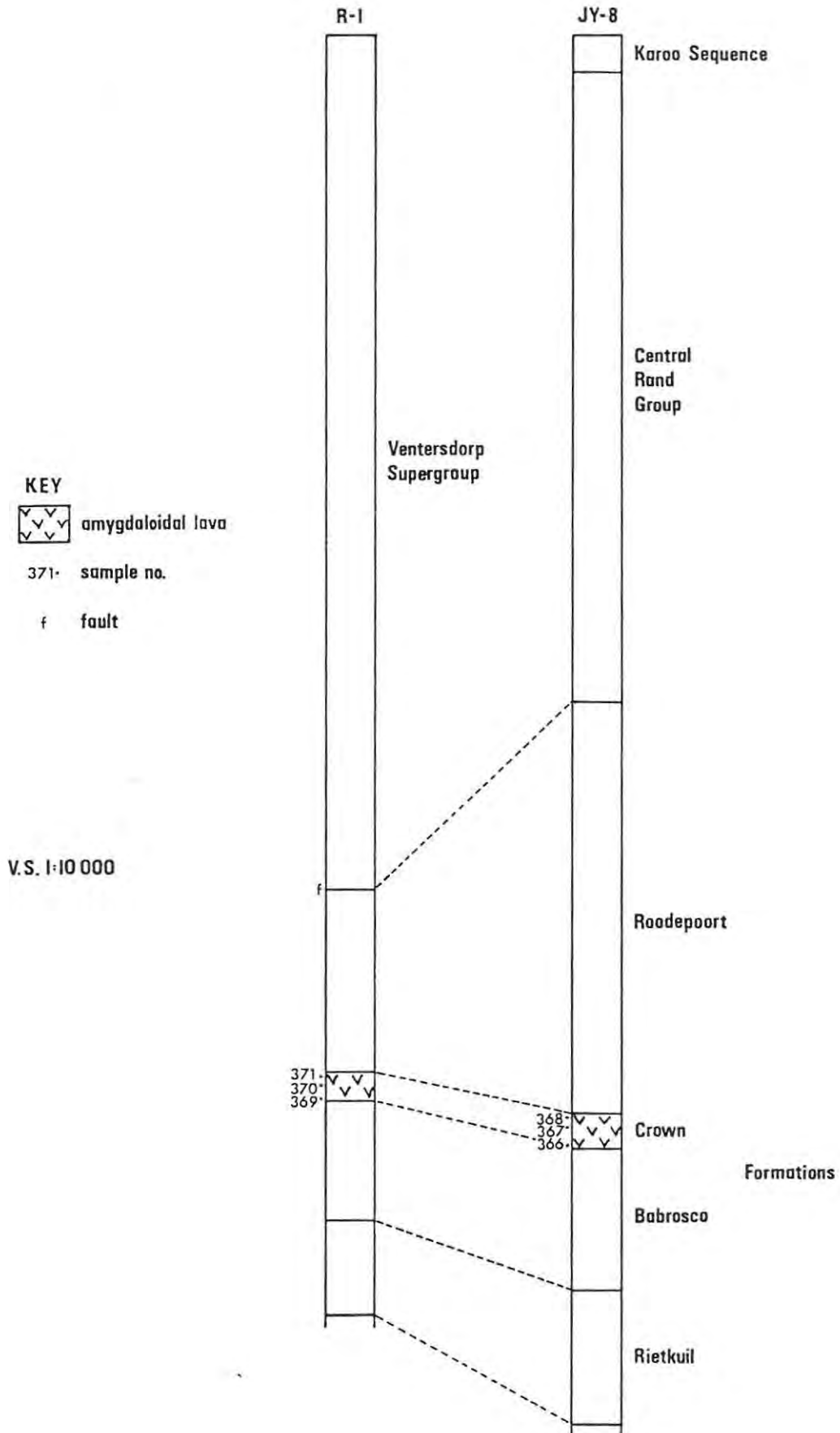
(a) Dominion Group boreholes



(b) Ventersdorp Supergroup boreholes



(c) Crown Formation boreholes



contaminated by metal smears from the core barrels and core trays. Any vein-filling material or amygdalae were also removed at this stage. The rock fragments were then passed through a jaw crusher, and ground to approximately 200 mesh in a Hertzog swing mill, using a Mn-steel container. The final crush was quartered and seven grams were finely ground by hand, using an agate pestle and mortar.

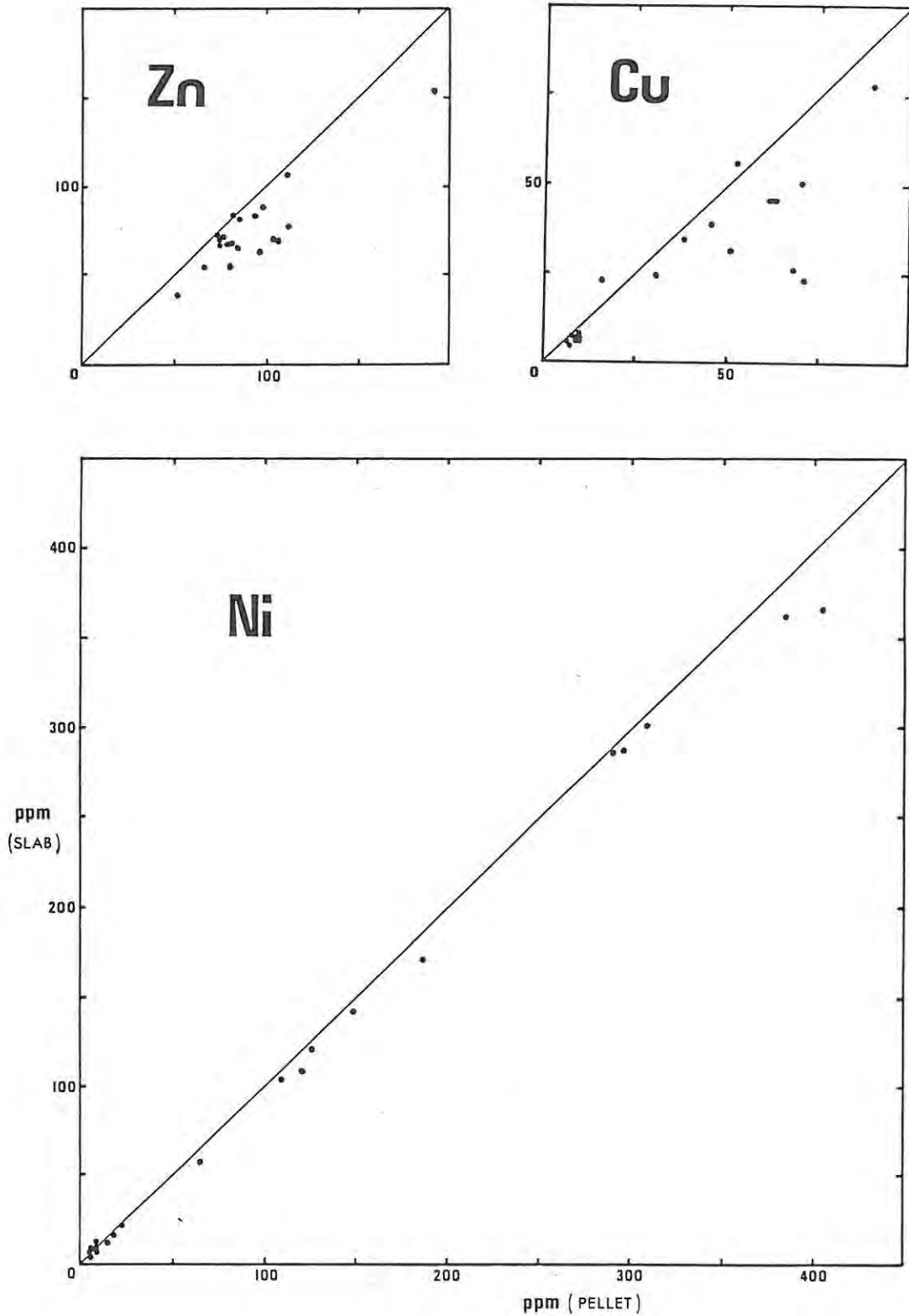
Of the seven grams, two were used for making a Norrish fusion disc for major element analysis, according to the method described in Norrish and Hutton (1969). This entailed mixing 1.5 grams of Spectroflux with 0.28 grams of sample which had been ashed at 1000°C to eliminate all volatile material. The mixture was melted in a platinum crucible at more than 1000°C and poured into a graphite mould to form a thin glass disc.

The remaining five grams were used for making a briquette of sample contained in a boric acid/bakelite case, which was compressed under 15 tons/sq. in. using a steel barrel and plunger. The briquettes were used for the analysis of Na₂O and the trace elements Ba, Nb, Zr, Y, Sr, Rb, Zn, Cu, Ni, Co, Cr, V, La, Ce and Nd.

An alternative to the briquette method - the use of polished slabs - was tested. A disc was simply sawn from the borehole core, ground flat and polished on one side, and inserted in the spectrometer for trace element analysis. A comparison between the two methods for a zinc-copper-nickel run is shown in Figure 2.3, in which concentrations determined on briquettes are plotted against concentrations obtained from the polished slabs. Nickel shows a far better correlation than either zinc or copper, but in general, the slab method resulted in lower values than those obtained from the pressed pellets. This discrepancy is to be expected, as a slab will very rarely display the same degree of homogeneity as a pressed pellet, unless the rock is extremely fine-grained.

Although this technique has not been pursued further, it may prove in practice that the time saved in using the slab method more than compensates for the loss of precision. This method might be viable for monitoring drilling, when decisions as to whether a particular borehole should be continued or abandoned cannot rest on elaborate and lengthy laboratory procedures.

Figure 2.3 Concentrations of Zn, Cu and Ni determined on slabs vs pellets



Major and trace element analyses were performed by X-ray fluorescence, using a Philips PW 1410 X-ray spectrometer. Both international and locally calibrated standards were used for calibration. International standards used include the following: AGV-1, BCR-1, G-2, GSP-1, JG-1, and PCC-1, while Rhodes University in-house standards used are CAR-08, KRF-13 and PRC-1. A number of standards calibrated at N.I.M. (now Mintek) were also used, viz. NIM-G, NIM-N, NIM-P, S-9, S-10, S-12 and S-15. Details of analytical techniques can be found in Appendix II.

Data reduction was performed initially on an ICL 1904S computer and subsequently on a CDC Cyber 170-825, using programs compiled at the Geochemistry Department, University of Cape Town, and modified for use at Rhodes University by Dr J.S. Marsh of the Rhodes Geology Department.

Major and trace element analyses for 326 samples are listed in Table 2.4. All iron was analysed as Fe_2O_3 . Both major and trace element values have been recalculated to 100%, volatile-free. It was considered justified and indeed necessary to recalculate trace element as well as major element data, in the light of the high degree of hydration and carbonation of a number of the samples. Five samples had a volatile content of over 10%, the highest being 12.88% for sample AR-496. In cases such as these and many others, if recalculation of trace elements was not effected, major and trace element contents might prove irreconcilable. The original total ("TCTAL") and volatile content ("LCI") are provided for each sample in Table 2.4, should back-calculation to the original values be required.

The numerical part of the sample names indicates the order in which samples were taken, while the alphabetic prefix indicates the group and the formation to which each sample belongs, as shown in Table 2.3.

Major element values are expressed in weight percent, trace elements in parts per million. Trace elements listed as 0 ppm had values below the lower limit of determination (LLD). Those marked with an asterisk (*) had values below the LLD, but above the detection limit (DL), which is half the LLD (see Table 2.5).

Table 2.3 Derivation of sample names

GROUP	FORMATION	SYMBOL
	AR - Allanridge	AR
P - Platberg	R - Rietgat M - Makwassie G - Goedgenoeg	PR PM PG
K - Klipriviersberg	L - Loraine-Edenville O - Orkney A - Alberton	KL KO KA
J - Jeppestown	C - Crown	JC
D - Dominion	P - Porphyries B - Basic Lavas	DP DB
Unknown	IN - Intrusive	IN

In X-ray fluorescence analysis the precision of any determination may be considered to be a function of:

- (i) the reproducibility of the sample preparation techniques, and
- (ii) X-ray counting statistics.

The successful duplication of 45 analyses has demonstrated that the conditions required under (i) are sufficiently satisfied. X-ray counting statistics, expressed in terms of average counting error (CE) and lower limit of determination (LLD) for each batch of trace element analyses processed, are shown in Table 2.5, giving a measure of the precision attained during analysis. The samples were analysed in three separate batches consisting of 40, 48 and 238 samples respectively, with the exception of the rare earths La, Ce and Nd, which were analysed in two batches of 40 and 286 samples respectively.

Table 2.4 Major and trace element analyses

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE:	DB- 1	DB- 2	DB- 3	DP- 4	DP- 5	DP- 6	DB- 7	DB- 8	DB- 9	DB- 10
SiO2	57.76	58.44	58.78	68.32	70.84	69.87	52.21	52.19	53.77	57.55
TiO2	1.084	1.041	1.004	0.849	0.809	0.830	0.883	0.889	0.843	1.171
Al2O3	14.58	13.85	13.66	12.74	12.50	12.65	13.29	13.08	13.09	14.46
Fe2O3	11.35	11.48	11.72	7.35	5.66	6.71	12.87	13.68	11.48	10.05
MnO	0.25	0.15	0.14	0.11	0.08	0.07	0.18	0.23	0.17	0.15
MgO	3.51	3.54	3.77	2.50	1.73	1.70	7.71	10.07	9.52	4.94
CaO	7.35	6.82	6.77	2.05	1.45	2.20	10.42	6.68	7.36	5.57
Na2O	5.33	3.03	3.17	3.65	3.42	2.70	1.67	2.85	3.43	5.11
K2O	0.50	1.37	1.34	2.16	3.22	3.03	0.08	0.11	0.15	0.57
P2O5	0.255	0.265	0.257	0.271	0.278	0.243	0.191	0.194	0.180	0.421
TOTAL:	99.21	99.86	99.75	99.69	99.49	99.11	99.52	99.30	100.12	99.55
LOI	5.72	1.65	1.48	1.52	1.68	2.33	6.89	2.62	2.13	1.71
BA	204	395	484	846	853	889	25	48	77	244
NB	8.5	7.5	7.1	12.9	13.4	15.0	4.2	2.4	5.3	7.7
ZR	173	162	171	256	277	275	91	89	93	174
Y	30.1	26.4	27.7	28.9	27.3	27.1	22.7	25.0	22.7	30.2
SR	286	460	465	214	111	71	386	132	165	147
RB	11.4	51.0	44.5	36.9	81.5	114.7	1.3	2.2	1.4	9.8
ZN	125	100	103	41	87	82	107	126	99	95
CU	28	16	42	10	9	7	58	60	43	53
NI	19	23	25	12	9	12	310	446	322	125
CO	35	54	55	26	16	21	68	87	71	58
CR	3	6	12	12	2*	4	1344	1510	1290	34
V	27	26	245	106	53	87	236	257	221	230
LA	48	25	27	44	46	35	10	13	7*	21
CE	28	59	50	77	94	76	29	27	15	45
ND	28	24	28	37	40	34	17	19	11	25

SAMPLE:	DB- 11	DB- 12	DB- 13	DB- 14	DB- 15	DB- 16	DB- 17	DB- 18	DB- 19	DB- 20
SiO2	54.10	61.12	57.18	60.41	53.72	52.45	50.82	52.78	52.70	52.51
TiO2	1.323	1.684	2.076	1.596	0.787	0.758	0.744	0.682	0.660	0.641
Al2O3	15.34	11.41	13.16	12.98	14.26	13.52	13.93	14.45	16.37	14.55
Fe2O3	12.69	11.45	14.78	10.94	11.32	11.57	13.93	10.48	9.78	10.43
MnO	0.16	0.10	0.18	0.16	0.14	0.16	0.18	0.18	0.14	0.13
MgO	5.21	1.48	2.45	2.99	7.26	8.86	7.83	8.27	7.36	8.53
CaO	5.56	10.67	4.17	6.03	8.06	9.44	11.13	9.62	8.75	9.52
Na2O	4.76	0.29	3.74	3.63	3.18	2.73	1.44	2.69	3.94	2.96
K2O	0.44	1.16	1.51	0.73	1.08	0.64	0.19	0.69	0.65	0.62
P2O5	0.397	0.637	0.778	0.546	0.189	0.168	0.168	0.163	0.152	0.109
TOTAL:	99.74	100.03	98.97	99.08	99.31	99.99	99.97	99.99	99.32	99.86
LOI	1.77	1.89	2.72	2.62	1.91	2.02	3.69	2.94	2.35	2.53
BA	149	549	557	176	676	406	138	278	427	213
NB	9.3	12.5	16.5	12.9	5.6	2.9	2.8	2.6	4.5	3.4
ZR	196	260	351	259	105	90	89	80	95	82
Y	33.2	44.6	34.4	41.2	22.6	21.3	23.2	17.7	20.5	18.4
SR	121	2015	179	247	462	444	523	435	564	484
RB	7.9	37.9	90.7	28.4	17.2	11.0	3.9	13.9	14.1	15.8
ZN	123	93	142	116	104	90	96	83	85	77
CU	36	34	12	32	60	55	85	70	66	67
NI	155	19	28	67	262	290	360	318	240	258
CO	67	43	53	48	69	66	75	65	58	62
CR	41	4	17	7	772	926	1197	784	548	850
V	243	291	326	270	208	209	217	194	186	191
LA	20	34	43	31	12	13	8*	8	9	12
CE	50	79	88	86	25	27	22	25	18	20
ND	29	45	52	45	18	17	16	8	14	4

SAMPLE:	DB- 21	DB- 22	DB- 23	DB- 24	DP- 25	DB- 26	DB- 27	AR- 28	DB- 29	DB- 30
SiO2	56.13	57.20	54.96	62.01	70.63	56.82	53.75	57.36	58.35	57.48
TiO2	0.624	0.644	0.715	1.430	0.795	0.625	1.251	1.413	1.541	1.134
Al2O3	15.26	15.50	16.11	12.75	12.53	15.94	14.79	15.51	13.99	14.17
Fe2O3	9.47	9.19	9.95	12.30	6.28	7.95	13.58	16.43	15.14	14.19
MnO	0.11	0.12	0.14	0.16	0.07	0.12	0.17	0.15	0.13	0.16
MgO	5.59	5.18	5.75	1.94	1.06	5.27	3.67	2.80	2.85	3.24
CaO	8.00	7.66	7.41	4.89	3.60	7.29	8.90	2.44	4.21	5.60
Na2O	3.62	3.49	3.74	2.85	2.89	3.50	3.01	2.90	2.61	2.94
K2O	1.04	0.84	1.04	1.20	1.88	2.00	0.39	0.83	0.64	0.66
P2O5	0.165	0.171	0.195	0.465	0.255	0.177	0.247	0.171	0.458	0.314
TOTAL:	99.58	99.72	99.75	98.55	98.99	99.16	99.11	97.87	99.13	98.54
LOI	2.05	2.09	1.96	2.72	1.29	2.08	3.08	3.14	3.32	3.25
BA	341	327	504	373	637	567	306	341	423	214
NB	4.7	4.5	4.7	11.0	13.5	3.9	8.5	11.3	10.4	6.9
ZR	117	121	133	258	297	106	202	230	237	158
Y	19.7	20.7	22.7	35.5	30.8	18.9	33.0	25.9	38.2	27.4
SR	503	530	628	406	401	593	423	134	413	280
RB	22.2	17.0	20.2	42.1	37.3	52.3	14.8	15.6	21.4	24.7
ZN	78	78	88	73	69	77	117	146	149	146
CU	75	66	52	23	12	73	56	23	33	57
NI	90	94	98	14	12	131	35	27	21	33
CO	50	49	52	48	18	54	52	68	67	63
CR	53	51	49	4	9	52	0	0	10	10
V	161	164	169	285	48	166	223	259	435	324
LA	13	15	17	42	45	14	21	18	32	23
CE	26	29	31	76	85	35	36	72	72	42
ND	19	16	18	45	38	16	30	34	43	27

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE:	DB- 31	DP- 32	DP- 33	DP- 34	DP- 35	DP- 36	DP- 37	DP- 38	DP- 39	DP- 40
SI02	55.68	67.18	58.44	64.02	71.41	71.65	70.04	71.04	71.61	76.74
TIO2	1.085	0.891	1.489	0.894	0.774	0.819	0.820	0.809	0.788	0.747
AL2O3	14.52	13.25	21.63	14.70	11.73	12.99	12.88	12.44	12.90	12.00
FE2O3	15.33	6.62	6.58	12.54	9.19	5.41	6.26	5.81	7.13	1.81
MNO	0.16	0.08	0.03	0.11	0.08	0.07	0.06	0.12	0.06	0.02
MGO	3.77	1.14	1.40	2.17	1.35	0.40	0.60	0.48	0.74	0.32
CAO	7.06	4.15	1.70	1.14	0.67	1.24	1.85	2.23	0.90	0.51
NA2O	5.33	4.03	1.29	1.01	1.85	4.12	4.16	3.48	4.26	3.45
K2O	0.77	2.35	7.12	3.10	2.69	3.03	3.07	3.31	1.36	4.14
P2O5	0.300	0.292	0.517	0.319	0.255	0.279	0.267	0.274	0.246	0.259
TOTAL:	98.90	98.50	98.91	99.87	99.44	99.93	99.37	97.89	99.42	99.05
LOI	3.11	3.60	3.02	3.06	2.12	1.46	1.98	2.54	1.62	0.80
BA	388	805	1121	390	430	1042	769	861	311	1207
NB	5.6	15.1	28.5	14.7	11.7	13.3	14.1	13.5	14.4	13.5
ZR	153	313	545	315	279	290	296	279	290	271
Y	29.2	30.2	60.7	33.8	27.5	29.2	30.6	28.4	31.7	24.6
SR	691	262	66	51	94	233	201	164	87	108
RB	28.6	40.8	204.4	90.9	90.4	68.2	56.1	74.5	37.5	112.0
ZN	116	86	46	125	183	58	65	83	127	51
CU	74	8	8	8	20	11	11	10	7	7
NI	38	7	16	10	11	10	9	7	10	6
CO	97	19	15	24	28	14	16	15	20	9
CR	7	0	3	2*	4	4	4	U	8	0
V	325	59	97	65	54	54	52	53	55	41
LA	25	72	21	118	40	50	52	46	48	49
CE	50	124	59	194	72	93	100	92	91	100
ND	28	54	36	77	34	40	44	38	41	45

SAMPLE:	DB- 41	KL- 42	KL- 43	KL- 44	KL- 45	KL- 46	DP- 47	DP- 48	DP- 49	DB- 50
SI02	57.78	57.72	50.28	50.53	50.86	50.91	69.37	67.68	68.38	62.50
TIO2	1.324	0.539	0.459	0.423	0.433	0.462	0.749	0.766	0.794	0.945
AL2O3	13.43	17.38	14.40	14.37	15.01	14.85	12.65	12.83	12.76	14.67
FE2O3	15.09	12.72	10.27	9.56	9.51	10.10	7.33	8.16	8.74	8.65
MNO	0.14	0.16	0.14	0.14	0.15	0.15	0.09	0.11	0.18	0.16
MGO	3.02	5.03	10.58	10.19	9.67	9.06	0.70	0.50	0.26	1.94
CAO	5.80	0.79	11.07	12.57	12.19	12.05	2.02	2.87	1.90	6.47
NA2O	2.72	3.95	1.02	1.54	1.43	1.58	4.01	4.39	3.33	3.14
K2O	0.55	1.66	1.77	0.63	0.70	0.79	2.67	2.48	3.41	1.16
P2O5	0.148	0.045	0.059	0.035	0.044	0.041	0.211	0.208	0.234	0.251
TOTAL:	99.20	99.39	99.85	100.34	101.46	99.46	99.68	99.47	98.63	98.59
LOI	5.90	3.65	2.67	2.54	2.62	2.49	2.16	2.86	2.39	4.99
BA	226	214	315	108	111	130	628	492	669	788
NB	7.7	2.4	0.7	0.4	0.5	1.2	14.8	15.7	14.5	9.5
ZR	144	42	34	32	33	37	246	248	256	187
Y	28.5	6.4	17.9	16.7	16.5	17.9	35.6	36.7	32.9	29.8
SR	287	62	209	192	173	182	237	346	165	174
RB	15.9	46.4	63.0	27.6	31.0	35.1	87.8	90.1	126.8	44.4
ZN	133	118	85	59	61	73	94	116	82	76
CU	49	155	85	75	73	87	13	113	9	64
NI	19	143	246	220	192	171	12	11	9	21
CO	64	60	59	55	55	58	19	21	22	40
CR	4	164	814	1119	755	465	5	4	0	2*
V	479	237	229	225	214	234	17	20	25	173
LA	23	0	0	5*	0	0	43	42	46	25
CE	38	0	0	0	0	0	82	83	92	51
ND	20	0	4*	0	0	15	39	38	35	26

SAMPLE:	DP- 51	DB- 52	DB- 53	DB- 54	DB- 55	DB- 56	DB- 57	DB- 58	DP- 59	DB- 60
SI02	69.84	51.82	55.56	57.93	53.89	55.50	53.14	58.31	70.87	60.28
TIO2	0.785	0.806	1.152	1.669	0.752	0.664	0.647	0.970	0.771	0.990
AL2O3	12.65	12.68	14.36	12.93	13.68	14.93	14.93	14.40	12.31	13.54
FE2O3	6.23	12.51	11.65	12.21	10.22	8.96	9.60	10.28	5.57	10.42
MNO	0.08	0.21	0.15	0.14	0.16	0.14	0.16	0.14	0.06	0.13
MGO	0.97	11.36	4.67	2.18	7.54	6.17	7.56	3.52	1.21	3.45
CAO	2.12	7.57	7.25	7.07	10.25	9.13	9.74	3.22	1.85	7.70
NA2O	4.40	2.51	3.99	3.54	2.61	3.64	2.77	2.88	1.43	1.96
K2O	2.64	0.35	0.85	1.71	0.73	0.70	1.32	0.97	5.68	1.20
P2O5	0.265	0.184	0.565	0.619	0.170	0.144	0.128	0.315	0.260	0.320
TOTAL:	99.49	99.66	99.79	98.50	100.42	100.55	99.81	99.83	99.33	100.86
LOI	0.80	2.63	1.24	3.61	2.59	1.41	4.31	1.73	1.26	1.76
BA	589	190	271	856	569	372	542	552	1695	1011
NB	12.4	3.1	7.5	13.0	85.1	3.7	3.2	6.2	11.3	6.6
ZR	272	76	167	270	85	70	82	163	259	167
Y	29.0	18.8	30.8	41.3	10.4	18.4	19.7	25.2	25.3	26.5
SR	158	244	419	402	574	685	420	743	254	704
RB	83.9	8.8	23.4	65.3	11.2	11.0	22.9	14.5	107.1	18.5
ZN	86	168	105	122	84	74	77	91	112	89
CU	12	52	39	39	80	68	64	57	14	43
NI	10	364	86	54	513	182	242	46	12	27
CO	21	76	56	53	68	50	64	53	19	51
CR	5	1301	41	11	428	682	670	20	4	10
V	69	233	235	291	218	188	190	214	70	222
LA	47	9	20	38	8*	8	12	19	50	27
CE	45	18	50	77	17	20	21	44	91	42
ND	36	14	32	47	14	13	16	30	42	27

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE:	DB- 61	DB- 62	DB- 63	DP- 64	DP- 65	DB- 66	DB- 67	DB- 68	DB- 69	DB- 72
SiO2	56.49	58.47	55.05	71.19	70.82	56.32	57.38	56.39	60.67	60.33
TiO2	U.603	1.435	1.210	0.749	U.783	1.155	1.172	1.138	1.958	1.849
AL2O3	15.51	12.92	15.35	12.22	12.38	14.40	14.94	14.16	11.94	11.42
FE2O3	8.53	14.02	13.09	5.56	5.58	11.56	11.68	11.76	13.03	13.58
MNO	U.12	U.16	U.22	U.07	U.07	U.17	U.17	U.14	U.14	U.14
MGO	5.75	2.60	2.12	0.61	U.64	4.61	5.94	4.69	2.28	2.29
CAO	8.66	6.63	7.97	2.46	2.31	6.18	3.65	6.33	5.24	5.18
NA2O	2.96	2.34	3.96	3.83	4.35	4.06	3.43	3.98	2.89	2.63
K2O	1.21	1.01	0.69	3.06	2.80	1.19	1.27	1.07	1.14	2.11
P2O5	0.165	0.420	0.353	U.255	U.267	0.354	0.372	0.332	U.710	U.677
TOTAL:	100.50	100.14	98.71	100.17	99.50	98.40	97.99	99.73	99.32	99.99
LOI	1.78	3.05	4.68	1.66	1.74	1.67	2.93	1.54	2.76	2.06
BA	362	536	324	877	737	415	508	386	664	688
NB	3.8	9.2	7.6	11.9	13.3	7.8	8.4	8.0	14.8	13.9
ZR	95	214	165	267	271	173	174	168	316	285
Y	19.1	34.4	30.3	27.6	28.0	31.9	29.4	31.7	48.0	44.0
SR	750	514	715	299	300	437	239	485	375	264
RB	58.4	45.5	29.6	73.3	67.5	29.4	27.4	36.1	53.5	106.2
ZN	75	106	107	73	78	111	240	108	137	149
CU	86	34	65	15	11	79	70	78	26	18
NI	108	19	37	8	11	87	96	93	28	23
CO	54	58	72	17	16	56	62	57	49	57
CR	60	4	10	4	3	42	46	52	13	10
V	174	392	350	60	53	247	274	244	312	326
LA	16	33	22	52	50	20	24	18	40	37
CE	27	72	47	98	92	40	30	45	86	85
ND	18	39	28	44	43	29	32	27	50	52

SAMPLE:	DB- 73	DB- 75	DB- 76	DB- 77	DB- 78	DB- 79	DP- 80	DP- 81	DB- 82	DB- 83
SiO2	59.41	60.46	59.97	60.74	48.79	53.20	70.60	71.32	53.12	52.92
TiO2	1.889	1.823	1.706	1.632	1.037	U.722	U.746	U.744	U.682	U.679
AL2O3	12.00	12.49	12.47	12.98	20.33	16.00	12.25	12.31	13.28	13.57
FE2O3	13.87	12.64	13.05	10.65	14.60	9.96	5.51	5.99	11.49	10.99
MNO	U.15	U.14	U.15	U.15	U.20	U.14	U.07	U.09	U.18	U.17
MGO	2.15	2.37	3.21	2.72	7.98	8.32	0.88	0.83	8.46	8.38
CAO	4.95	3.52	3.05	6.44	3.09	6.74	1.80	4.58	10.18	10.02
NA2O	3.32	2.86	2.65	2.86	2.99	3.19	2.20	2.14	2.19	2.75
K2O	1.56	1.04	1.16	1.22	U.78	1.58	5.70	1.76	0.29	0.38
P2O5	U.702	0.684	0.592	U.596	U.204	0.147	0.246	0.241	U.136	U.136
TOTAL:	99.70	98.80	99.42	99.35	99.44	99.18	99.89	99.57	99.87	101.33
LOI	2.36	2.68	3.55	4.54	4.70	2.30	0.69	1.22	3.27	2.27
BA	617	473	549	711	184	1090	1230	645	135	230
NB	14.9	12.9	13.1	12.3	5.2	5.0	13.5	13.6	2.9	3.6
ZR	300	291	278	268	109	89	267	277	80	80
Y	46.6	45.5	43.9	37.1	26.0	20.4	27.4	29.2	17.5	19.0
SR	263	530	306	510	227	242	165	565	323	391
RB	43.8	39.8	44.7	30.3	21.2	29.7	121.0	33.6	5.1	8.0
ZN	146	128	141	119	155	92	78	74	85	75
CU	32	35	30	35	36	41	9	13	64	78
NI	25	32	31	34	601	220	12	11	318	320
CO	50	53	52	44	110	62	17	18	70	67
CR	10	9	9	8	1389	734	8	5	1012	974
V	310	315	311	264	269	204	50	49	204	198
LA	38	36	31	37	8	8	45	47	8	8
CE	81	76	75	65	30	23	94	100	19	24
ND	50	47	44	47	17	14	41	43	7*	14

SAMPLE:	DB- 85	KL- 86	KL- 87	DB- 88	DP- 89	KL- 90	PR- 91	DB- 92	DB- 93	DB- 94
SiO2	58.80	51.23	52.74	57.23	71.54	51.25	52.27	56.22	51.70	53.20
TiO2	U.794	U.409	U.614	1.167	U.832	U.468	1.880	1.598	U.763	U.676
AL2O3	14.76	14.99	14.58	14.61	12.47	14.64	13.30	13.05	13.06	13.70
FE2O3	9.24	9.15	11.21	13.90	6.58	9.11	14.44	14.50	12.36	10.36
MNO	U.14	U.15	U.17	U.17	U.06	U.16	U.26	U.18	U.18	U.14
MGO	4.86	9.12	7.69	4.43	1.17	10.13	7.75	3.41	9.15	8.47
CAO	4.95	12.18	9.99	4.15	U.78	10.71	7.08	6.91	9.54	9.46
NA2O	2.55	1.67	2.35	2.55	3.70	1.80	1.92	2.74	2.50	3.08
K2O	3.69	1.07	0.80	1.50	2.61	1.69	0.11	U.84	U.57	U.78
P2O5	U.219	U.033	U.064	U.292	U.259	U.046	0.987	0.546	U.172	U.135
TOTAL:	98.65	100.80	100.62	99.40	99.56	100.58	100.00	98.75	99.54	101.02
LOI	1.74	2.08	2.08	4.32	1.18	2.46	6.82	2.81	3.54	2.46
BA	2752	285	177	426	816	1470	105	510	223	504
NB	7.3	0.1	1.2	6.4	13.8	0.7	17.2	11.6	3.6	2.0
ZR	157	28	52	169	286	34	284	257	80	75
Y	23.3	15.2	21.4	28.9	27.4	16.7	57.4	42.1	20.3	17.5
SR	334	168	168	322	143	224	216	422	343	422
RB	59.7	41.9	39.2	41.5	73.1	101.0	6.5	31.4	14.7	18.3
ZN	90	61	78	122	84	69	167	137	93	77
CU	59	74	95	38	10	68	24	43	71	41
NI	64	181	117	28	10	186	187	40	321	292
CO	50	56	59	72	17	56	71	63	79	66
CR	28	774	120	7	12	932	659	19	1158	858
V	196	212	260	354	54	236	283	302	239	205
LA	17	0	5*	29	47	4*	52	40	9	10
CE	43	0	8*	68	88	0	122	84	24	20
ND	23	0	7	52	40	5*	64	45	16	16

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE	DB- 95	DP- 96	DB- 97	DP- 98	DP- 99	KA-100	KA-101	KA-102	KA-103	KA-104
SiO2	56.08	71.09	56.94	70.79	71.46	56.26	57.07	55.49	57.29	55.29
TiO2	0.855	0.811	1.112	0.764	0.678	1.040	1.080	1.094	1.040	1.076
Al2O3	15.355	12.38	13.44	12.63	11.45	14.44	14.45	15.62	14.47	14.975
Fe2O3	11.52	5.95	14.25	7.16	8.82	11.02	11.90	11.60	11.19	10.75
MnO	0.14	0.05	0.15	0.05	0.08	0.14	0.12	0.11	0.11	0.11
MgO	0.13	0.01	4.43	0.73	1.30	0.14	0.16	0.16	0.16	0.16
CaO	0.84	1.16	6.31	0.98	1.04	0.77	0.77	0.39	0.39	0.39
Na2O	3.44	3.37	2.66	2.36	2.76	0.67	0.69	0.09	0.08	0.08
K2O	1.30	3.93	0.21	4.36	2.22	0.91	0.75	0.75	0.75	0.75
P2O5	0.250	0.258	0.140	0.205	0.185	0.144	0.140	0.159	0.153	0.183

TOTAL:	98.98	100.14	100.44	99.43	100.15	100.74	100.50	99.95	100.76	99.65
LOI	2.33	1.64	6.58	1.78	2.41	10.35	2.99	3.47	2.83	3.51

BA	590	1019	104	833	482	1024	340	226	231	495
NB	5.4	13.6	5.5	14.5	14.0	4.9	5.1	6.2	6.3	6.2
ZR	152	275	114	244	220	113	119	126	120	125
Y	34.6	26.9	24.4	29.5	29.3	20.7	21.5	23.0	23.5	23.5
SR	539	173	282	71	54	507	278	232	245	392
RB	24.4	84.8	6.9	159.3	63.0	38.5	18.0	23.4	23.2	31.6
ZN	101	57	158	59	91	82	85	86	87	84
CU	35	11	61	15	14	8	8	8	7	10
NI	71	12	51	18	10	11	12	12	15	16
CO	62	17	67	42	21	61	58	65	61	57
CR	22	7	79	3	12	13	14	89	84	66
V	208	59	342	41	19	191	192	201	193	169
LA	21	48	21	39	39	11	5*	13	12	11
CE	47	89	36	116	79	25	28	33	28	32
ND	26	41	17	45	34	14	17	19	14	15

SAMPLE	KA-105	KA-106	KA-107	KO-108	KO-109	KO-110	KO-111	KO-112	KO-113	KO-114
SiO2	53.96	55.46	56.70	53.85	53.80	54.30	52.31	53.26	53.20	51.77
TiO2	1.097	1.017	1.114	1.032	1.078	1.046	1.097	0.996	1.025	1.022
Al2O3	15.18	15.00	15.59	15.07	14.90	14.77	15.54	14.96	14.58	15.43
Fe2O3	11.61	10.66	11.56	12.86	13.94	13.22	14.46	12.77	12.42	14.15
MnO	0.12	0.14	0.11	0.15	0.17	0.15	0.17	0.15	0.17	0.17
MgO	4.60	4.71	4.14	5.29	5.53	5.19	5.79	4.46	5.22	4.91
CaO	8.64	8.18	5.91	6.89	5.81	6.49	5.66	9.82	9.69	8.25
Na2O	3.22	3.38	4.02	3.99	3.43	3.91	3.90	2.69	2.73	2.89
K2O	1.40	1.29	0.70	0.75	1.22	0.79	0.95	0.76	0.83	1.18
P2O5	0.169	0.171	0.162	0.122	0.127	0.120	0.126	0.121	0.149	0.131

TOTAL:	100.77	100.23	99.35	100.56	100.06	99.64	100.18	100.58	98.94	100.69
LOI	2.13	3.36	3.17	1.99	2.19	2.00	2.11	1.89	2.38	2.15

BA	373	344	155	289	305	249	278	251	262	360
NB	5.1	4.9	8.1	5.6	4.2	5.6	5.4	6.6	6.8	4.0
ZR	123	117	131	105	103	104	109	98	105	101
Y	22.2	21.7	23.2	24.3	23.6	23.4	24.8	23.4	23.9	22.4
SR	493	517	147	350	332	369	282	671	497	304
RB	46.0	46.6	24.6	27.3	43.5	28.1	34.3	27.0	29.4	46.7
ZN	84	82	96	101	105	95	107	93	90	126
CU	114	108	105	104	104	113	112	119	92	64
NI	162	159	156	151	163	160	163	152	153	156
CO	64	57	63	74	75	74	79	75	66	64
CR	97	66	77	67	72	79	72	65	54	73
V	180	168	175	233	249	239	258	234	214	246
LA	10	10	17	12	9	12	27	18	8	12
CE	34	31	32	27	21	26	27	18	30	22
ND	15	12	16	12	13	13	15	10	10	15

SAMPLE	KO-115	KO-116	KO-117	KO-118	KO-119	KO-120	KO-121	KO-122	KO-123	KO-124
SiO2	55.53	54.53	54.85	53.26	51.03	51.46	49.50	54.79	54.30	53.57
TiO2	0.991	1.025	1.013	0.931	1.068	1.022	0.940	0.934	0.879	0.887
Al2O3	13.96	14.75	14.55	15.05	15.05	14.32	16.75	14.77	14.27	14.57
Fe2O3	11.47	13.23	13.16	11.78	13.90	13.31	10.01	12.35	11.83	11.62
MnO	0.14	0.14	0.14	0.13	0.17	0.18	0.10	0.18	0.16	0.14
MgO	5.29	5.77	5.44	4.80	5.74	6.21	2.33	5.39	4.58	4.50
CaO	7.49	5.67	6.02	10.66	8.13	6.48	19.61	6.82	5.89	7.63
Na2O	2.65	3.97	3.55	2.91	3.44	3.91	0.25	3.17	7.08	3.08
K2O	2.34	0.79	1.15	0.34	1.32	0.86	0.38	1.46	0.90	2.27
P2O5	0.138	0.124	0.123	0.134	0.143	0.138	0.150	0.125	0.108	0.115

TOTAL:	99.21	100.45	100.35	99.84	101.45	103.95	100.11	99.27	103.66	102.55
LOI	2.18	1.94	1.85	2.34	2.39	2.44	6.78	2.82	2.15	1.89

BA	691	392	361	148	477	358	126	494	218	1005
NB	6.0	5.5	5.2	5.1	5.6	5.6	5.1	5.1	4.2	3.9
ZR	100	103	98	91	103	100	100	97	90	91
Y	22.2	24.0	22.3	22.1	23.4	23.1	21.0	21.8	19.6	19.9
SR	356	353	303	781	504	214	266	258	174	249
RB	88.5	31.4	41.0	10.9	50.0	32.5	15.2	56.3	33.9	90.4
ZN	86	98	93	76	96	95	61	92	81	80
CU	137	103	106	65	121	77	117	96	89	91
NI	146	157	156	143	155	154	140	142	121	126
CO	64	78	71	63	70	67	55	66	58	63
CR	34	74	97	49	77	46	64	22	35	26
V	222	243	229	228	234	222	256	221	200	214
LA	10	18	12	9	11	8	11	10	11	16
CE	27	20	25	26	29	20	23	25	23	23
ND	10	15	15	12	16	11	80	11	0	15

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLEI	KO-125	KO-126	KO-127	KO-128	KO-129	KO-130	KL-131	KL-132	KL-133	KL-134
SI02	55.61	51.87	55.34	55.06	51.11	55.65	54.10	56.19	57.24	56.31
TI02	0.908	0.925	0.914	0.903	0.914	0.902	0.769	0.763	0.780	0.649
AL2O3	14.89	12.27	14.42	14.61	13.59	14.59	13.79	14.10	13.33	12.53
FE2O3	11.99	12.87	12.43	11.99	13.72	11.55	11.51	11.05	11.09	10.72
MNO	0.13	0.16	0.16	0.15	0.16	0.16	0.18	0.13	0.12	0.18
HGO	4.90	5.47	4.52	4.59	4.85	4.47	6.62	6.48	6.19	6.54
CAO	6.45	8.24	6.60	7.18	9.28	7.58	7.12	7.06	6.82	9.15
NA2O	3.89	3.22	3.70	3.04	3.22	3.32	3.72	3.75	3.45	3.12
K2O	1.13	1.85	1.78	2.37	0.91	2.63	2.09	0.38	0.89	0.72
P2O5	0.114	0.125	0.125	0.108	0.134	0.136	0.098	0.095	0.092	0.082

TOTAL:	101.23	100.43	100.87	100.50	99.91	99.57	100.47	100.21	100.48	99.79
LOI	1.69	1.98	1.62	1.44	2.05	2.25	1.79	2.10	1.90	2.02

BA	299	453	477	653	178	816	328	134	209	179
NBR	3.2	4.5	5.3	3.6	5.7	4.6	3.2	3.5	4.9	2.0
ZR	96	96	97	97	107	97	86	85	84	68
Y	23.0	21.6	25.1	22.8	25.5	21.2	19.1	20.0	20.1	17.8
SR	263	247	251	410	407	411	64	227	141	155
RB	44.0	65.9	68.0	55.2	33.9	108.9	74.7	12.7	29.8	19.5
ZN	87	94	86	82	104	87	90	77	84	77
CU	92	96	91	97	103	91	79	78	123	71
NI	127	133	113	111	131	111	172	180	164	190
CO	64	63	63	63	72	59	64	60	64	59
CR	29	26	26	20	28	16	233	272	238	400
V	216	209	212	222	235	215	193	187	193	195
LA	10	8	10	10	10	9	9	6	8	11
CE	22	23	30	15	28	31	22	3	17	15
ND	11	12	15	9	12	9	11	13	9	10

SAMPLEI	KL-135	KL-136	KL-137	KL-138	KL-139	KL-140	KL-141	KL-142	KL-143	KL-144
SI02	53.10	53.47	51.80	51.82	53.96	53.32	52.34	50.97	53.12	54.56
TI02	0.887	0.655	0.683	0.694	0.700	0.533	0.571	0.627	0.612	0.616
AL2O3	10.35	13.61	14.10	15.13	14.69	11.30	11.87	13.46	12.96	13.69
FE2O3	11.88	11.83	12.28	11.94	12.15	10.99	11.69	12.06	11.46	11.20
MNO	0.18	0.16	0.19	0.20	0.17	0.21	0.18	0.18	0.17	0.18
HGO	13.62	7.28	7.71	8.84	7.55	9.91	10.60	8.33	8.32	7.59
CAO	8.29	9.54	9.72	8.46	6.91	9.44	9.21	10.55	9.31	7.18
NA2O	1.85	2.09	2.82	3.37	3.08	2.39	1.38	3.08	2.69	3.09
K2O	0.18	1.30	0.59	1.46	0.70	1.85	1.68	0.68	1.29	1.83
P2O5	0.046	0.075	0.099	0.087	0.097	0.059	0.078	0.064	0.070	0.076

TOTAL:	101.14	100.44	99.26	100.01	99.99	101.31	99.12	99.80	100.96	98.85
LOI	3.35	2.18	3.31	2.94	3.42	1.55	2.74	2.69	2.20	2.29

BA	70	403	227	644	186	468	599	208	534	641
NBR	1.3	1.9	3.5	2.2	3.9	0.7*	1.7*	1.8	2.6	2.0
ZR	42	63	70	71	74	45	52	56	56	64
Y	12.5	18.5	18.2	18.4	18.1	14.6	14.9	17.4	17.6	18.8
SR	34.8	561	244	173	20.8	107	270	208	209	176
RB	8.8	60.4	22.6	43.3	20.8	55.2	79.7	20.3	56.9	78.6
ZN	78	79	107	81	98	81	76	93	74	80
CU	52	80	67	80	52	55	64	79	73	74
NI	43	206	196	166	173	281	308	221	225	220
CO	78	70	71	67	69	66	75	71	67	67
CR	2086	381	344	215	251	783	900	475	482	446
V	171	211	217	197	155	201	219	224	214	217
LA	0	10	5*	10	9	6*	4*	6*	9	5*
CE	10*	18	14	21	16	13*	9*	21	20	13
ND	0	11	0	10	6*	6*	0	11	6*	6*

SAMPLEI	KL-145	KL-146	KL-147	KL-148	KL-149	KL-150	KL-151	KL-152	KL-153	KL-154
SI02	54.50	51.77	53.95	54.12	55.26	52.11	56.82	54.87	52.25	53.59
TI02	0.604	0.519	0.595	0.571	0.553	0.670	0.642	0.696	0.737	0.477
AL2O3	13.78	12.72	13.07	13.51	13.68	14.80	14.20	14.73	15.38	12.73
FE2O3	11.46	11.08	11.63	11.10	11.01	11.85	10.23	10.64	12.10	10.79
MNO	0.17	0.19	0.24	0.15	0.15	0.15	0.13	0.15	0.23	0.18
HGO	7.65	12.87	7.85	7.93	8.35	7.36	5.23	6.67	7.11	10.75
CAO	9.86	7.19	7.50	8.16	8.11	8.54	6.91	7.96	6.29	7.75
NA2O	0.15	1.98	3.48	3.11	3.03	3.01	3.38	3.42	3.31	2.16
K2O	1.73	1.60	1.59	1.26	1.78	1.43	2.27	1.20	2.49	1.52
P2O5	0.071	0.078	0.083	0.071	0.068	0.082	0.131	0.105	0.102	0.070

TOTAL:	99.50	100.01	100.87	99.83	100.35	98.81	99.56	100.13	99.61	100.55
LOI	1.98	4.00	0.27	2.45	2.20	2.79	2.02	2.48	2.57	3.12

BA	231	433	249	337	866	669	735	331	680	435
NBR	6.1	2.3*	1.3	2.2	1.7	1.5	2.4	1.7*	3.0	1.6
ZR	113	60	62	60	55	73	75	77	81	56
Y	23.4	12.7	15.6	17.4	15.2	20.8	18.6	17.5	20.2	14.6
SR	34	109	49	29	21	258	314	524	140	361
RB	35.1	60.4	37.6	56.2	72.8	42.2	74.3	43.8	59.3	60.5
ZN	106	118	77	76	71	80	79	72	97	92
CU	69	48	63	65	60	62	80	72	82	56
NI	230	422	209	231	223	209	145	153	163	294
CO	63	26	65	66	63	66	56	60	67	64
CR	421	136	449	526	343	423	156	158	171	1200
V	225	190	204	207	207	213	182	197	207	156
LA	0	6*	9	7	7	6*	9	8	9	9
CE	32	8*	16	18	16	18	23	20	18	18
ND	17	0	9	6*	7	10	12	8	9	9

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE:	KL-155	KL-156	KL-157	KL-158	KL-159	PG-160	PG-161	PG-162	PG-163	PG-164
SiO2	54.46	54.02	54.23	50.81	53.82	62.04	60.37	60.07	56.13	51.85
TiO2	0.514	0.409	0.416	0.458	0.409	1.198	1.284	1.251	1.628	1.644
Al2O3	13.01	10.64	10.38	11.25	10.57	13.78	13.64	14.54	13.71	14.33
Fe2O3	10.22	10.96	10.86	12.05	11.24	9.83	9.01	9.68	11.48	11.80
MnO	0.16	0.19	0.17	0.18	0.19	0.09	0.11	0.13	0.15	0.16
MgO	11.22	13.43	14.27	14.89	14.07	13.94	4.25	3.37	6.80	6.96
CaO	6.36	8.23	7.68	8.37	7.94	7.94	4.78	3.38	6.99	6.65
Na2O	2.33	1.80	1.64	1.54	1.40	1.40	3.29	3.56	2.66	2.46
K2O	1.45	0.29	0.12	0.40	0.28	1.18	0.73	0.49	0.86	1.53
P2O5	0.084	0.052	0.042	0.060	0.066	0.320	0.540	0.535	0.795	0.815
TOTAL:	99.55	100.63	99.11	102.44	99.23	98.80	99.60	98.33	98.90	98.36
LOI	3.84	3.88	4.40	4.32	4.42	3.58	2.56	2.77	3.67	4.39
BA	430	111	49	146	120	777	1085	1010	886	806
NB	2.7	0.7*	1.1	1.6	0.0	17.4	17.5	16.6	14.2	13.3
ZR	63	38	41	45	41	302	310	313	253	254
Y	14.1	11.6	12.4	14.4	12.2	47.9	48.1	48.0	51.1	51.7
SR	443	215	137	126	180	494	351	228	401	103
RB	52.8	11.0	4.1	9.4	5.9	27.9	60.5	60.9	23.9	35.0
ZN	78	84	114	101	112	120	112	149	141	131
CU	53	51	43	51	40	233	19	29	48	31
NI	304	385	402	344	405	96	88	93	159	162
CO	70	75	76	72	78	41	37	40	50	26
CR	1287	1935	2289	1577	1771	250	255	258	337	388
V	153	163	171	190	171	159	160	175	225	236
LA	6*	3*	6	5*	4*	123	130	129	116	117
CE	14	10*	13*	10*	9*	60	59	66	66	66
ND	5*	4*	6*	6*	0					

SAMPLE:	PG-165	PG-166	PG-167	PG-168	PG-169	PG-170	PG-171 A	PG-171 B	PG-172	PG-173
SiO2	53.31	53.50	57.48	54.17	56.97	57.07	55.08	57.13	56.13	61.89
TiO2	1.637	1.579	1.686	1.614	1.569	1.667	1.666	1.611	1.549	1.460
Al2O3	13.49	14.84	14.49	14.70	13.97	15.55	15.71	14.15	15.32	13.70
Fe2O3	12.81	11.67	10.85	12.31	10.55	10.35	9.71	10.04	9.25	7.77
MnO	0.16	0.14	0.14	0.18	0.14	0.11	0.11	0.13	0.12	0.09
MgO	7.09	6.20	5.28	7.13	3.71	5.61	3.71	3.61	3.83	2.88
CaO	6.74	6.95	5.70	5.24	8.74	4.86	7.29	8.15	7.42	6.18
Na2O	2.82	3.73	3.41	3.80	3.12	3.10	3.61	3.01	3.44	3.17
K2O	1.18	0.65	0.18	1.05	0.51	0.88	2.54	1.41	2.11	2.16
P2O5	0.760	0.742	0.793	0.799	0.735	0.805	0.792	0.765	0.794	0.713
TOTAL:	100.19	98.40	98.26	99.15	100.13	98.17	99.45	99.71	98.63	99.52
LOI	3.12	2.77	4.21	3.53	3.85	4.73	5.28	3.76	4.77	2.86
BA	522	443	172	639	384	977	1029	689	999	1131
NB	13.8	16.6	16.9	17.2	16.5	18.7	18.4	17.2	17.4	18.2
ZR	244	272	304	292	287	333	325	307	314	309
Y	30.7	51.9	55.8	51.6	50.1	56.9	58.7	51.6	50.5	45.4
SR	566	470	564	502	1022	469	555	830	587	596
RB	26.1	12.4	3.0	19.0	8.3	21.3	44.0	24.3	41.8	44.5
ZN	144	131	132	140	109	125	124	125	117	88
CU	43	33	29	36	38	41	38	36	33	42
NI	161	159	145	167	118	105	109	103	124	107
CO	53	49	48	51	46	43	44	42	45	36
CR	369	371	376	380	242	167	191	186	253	207
V	236	247	219	214	200	204	204	188	197	168
LA	53	51	58	47	54	46	65	56	60	53
CE	113	112	129	116	120	112	141	124	123	123
ND	58	60	71	59	64	60	75	65	65	63

SAMPLE:	PG-174	PG-175	PG-176	PG-177	PM-178	PM-179	PM-180	PR-181	PR-182	PR-183
SiO2	50.47	60.16	59.27	78.24	64.95	66.79	60.51	56.33	57.53	56.32
TiO2	1.733	1.431	1.399	0.254	1.122	1.172	1.300	1.434	1.280	1.440
Al2O3	17.27	14.83	14.57	5.91	13.97	14.21	15.38	15.39	13.66	15.74
Fe2O3	13.65	9.20	8.62	9.05	7.79	5.75	8.83	10.69	9.84	8.86
MnO	0.16	0.11	0.12	0.13	0.08	0.08	0.09	0.16	0.10	0.11
MgO	6.26	4.22	2.86	3.85	3.38	2.08	3.69	6.52	5.79	4.14
CaO	6.09	4.03	8.12	2.03	2.21	4.52	5.45	5.33	6.61	6.78
Na2O	2.69	3.14	3.02	0.09	1.31	4.04	3.99	3.03	3.41	3.35
K2O	0.88	2.23	1.38	0.41	4.72	0.80	1.19	0.43	1.17	2.54
P2O5	0.796	0.658	0.638	0.036	0.461	0.542	0.560	0.683	0.609	0.735
TOTAL:	98.70	98.48	98.93	99.10	99.72	98.90	98.41	98.83	98.62	99.54
LOI	4.28	3.76	4.42	3.84	3.00	1.96	2.71	5.97	3.28	5.82
BA	1034	984	590	204	1670	171	1054	293	809	1291
NB	21.6	19.8	16.6	26.7	30.2	32.1	32.6	14.3	13.5	14.5
ZR	376	326	316	360	491	549	560	276	245	245
Y	61.0	48.6	68.8	57.4	62.1	73.4	66.0	42.1	39.6	44.5
SR	670	303	805	67	113	926	1270	312	385	232
RB	18.0	45.6	29.6	12.1	158.2	34.0	27.6	10.8	24.4	53.9
ZN	161	105	107	57	99	93	100	117	100	99
CU	40	30	31	10	12	13	23	27	36	33
NI	129	99	89	26	30	17	31	146	170	105
CO	54	42	39	21	23	20	29	50	50	42
CR	238	196	168	0	26	15	24	386	520	150
V	221	182	174	25	98	89	140	179	169	189
LA	64	56	56	52	73	100	94	48	47	45
CE	149	125	126	89	174	205	205	102	101	109
ND	79	61	62	33	81	94	94	53	51	55

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE	PR-184	PR-185	PR-186	PR-187	PK-188	PR-189	PR-190	PR-191	PR-192	PR-193
SiO2	47.50	54.51	52.79	56.45	56.96	55.72	57.43	52.23	59.41	56.38
TiO2	2.202	1.831	1.924	1.713	1.711	1.642	1.602	1.601	1.444	1.386
Al2O3	16.47	13.66	14.66	14.31	14.62	14.59	13.91	16.79	15.85	14.89
Fe2O3	14.26	12.92	12.52	11.11	11.21	10.90	10.09	13.11	6.20	11.90
MnO	0.16	0.18	0.15	0.13	0.13	0.14	0.13	0.14	0.07	0.16
MgO	5.41	6.48	5.68	4.43	4.50	5.35	4.24	5.50	2.59	5.91
CaO	9.39	6.22	7.97	5.72	6.51	6.55	6.22	5.60	8.14	5.91
Na2O	3.23	2.65	2.69	3.17	2.94	4.02	3.30	3.29	3.97	2.60
K2O	0.21	0.58	0.63	0.13	0.53	0.15	0.28	0.92	1.58	0.89
P2O5	1.168	0.966	0.975	0.831	0.877	0.842	0.788	0.820	0.657	0.674
TOTAL	99.10	98.92	99.21	99.20	98.14	98.33	99.90	98.10	98.64	99.18
LOI	6.76	3.89	6.79	3.55	4.65	2.92	3.16	5.32	6.45	6.77
BA	312	696	482	1182	398	164	1599	799	832	435
NB	19.8	17.7	21.2	19.9	19.1	21.2	20.6	18.8	20.4	17.7
ZR	321	305	334	348	355	385	363	369	387	351
Y	67.6	55.8	59.2	56.7	55.4	59.8	54.0	58.2	48.4	50.6
SR	775	740	448	291	829	813	702	963	679	562
RB	8.4	12.8	12.8	40.1	7.9	3.2	39.7	15.0	63.6	24.3
ZN	162	152	143	129	119	146	121	151	75.6	116
CU	58	44	41	40	47	31	36	41	39	31
NI	153	166	133	120	101	115	106	133	82	115
CO	62	60	51	48	46	46	43	56	35	48
CR	362	479	306	280	211	246	300	318	184	294
V	314	234	229	234	227	218	202	240	184	195
LA	65	56	57	66	64	60	65	64	63	53
CE	144	132	131	131	139	139	147	136	133	123
ND	78	67	71	76	68	70	74	71	63	60

SAMPLE	PH-194	AR-195	AR-196	AR-197	AR-198	AR-199	KA-200	KO-201	KO-202	KL-203
SiO2	53.93	58.33	56.70	55.38	61.26	57.15	57.47	56.84	54.06	54.55
TiO2	1.465	1.242	1.229	1.167	1.182	1.146	1.074	0.982	1.012	0.674
Al2O3	15.48	14.35	14.70	14.88	13.66	13.80	14.42	13.92	14.82	13.97
Fe2O3	10.72	12.04	11.86	15.12	10.05	11.46	11.62	12.38	13.31	11.58
MnO	0.12	0.16	0.15	0.15	0.09	0.12	0.12	0.14	0.15	0.16
MgO	7.36	3.27	2.21	4.05	3.20	4.18	4.62	4.48	5.36	6.94
CaO	7.73	5.99	8.17	5.36	3.25	6.45	6.37	6.74	6.99	7.63
Na2O	1.05	4.12	4.15	3.16	3.80	4.05	2.70	2.21	3.21	3.05
K2O	1.47	0.30	0.62	0.53	1.30	1.42	1.45	2.18	0.97	1.38
P2O5	0.671	0.207	0.218	0.198	0.210	0.228	0.161	0.127	0.127	0.075
TOTAL	100.43	99.30	100.36	99.36	100.13	99.98	98.77	99.32	100.28	98.26
LOI	9.30	5.92	7.77	6.12	5.19	3.32	6.48	2.02	2.14	2.12
BA	385	524	259	271	362	565	1189	617	254	518
NB	20.2	10.4	10.3	10.3	10.5	17.7	5.5	4.3	4.5	3.0
ZR	367	178	180	177	180	157	124	101	105	68
Y	51.6	27.5	26.9	28.2	28.0	27.4	24.4	22.9	23.7	18.6
SR	217	723	690	802	678	1472	353	358	461	328
RB	51.2	10.7	24.2	11.8	29.3	40.0	39.8	86.9	56.4	58.6
ZN	129	115	111	158	93	94	82	89	97	77
CU	11	118	123	164	115	123	95	114	122	80
NI	115	110	106	116	113	143	163	148	151	190
CO	46	62	62	64	56	60	64	63	69	71
CR	286	10	13	14	16	20	98	57	62	390
V	213	164	180	186	164	177	199	213	221	217
LA	64	26	30	25	24	26	16	15	11	4*
CE	138	55	60	51	55	54	38	25	25	12*
ND	69	25	30	25	27	23	17	11	13	5*

SAMPLE	KL-204	KL-205	PG-206	PG-207	PG-208	PH-209	PH-210	PH-212	PH-213	PH-217
SiO2	53.60	54.90	58.23	53.67	52.50	59.75	68.29	66.24	68.03	65.41
TiO2	0.499	0.427	1.285	1.599	1.897	1.373	1.064	1.226	1.27	1.151
Al2O3	12.42	10.53	13.96	13.60	16.10	15.48	13.09	15.19	13.91	13.74
Fe2O3	10.17	10.43	10.15	11.84	12.41	8.59	6.05	6.38	5.86	7.45
MnO	0.17	0.18	0.12	0.16	0.16	0.05	0.05	0.09	0.02	0.10
MgO	9.22	12.24	5.15	6.42	3.74	3.76	0.77	1.66	1.72	1.71
CaO	7.99	8.82	5.43	8.15	3.86	2.44	2.65	2.04	2.51	4.09
Na2O	2.39	1.90	2.86	3.04	3.88	0.51	2.96	6.56	1.54	3.56
K2O	1.47	0.52	2.24	0.75	0.42	7.47	4.64	0.07	4.68	2.31
P2O5	0.073	0.060	0.564	0.782	1.029	0.584	0.436	0.545	0.517	0.478
TOTAL	98.91	101.56	100.78	99.72	99.92	99.43	99.07	98.63	99.14	99.36
LOI	2.54	3.16	2.78	3.74	5.15	2.92	2.04	1.70	2.74	2.62
BA	419	187	745	541	595	2784	2017	25	1721	1942
NB	2.2	0.7*	16.4	14.8	18.9	36.9	28.3	34.6	33.9	28.2
ZR	66	45	315	271	317	594	476	561	527	511
Y	16.5	13.9	48.0	51.7	68.4	79.8	66.4	76.7	78.1	73.2
SR	67	236	493	672	583	197	330	203	207	888
RB	51.4	23.6	53.2	17.5	7.6	207.2	112.3	2.0	234.4	55.8
ZN	84	60	121	143	136	123	96	99	74	112
CU	52	50	49	40	42	15	11	10	13	17
NI	238	338	98	158	168	338	25	17	27	26
CO	62	69	44	50	59	25	19	20	18	21
CR	1032	1697	272	392	398	27	20	21	27	21
V	180	167	172	209	223	120	91	98	109	100
LA	10	4*	60	49	49	104	88	107	128	98
CE	15	12*	135	114	159	208	192	205	243	193
ND	11	5*	64	59	82	96	85	94	102	93

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE:	PM-219	PM-222	PM-228	PM-231	PM-233	PR-234	AR-236	AR-238	AR-241	AR-244
SiO2	65.71	68.23	66.00	61.88	59.78	58.46	56.15	58.43	55.90	51.43
TiO2	1.249	1.082	1.008	1.341	1.187	1.346	1.322	1.193	1.251	1.228
Al2O3	14.55	13.54	12.96	16.87	14.78	14.26	15.35	14.54	14.79	14.31
Fe2O3	9.86	6.04	6.42	7.70	13.11	9.91	5.09	10.15	12.46	14.47
MnO	0.12	0.10	0.09	0.08	0.13	0.13	0.25	0.11	0.18	0.22
MgO	3.28	1.70	1.06	2.31	3.22	4.71	2.02	2.28	3.55	5.38
CaO	1.84	3.85	7.93	2.84	2.01	6.61	14.66	8.29	6.83	9.14
Na2O	2.86	3.94	3.42	3.33	2.24	3.21	0.16	4.26	3.67	2.99
K2O	1.98	1.05	0.60	2.77	1.00	0.74	4.59	0.54	1.14	0.58
P2O5	0.538	0.460	0.453	0.669	0.531	0.636	0.228	0.208	0.228	0.252
TOTAL:	98.99	99.65	99.41	98.25	98.63	99.61	99.60	99.62	99.21	98.34
LOI	2.93	1.99	3.15	2.98	4.29	4.57	12.61	6.95	6.50	8.62
BA	660	337	229	1104	368	686	998	386	379	233
NB	32.8	28.8	23.6	40.1	31.0	15.9	10.1	11.0	9.4	7.2
ZR	540	489	464	688	315	341	187	179	179	174
Y	69.2	69.9	65.4	84.6	63.1	49.1	30.1	27.6	28.4	26.2
SR	178	962	2943	488	428	1395	403	874	305	342
RB	82.7	45.5	22.2	103.2	34.7	9.6	17.4	16.3	25.8	15.2
ZN	151	113	78	140	204	115	130	129	96	121
CU	14	16	28	17	20	109	130	104	146	193
NI	27	22	28	28	17	42	31	54	64	64
CO	26	16	19	19	33	283	21	11	32	78
CR	24	23	30	25	17	184	166	176	211	211
V	108	95	90	124	94	63	184	30	27	31
LA	43	94	83	140	38	134	28	57	61	48
CE	112	193	176	282	132	68	60	31	29	27
ND	65	88	83	186	61					

SAMPLE:	DB-247	DB-251	DB-254	DB-257	DB-259	DP-261	DP-263	DP-265	DP-269	DB-273
SiO2	59.58	54.52	51.84	53.85	49.95	69.32	55.93	78.93	68.93	60.28
TiO2	1.660	0.664	0.691	0.624	1.319	0.845	1.080	0.660	0.788	1.147
Al2O3	12.45	15.89	14.76	14.37	17.25	13.17	15.11	10.93	12.63	13.70
Fe2O3	12.81	9.46	12.34	9.42	17.15	6.74	11.78	2.43	7.46	13.20
MnO	0.18	0.15	0.17	0.16	0.21	0.09	0.12	0.03	0.11	0.17
MgO	5.18	6.76	7.85	8.20	5.33	1.57	5.44	0.85	0.50	2.39
CaO	6.66	7.39	8.57	9.87	4.76	1.70	5.55	0.43	0.79	2.65
Na2O	1.72	3.93	3.40	2.33	2.72	4.36	2.26	0.20	3.21	4.29
K2O	1.20	1.08	0.43	1.03	0.75	0.92	0.46	0.34	0.21	2.01
P2O5	0.567	0.166	0.140	0.138	0.362	0.286	0.156	0.208	0.219	0.168
TOTAL:	99.83	99.28	100.54	100.18	100.25	99.16	99.13	99.48	99.53	98.09
LOI	3.72	2.02	2.24	2.04	4.22	2.21	2.57	1.55	2.84	5.35
BA	951	662	232	321	278	516	475	1013	642	376
NB	11.7	3.4	2.5	2.9	8.2	12.8	16.1	11.4	14.5	5.4
ZR	258	97	85	80	172	271	353	233	245	139
Y	40.0	18.9	18.7	17.3	30.7	25.7	36.0	23.4	33.8	24.2
SR	505	441	325	463	337	159	203	43	213	244
RB	31.8	16.8	6.5	29.2	13.3	40.1	56.3	150.1	97.6	91.4
ZN	121	81	94	76	166	68	93	40	90	116
CU	31	57	77	76	6	6	6	10	10	4
NI	31	216	389	25	44	9	10	13	18	26
CO	50	57	69	61	78	19	17	10	3	6
CR	10	299	883	746	5	3	9	9	18	61
V	281	180	201	182	402	76	59	50	26	387
LA	51	13	8	11	29	28	30	27	47	25
CE	80	27	21	23	60	77	97	51	88	51
ND	45	12	8	7	27	35	42	22	55	18

SAMPLE:	DP-276	DB-278	DB-281	DB-285	DB-286	DP-288	KA-291	KO-295	AR-300	AR-304
SiO2	73.65	55.35	52.78	55.49	56.56	69.39	56.60	55.55	61.89	60.38
TiO2	0.322	1.252	0.671	0.818	1.413	0.814	1.047	1.074	1.222	1.271
Al2O3	12.70	13.79	14.39	15.69	14.31	12.57	15.39	14.66	14.36	14.80
Fe2O3	5.04	11.28	10.63	9.85	14.26	6.39	11.74	13.47	9.75	7.64
MnO	0.07	0.18	0.16	0.14	0.18	0.09	0.12	0.15	0.12	0.13
MgO	0.68	5.93	8.35	5.20	3.20	0.94	0.07	0.23	3.64	2.88
CaO	0.98	5.01	9.81	7.88	6.02	3.46	5.53	3.22	2.80	3.60
Na2O	3.14	3.99	2.51	3.22	6.90	3.36	3.92	3.24	2.81	3.62
K2O	3.09	0.83	0.55	1.48	0.81	2.72	0.74	0.90	3.18	3.46
P2O5	0.125	0.383	0.151	0.438	0.344	0.259	0.140	0.124	0.220	0.227
TOTAL:	99.41	98.10	100.47	96.24	98.76	98.98	100.53	99.35	99.69	98.69
LOI	1.86	1.86	2.11	1.64	3.06	3.38	3.48	2.89	3.52	4.96
BA	370	416	230	837	501	625	304	233	707	741
NB	14.8	8.1	3.3	5.4	8.5	13.4	5.5	5.4	11.9	12.2
ZR	269	185	78	150	223	272	122	110	187	193
Y	24.4	33.8	17.7	24.8	34.5	28.7	21.4	21.3	23.1	24.3
SR	89	203	428	607	409	275	360	197	71	95
RB	102.5	11.6	12.9	26.5	14.0	73.4	35.8	31.9	69.3	75.1
ZN	51	132	84	91	140	80	92	105	84	64
CU	13	48	71	67	39	5	72	68	10	5
NI	8	115	292	23	25	6	167	161	113	9
CO	11	61	63	33	62	20	59	76	57	49
CR	5	47	801	29	0	0	81	60	18	9
V	21	276	199	210	372	73	180	208	179	147
LA	55	22	6	19	34	33	14	9	14	16
CE	58	35	22*	50	69	83	33	23	37	38
ND	38	31	7*	42	34	35	10	7	18	19

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE I	AR-305	AR-307	PG-312	PG-314	PG-316	PG-319	PG-322	PG-325	PG-328	PM-329
SiO2	58.00	57.97	59.07	50.68	65.42	57.89	59.32	56.96	60.06	54.93
TiO2	1.295	1.201	1.242	1.732	1.289	1.281	1.364	1.530	1.267	1.729
AL2O3	14.79	14.27	14.15	19.44	13.48	13.95	15.14	13.37	15.34	18.55
FE2O3	14.04	10.00	11.34	14.30	8.10	9.99	9.87	11.01	8.68	10.18
MNO	0.13	0.16	0.14	0.12	0.14	0.18	0.16	0.14	0.11	0.12
MGO	4.78	4.95	5.34	4.60	2.35	3.97	5.61	6.66	4.66	4.07
CAO	2.43	6.23	4.28	2.76	4.99	7.03	3.08	5.62	4.38	3.16
NA2O	2.78	3.45	2.72	3.00	2.81	2.54	4.11	3.19	4.38	2.93
K2O	1.58	1.50	0.87	2.63	2.87	2.62	0.81	0.57	0.73	3.64
P2O5	0.176	0.266	0.551	0.748	0.541	0.550	0.542	0.754	0.591	0.687
TOTAL:	99.83	98.33	99.92	98.60	98.73	98.75	98.42	97.62	99.23	97.08
LOI	3.87	4.84	5.83	4.25	2.10	3.22	2.87	2.56	2.88	2.90
BA	301	349	865	1042	1198	1475	522	549	413	2147
NB	9.5	7.6	18.4	32.6	24.4	18.1	22.8	18.0	21.3	45.8
ZR	180	185	355	705	493	323	455	297	410	748
Y	22.2	28.9	50.1	95.5	64.2	48.5	54.4	52.2	52.6	92.9
SR	55	42	42	30	77	85	41	49	54	52
RB	33.8	24.6	21.4	30.0	27.3	21.6	21.8	13.6	33.8	33.7
ZN	116	103	136	174	100	114	110	130	102	139
CU	7	4	12	15	1	2	1	1	1	1
NI	162	198	70	39	3	9	6	16	1	3
CO	21	6	40	100	268	43	38	47	35	31
CR	72	128	171	100	1	253	213	442	120	39
V	197	208	189	253	151	181	158	219	182	157
LA	14	24	65	90	65	62	51	49	81	123
CE	41	57	136	196	199	126	121	113	163	263
ND	20	29	64	96	93	60	58	60	75	119

SAMPLE I	PM-332	PM-335	PM-338	AR-339	AR-341	AR-343	AR-345	KA-346	KA-347	KA-348
SiO2	66.91	68.75	67.40	59.05	55.76	52.42	57.15	52.56	58.13	53.26
TiO2	1.090	1.116	1.082	1.190	1.332	1.174	1.071	1.129	1.005	1.108
AL2O3	13.24	12.75	13.61	14.42	14.67	14.27	14.91	14.87	14.13	15.13
FE2O3	6.76	6.60	6.13	9.93	11.70	12.76	11.82	12.51	10.10	12.27
MNO	0.14	0.07	0.08	0.18	0.14	0.16	0.15	0.14	0.14	0.13
MGO	1.56	1.76	2.42	3.35	4.03	4.91	4.81	7.29	4.89	5.16
CAO	2.94	2.33	3.95	8.06	7.93	9.36	5.37	8.24	6.43	8.13
NA2O	3.30	3.16	2.79	1.18	3.98	4.06	3.93	2.04	4.66	3.80
K2O	3.56	3.78	2.10	2.41	0.97	0.65	0.58	1.07	0.58	0.87
P2O5	0.462	0.468	0.459	0.225	0.271	0.241	0.224	0.152	0.137	0.165
TOTAL:	99.92	100.18	100.57	99.16	100.07	101.05	98.79	101.77	100.21	100.72
LOI	2.06	1.85	3.62	8.74	6.43	5.06	3.58	10.11	2.25	2.23
BA	1330	1310	509	509	208	349	347	445	204	591
NB	28.8	31.9	30.4	11.4	9.0	7.3	6.5	2.0	5.2	6.4
ZR	478	468	452	178	178	168	167	60	108	118
Y	67.9	67.9	63.9	28.3	25.3	28.3	24.4	17.7	20.7	23.7
SR	493	202	224	209	283	606	499	280	257	588
RB	88.0	91.2	85.5	98.5	23.9	17.4	17.4	87.7	11.7	26.7
ZN	104	87	91	96	127	93	135	81	67	105
CU	10	10	8	114	92	50	57	70	77	122
NI	22	23	20	102	176	184	165	211	101	167
CO	20	19	18	26	78	71	71	69	34	63
CR	11	17	20	18	48	57	55	413	124	64
V	96	106	106	167	215	201	195	211	179	171
LA	93	84	70	25	23	23	24	11	13	11
CE	191	177	140	51	51	42	51	17	25	34
ND	88	79	68	22	26	26	27	8	13	16

SAMPLE I	KA-349	KA-350	KO-352	KO-355	KO-356	AR-359	AR-360	AR-361	AR-362	AR-365
SiO2	55.12	54.83	54.75	53.50	58.88	57.76	53.96	55.38	55.09	54.18
TiO2	0.973	1.092	0.984	0.919	0.925	1.240	1.241	1.321	1.181	1.102
AL2O3	15.54	14.98	14.62	15.10	15.77	14.38	15.06	14.14	14.15	13.30
FE2O3	10.83	11.12	12.45	13.37	11.07	9.93	13.16	12.89	11.75	12.04
MNO	0.15	0.15	0.15	0.18	0.11	0.13	0.18	0.13	0.14	0.16
MGO	4.41	4.82	4.65	5.79	3.57	3.12	5.00	4.71	4.98	5.35
CAO	7.43	6.41	7.42	6.21	5.17	3.71	5.93	6.55	8.13	8.75
NA2O	4.57	3.67	3.52	3.21	3.64	3.44	4.11	3.80	5.26	3.34
K2O	0.84	2.77	1.52	1.61	0.75	3.08	1.14	1.80	1.06	1.47
P2O5	0.144	0.169	0.132	0.106	0.114	0.207	0.220	0.282	0.259	0.241
TOTAL:	101.66	102.31	100.51	101.55	101.62	100.63	100.97	100.22	99.55	99.73
LOI	1.94	2.18	1.96	2.42	2.57	5.56	4.22	6.20	5.09	4.79
BA	235	730	408	1208	164	649	350	355	255	596
NB	5.8	6.0	2.9	4.0	4.1	10.6	9.6	8.7	8.7	5.5
ZR	116	120	97	96	96	183	180	194	176	170
Y	22.3	24.4	22.2	19.6	19.1	25.1	26.5	31.6	26.0	27.7
SR	244	332	466	317	728	426	311	122	519	579
RB	24.5	86.8	52.3	37.4	22.1	57.1	34.0	29.6	16.2	20.2
ZN	94	84	94	104	67	99	125	106	98	106
CU	92	97	99	77	16	104	118	106	83	90
NI	100	152	139	128	126	112	128	164	209	220
CO	61	62	65	69	47	54	75	67	69	67
CR	64	66	59	25	29	15	17	53	147	176
V	171	186	225	235	188	164	187	207	201	195
LA	14	19	11	11	10	23	23	23	25	20
CE	33	34	23	27	25	55	47	50	51	53
ND	16	17	15	13	12	26	23	27	26	28

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE I	AR-364	AR-365	JC-366	JC-367	JC-368	JC-369	JC-370	JC-371	KA-373	KA-376
SI02	57.10	57.86	64.68	64.52	66.82	65.09	65.58	50.22	55.74	52.77
TI02	1.135	1.002	0.998	1.046	1.042	1.026	1.017	1.670	1.015	1.174
AL2O3	14.05	14.37	12.92	12.30	11.94	12.34	12.88	19.57	14.36	16.66
FE2O3	9.60	10.49	11.46	11.69	12.65	11.92	11.28	16.75	10.59	12.23
MNO	0.11	0.13	0.13	0.17	0.17	0.17	0.16	0.26	0.17	0.16
MGO	4.32	3.37	0.84	0.86	1.09	0.86	0.89	1.61	4.66	4.66
CAO	7.05	6.70	3.27	3.25	2.90	3.37	1.96	3.26	9.03	6.81
NA2O	4.03	3.77	3.65	4.20	2.72	3.51	3.84	4.20	3.43	3.39
K2O	2.27	1.99	1.53	1.45	0.15	0.31	1.90	1.59	0.86	1.97
P2O5	0.248	0.231	0.503	0.504	0.512	0.517	0.496	0.869	0.140	0.172
TOTAL:	101.00	101.34	99.51	99.82	98.31	98.58	100.17	98.88	99.79	98.45
LOI	3.52	5.82	2.40	2.29	2.52	2.91	2.07	3.20	2.42	5.54
BA	556	360	554	644	122	363	840	1210	403	355
NB	7.9	8.8	14.2	16.1	15.7	14.4	15.2	23.4	6.1	6.0
ZR	179	177	235	241	255	247	233	383	114	137
Y	26.4	25.8	51.7	52.1	54.3	51.4	46.8	72.3	22.1	22.4
SR	636	386	594	219	194	355	211	487	929	249
RB	28.6	24.0	48.2	66.8	4.3	12.1	55.5	34.5	26.7	63.7
ZN	90	84	161	155	164	164	160	232	79	95
CU	103	38	23	30	13	21	22	26	78	59
NI	121	133	13	14	15	11	12	18	113	173
CO	32	55	26	29	29	29	24	37	51	59
CR	50	57	9	6	3	6	9	8	115	86
V	189	202	15	18	18	17	17	27	172	188
LA	24	24	42	43	40	47	41	59	6	15
CE	50	60	87	88	90	98	81	124	23	32
ND	26	27	47	46	46	48	41	57	10	12

SAMPLE I	KL-382	KL-392	PR-395	PR-398	AR-400	AR-403	PG-406	PG-417	IN-419	PG-421
SI02	54.47	53.56	48.52	56.06	53.60	54.32	59.06	56.70	49.28	58.73
TI02	0.500	0.505	1.654	1.531	1.167	1.167	1.425	1.501	1.834	1.402
AL2O3	12.79	14.50	15.99	14.17	14.77	14.38	13.85	13.39	14.95	13.42
FE2O3	10.86	11.08	9.14	11.13	11.58	14.70	10.07	11.05	16.29	10.00
MNO	0.17	0.17	0.24	0.18	0.19	0.16	0.13	0.17	0.18	0.16
MGO	9.00	8.83	7.19	5.09	4.59	5.37	5.74	6.00	6.51	4.26
CAO	8.04	6.88	12.68	6.97	4.93	6.77	4.33	6.19	6.86	5.58
NA2O	2.15	3.05	3.58	2.48	2.02	1.50	2.25	2.83	3.33	3.48
K2O	1.95	1.36	0.50	1.66	1.88	1.22	2.58	1.50	0.46	2.35
P2O5	0.071	0.070	0.687	0.713	0.277	0.224	0.579	0.677	0.293	0.606
TOTAL:	101.17	98.95	98.24	98.82	98.54	98.55	98.99	98.03	100.18	99.55
LOI	2.34	3.17	10.65	6.76	9.59	9.78	5.82	3.31	2.60	1.74

SAMPLE I	KL-382	KL-392	PR-395	PR-398	AR-400	AR-403	PG-406	PG-417	IN-419	PG-421
BA	297	384	213	899	261	104	1005	821	231	845
NB	1.4*	0.0	19.6	16.4	6.6	9.2	18.7	14.9	6.7	18.7
ZR	59	66	336	288	181	180	286	276	159	348
Y	15.9	15.5	59.3	52.3	28.1	27.6	46.5	49.9	47.3	51.1
SR	217	144	268	966	524	207	374	759	484	616
RB	89.6	36.0	15.9	50.3	54.1	39.7	88.8	35.2	8.8	46.5
ZN	74	117	100	127	108	109	143	115	118	82
CU	53	49	3	54	37	97	28	27	85	62
NI	216	202	192	154	214	124	150	136	92	88
CO	66	65	51	48	67	70	46	49	69	43
CR	584	503	542	347	49	189	457	374	253	277
V	178	179	199	221	226	211	181	207	331	159
LA	7	8	49	53	21	21	49	46	9	56
CE	14	13	105	118	57	54	113	103	28	124
ND	6*	4*	58	61	29	20	56	57	18	62

SAMPLE I	PG-426	PG-433	AR-442	AR-446	AR-451	KA-452	KA-455	KO-457	KO-460	KO-462
SI02	55.85	54.54	53.98	58.88	58.44	53.91	54.76	56.09	54.26	54.11
TI02	1.614	1.580	1.298	1.076	0.891	1.046	1.081	1.008	0.951	0.946
AL2O3	14.06	16.76	16.66	14.67	14.37	13.76	15.21	13.69	14.89	14.95
FE2O3	10.67	10.46	12.58	10.64	8.11	12.30	11.84	13.35	12.14	12.94
MNO	0.17	0.09	0.17	0.15	0.14	0.15	0.14	0.16	0.13	0.15
MGO	5.50	5.62	5.17	4.02	4.49	6.57	4.50	5.54	5.17	4.91
CAO	6.02	5.16	5.60	4.62	6.95	9.52	7.07	5.92	6.62	5.99
NA2O	2.63	4.46	3.76	4.03	4.05	2.23	3.24	3.18	4.13	2.98
K2O	2.71	0.70	0.50	1.68	2.41	0.38	1.99	0.92	1.57	2.90
P2O5	0.782	0.633	0.278	0.237	0.164	0.139	0.163	0.125	0.123	0.121
TOTAL:	98.08	98.04	98.81	98.37	99.32	101.56	99.30	100.66	98.96	99.37
LOI	3.33	4.83	6.02	3.14	3.46	4.15	1.76	2.00	1.70	1.95

SAMPLE I	PG-426	PG-433	AR-442	AR-446	AR-451	KA-452	KA-455	KO-457	KO-460	KO-462
BA	1201	391	213	741	924	241	638	348	293	835
NB	17.2	28.6	8.1	8.7	6.6	4.0	5.6	6.3	4.6	4.7
ZR	309	515	191	177	141	106	122	100	101	100
Y	49.1	66.1	30.8	26.1	22.5	22.0	23.3	23.1	22.5	23.2
SR	527	430	332	393	451	566	512	258	213	190
RB	71.1	15.0	13.0	38.2	52.8	10.1	68.5	36.6	61.3	111.9
ZN	119	102	138	99	52	73	84	90	83	96
CU	27	18	5	66	95	109	103	79	95	103
NI	154	98	226	145	104	226	162	152	146	138
CO	49	41	82	66	49	64	63	69	64	68
CR	446	277	86	61	145	415	85	68	68	27
V	229	186	252	192	187	194	179	236	221	258
LA	53	70	20	23	21	14	13	10	11	10
CE	128	156	44	53	41	27	34	20	31	27
ND	66	75	23	22	17	14	15	9	16	17

OXIDES AND TRACE ELEMENTS RECALCULATED VOLATILE FREE, ALL FE AS FE2O3, TOTAL = ORIGINAL TOTAL:

SAMPLE I	KL-465	KL-467	KL-468	KL-469	KL-471	PR-472	PR-474	PR-476	PR-477	PR-479
SiO2	55.52	56.34	52.81	54.78	56.61	56.14	56.59	61.59	58.92	58.24
TiO2	0.878	0.736	0.400	0.637	0.599	1.067	1.264	1.350	1.427	1.355
Al2O3	14.31	14.47	8.59	14.92	17.75	17.01	13.54	14.46	14.47	13.74
Fe2O3	11.98	10.90	12.17	11.08	10.53	10.96	10.15	7.38	10.15	10.32
MnO	0.14	0.14	0.17	0.18	0.12	0.13	0.14	0.07	0.10	0.15
MgO	4.16	5.53	17.46	6.65	4.37	2.96	5.54	3.04	4.83	5.19
CaO	7.77	7.90	8.06	6.54	2.66	4.01	6.97	5.19	5.56	6.19
Na2O	3.40	3.09	0.19	3.61	4.30	4.23	2.74	3.74	3.20	2.66
K2O	1.70	0.82	0.11	1.53	2.92	2.24	2.52	3.59	0.68	1.72
P2O5	0.118	0.081	0.039	0.077	0.125	0.657	0.540	0.595	0.690	0.590

TOTAL:	100.06	100.83	99.59	99.56	100.81	100.86	101.68	100.88	99.19	100.31
LOI	1.25	2.08	4.95	3.71	5.88	3.03	4.08	4.07	4.53	4.35

BA	457	262	16	235	376	1054	987	1481	482	703
NB	3.6	3.4	1.2	5.3	3.5	33.3	15.1	24.0	16.1	17.7
ZR	97	78	34	63	84	750	288	421	283	311
Y	22.3	20.6	10.4	16.1	19.9	85.7	50.3	52.1	48.2	49.1
SR	400	498	20	90	173	328	741	367	664	545
RB	66.6	25.2	7.6	43.4	80.2	64.6	67.9	66.1	15.5	41.7
ZN	84	73	76	80	70	101	107	88	115	110
CU	19	88	45	65	113	10	37	12	18	23
NI	108	123	516	180	163	25	165	103	144	124
CO	58	57	88	64	58	29	46	36	46	40
CR	31	186	2618	336	437	17	421	302	457	432
V	221	206	178	224	215	158	196	159	201	187
LA	12	5*	0	5*	15	113	58	75	49	37
CE	14	19	8*	20	19	169	127	166	109	124
ND	12	10	0	10	13	84	61	81	58	61

SAMPLE I	PR-480	PR-481	PR-483	PR-484	PR-485	PR-486	PR-487	PR-489	PR-491	PR-492
SiO2	56.72	56.53	58.06	55.32	55.20	55.80	59.85	59.56	59.29	59.82
TiO2	1.486	1.426	1.241	1.469	1.513	1.377	1.435	1.381	1.376	1.744
Al2O3	15.33	14.73	13.96	13.62	13.78	12.02	13.65	13.96	14.02	13.68
Fe2O3	9.98	11.59	10.12	11.03	11.78	11.99	8.34	9.55	9.92	10.17
MnO	0.11	0.12	0.13	0.15	0.16	0.18	0.13	0.13	0.13	0.14
MgO	4.39	7.90	4.04	6.05	5.49	8.20	3.75	3.90	3.90	2.87
CaO	6.04	3.92	6.76	7.67	6.91	6.66	5.56	6.20	5.45	6.12
Na2O	3.31	2.92	2.05	1.90	2.88	1.93	3.50	6.62	2.95	2.75
K2O	1.98	0.17	3.10	2.14	1.57	1.24	1.03	2.06	2.34	1.91
P2O5	0.665	0.693	0.544	0.655	0.718	0.625	0.752	0.602	0.624	0.801

TOTAL:	100.48	98.72	100.20	101.03	99.68	99.21	100.41	98.78	100.31	100.66
LOI	5.45	3.64	5.00	3.10	5.87	2.93	2.92	3.51	3.73	5.82

BA	1025	83	1202	929	992	477	788	991	987	913
NB	19.6	14.6	16.8	14.8	14.5	14.7	16.8	17.2	19.6	25.6
ZR	349	299	310	276	273	244	318	350	354	426
Y	51.4	45.1	48.6	48.5	51.1	45.4	53.4	51.4	51.0	62.4
SR	488	291	537	529	460	400	955	837	1045	607
RB	49.1	3.2	78.3	50.1	33.1	24.3	23.0	45.4	38.8	39.5
ZN	124	116	111	114	126	123	78	104	109	104
CU	21	9	34	37	44	34	39	30	35	39
NI	92	165	92	157	139	220	68	78	80	50
CO	47	57	45	50	53	58	35	41	39	35
CR	257	546	241	462	415	936	126	252	276	253
V	195	218	169	199	223	195	183	170	168	233
LA	58	48	57	52	52	44	65	59	61	77
CE	142	122	124	106	113	98	136	128	133	163
ND	66	62	60	55	62	51	70	62	63	81

SAMPLE I	AR-494	AR-495	AR-496	AR-497	AR-498	AR-500
SiO2	51.80	52.68	53.74	51.76	56.77	56.74
TiO2	1.775	1.276	1.201	1.210	1.146	1.018
Al2O3	14.13	16.37	15.35	13.75	13.16	14.29
Fe2O3	13.37	12.73	12.23	13.67	11.41	11.16
MnO	0.17	0.11	0.13	0.16	0.14	0.18
MgO	5.31	5.06	4.07	5.78	5.61	4.96
CaO	7.63	7.49	9.17	7.76	5.64	5.86
Na2O	2.47	2.24	2.46	3.34	3.52	3.29
K2O	2.47	1.78	1.36	0.34	0.37	2.29
P2O5	0.879	0.266	0.264	0.229	0.233	0.220

TOTAL:	100.87	99.13	99.54	100.12	98.70	100.38
LOI	10.47	8.60	12.88	7.72	5.06	3.36

BA	580	416	175	97	202	705
NB	16.9	8.4	7.8	7.4	8.1	8.1
ZR	300	192	182	181	189	159
Y	54.8	32.1	28.7	24.2	32.0	28.0
SR	555	403	196	242	555	536
RB	88.5	58.8	48.2	10.9	52.7	41.8
ZN	132	117	97	106	105	95
CU	41	182	129	137	95	88
NI	166	210	157	213	166	99
CO	30	71	63	64	69	56
CR	426	61	51	213	56	183
V	233	265	258	222	208	233
LA	51	21	27	16	25	22
CE	118	53	57	48	50	50
ND	62	30	29	22	27	27

TABLE 2.5 X-ray counting statistics - LLD and CE - for each batch of samples processed

		Nb		Zr		Y		Sr		Rb	
BATCH	N	LLD	CE	LLD	CE	LLD	CE	LLD	CE	LLD	CE
1	40	2.42	.71	1.81	.76	2.19	.68	1.95	1.08	2.11	.69
2	48	2.42	.72	1.81	.80	2.20	.69	1.96	1.06	2.12	.70
3	238	.76	.23	.57	.25	.69	.22	.61	.36	.65	.22
		Co		Cr		V		Zn		Cu	
BATCH	N	LLD	CE	LLD	CE	LLD	CE	LLD	CE	LLD	CE
1	40	2.52	.83	2.56	1.31	3.34	1.41	2.16	1.10	2.45	1.04
2	48	2.56	.84	2.64	1.31	3.58	1.50	2.45	1.17	2.68	1.06
3	238	2.46	.82	2.54	1.41	3.27	1.45	2.67	1.26	2.80	1.17
		Ni		La		Ce		Nd		Ba	
BATCH	N	LLD	CE	LLD	CE	LLD	CE	LLD	CE	LLD	CE
1	40	3.26	1.59	5.84	1.56	11.58	2.98	6.47	1.65	14.0	6.0
2	48	3.46	1.61	6.55	1.62	12.32	3.11	6.64	1.67	14.7	5.8
3	238	3.38	1.63								

X-ray counting statistics are calculated as follows:

Counting error

Counting error is the standard deviation of the net peak count rate, i.e.

$$\sigma = \sqrt{\frac{R_P}{T_P} - \frac{R_B}{T_B}}$$

where R_P, R_B = count rates at peak and background positions

T_P, T_B = counting times at peak and background positions

Expressed in terms of concentration, this is:

$$\text{C.E.} = \frac{\sqrt{\frac{R_p}{T_p} - \frac{R_B}{T_B}}}{R_p - R_B} \cdot \text{concentration}$$

When $R_p \rightarrow R_B$, counting error $\rightarrow \infty$ so the concept of a lower limit of determination must be introduced.

Lower limit of determination

Due to the Gaussian distribution of X-ray emission intensity, the existence of a peak above background can only be confirmed if it has a count rate in excess of the mean background count rate, plus 3 standard deviations of the background count rate. Therefore,

$$\text{Detection limit} = 3 \sqrt{\frac{R_B}{T_B}}$$

Expressed in terms of concentration, this is

$$\text{D.L.} = \frac{3}{M} \sqrt{\frac{R_B}{R_T}}$$

Where M = counts per second per unit concentration.

At the detection limit counting errors are large, so the lower limit of determination (L.L.D.) is defined as twice the D.L., i.e.

$$\text{L.L.D.} = \frac{6}{M} \sqrt{\frac{R_B}{T_B}}$$

3. GEOCHEMISTRY

3.1 Classification

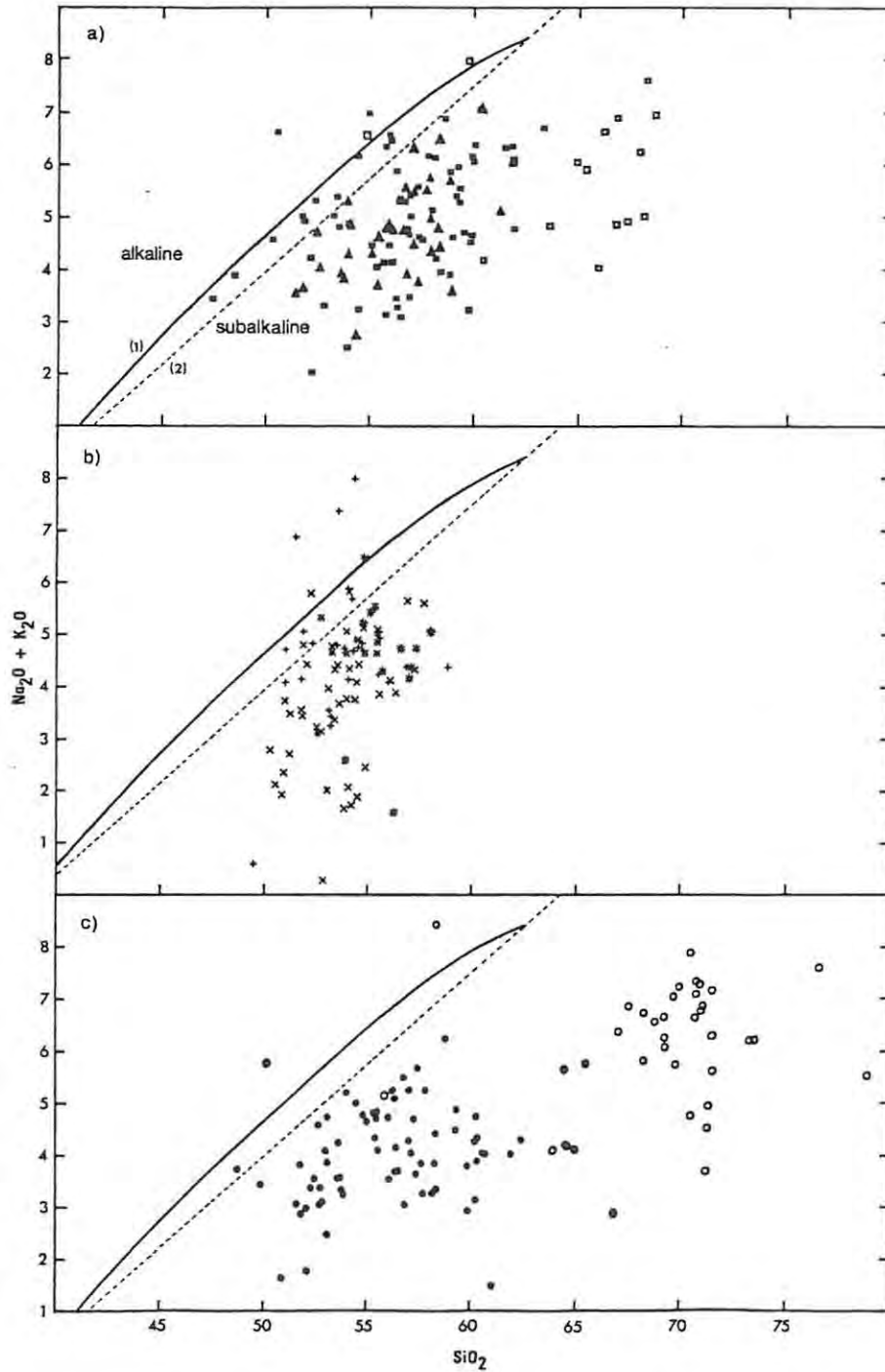
The classification scheme of Irvine and Baragar (1971) was selected to aid the classification of the Witwatersrand triad volcanic rocks for three important reasons:

- (1) The scheme is based on 2500 analyses of Precambrian (therefore applicable to the current study) and Cenozoic rocks.
- (2) The scheme can be applied to metamorphosed as well as unaltered volcanic rocks, as it is based on a chemical, rather than a mineralogical or modal composition.
- (3) Classification into the three main series - the tholeiitic basalt, calc-alkali and alkali olivine basalt series - has no genetic implications, as their boundaries appear to be natural divisions.

The initial distinction in Irvine and Baragar's classification scheme is that between alkaline and subalkaline rocks. For this purpose, a weight percent plot of $(\text{Na}_2\text{O} + \text{K}_2\text{O})$ against SiO_2 is employed, as shown in Figure 3.1. The curved line, marked (1), is Irvine and Baragar's dividing line between the alkaline and subalkaline fields. They found this line to be a more effective divider than that of MacDonald (1968) (marked (2)), which had been widely used.

The Witwatersrand triad volcanic rocks have been altered as a result of low-grade greenschist facies metamorphism. Variation diagrams (Figures 3.9a and 3.11a) show that a large amount of scatter is exhibited by Na_2O and K_2O concentrations, suggesting that these two oxides were fairly mobile. While it may be true that individual samples have lost or gained Na or K, the average compositions of formations as a whole should provide a good estimate of the original concentrations. Systematic trends are still evident, as well as clear compositional breaks between formations, e.g., between the Klipriviersberg Group and Platberg Group, and again at the top of the latter. Na_2O and K_2O concentrations thus seem to have remained basically unchanged overall and average concentrations can be regarded as valid for each formation, and can thus be used in the following classification diagram.

Figure 3.1 Alkalies vs silica diagram with dividing lines of (1) Irvine and Baragar (1971) and (2) MacDonald (1968)
a) Platberg Group and Allanridge Formation
b) Klipriviersberg Group
c) Dominion Group and Crown Formation



see pages 62 and 72 for symbols

From Figure 3.1 it is clear that almost all the Witwatersrand triad volcanic rocks are subalkaline. In Figure 3.1a, only two samples from the Makwassie Formation (\square), two from the Rietgat-Goedgenoeg Formations (\blacksquare) and none from the Allanridge Formation (\blacktriangle) plot in the alkaline field. In Figure 3.1b, one Alberton (*), three Orkney (+) and one Loraine-Edenville sample (x) plot in the alkaline field, while in Figure 3.1c, one Dominion Group porphyry and one Crown Formation sample are alkaline. Thus 96.6% of all the samples are subalkaline.

Having established the subalkaline nature of these rocks, the next subdivision is into either the tholeiitic basalt series, or the calc-alkali series. AFM triangular plots, with $A = Na_2O + K_2O$, $F = FeO + 0.8998 Fe_2O_3$, and $M = MgO$, all in weight percent, are shown in Figures 3.2a (Ventersdorp Supergroup) and 3.2b (Dominion Group and Crown Formation).

The Ventersdorp Supergroup samples straddle the dividing line, with the following proportion of samples plotting in the tholeiitic field:

Allanridge	- 51%	
Rietgat	- 35%	
Makwassie	- 29%	
Goedgenoeg	- 33%	Total: 40%
Loraine & Edenville	- 48%	
Orkney	- 45%	
Alberton	- 21%	

Figure 3.2a appears, at first glance, to suggest that all formations of the Ventersdorp Supergroup have calc-alkaline affinities. This needs to be examined more closely, however, as Irvine and Baragar (1971) themselves imply that the AFM plot does not satisfactorily separate tholeiitic from calc-alkaline types if the rocks are basic. In such instances, it is more important to look at the trend of samples from a particular formation. The Loraine-Edenville samples, for instance, display a clear tholeiitic trend towards FeO, even though most of the samples lie on the calc-alkaline side of the boundary line. The trends in some of the formations are less obvious, although the Dominion Group porphyries and basic rocks, and the Crown Formation are all clearly tholeiitic.

Figure 3.2 AFM triangular plots for
a) Ventersdorp Supergroup
b) Dominion Group and Crown Formation

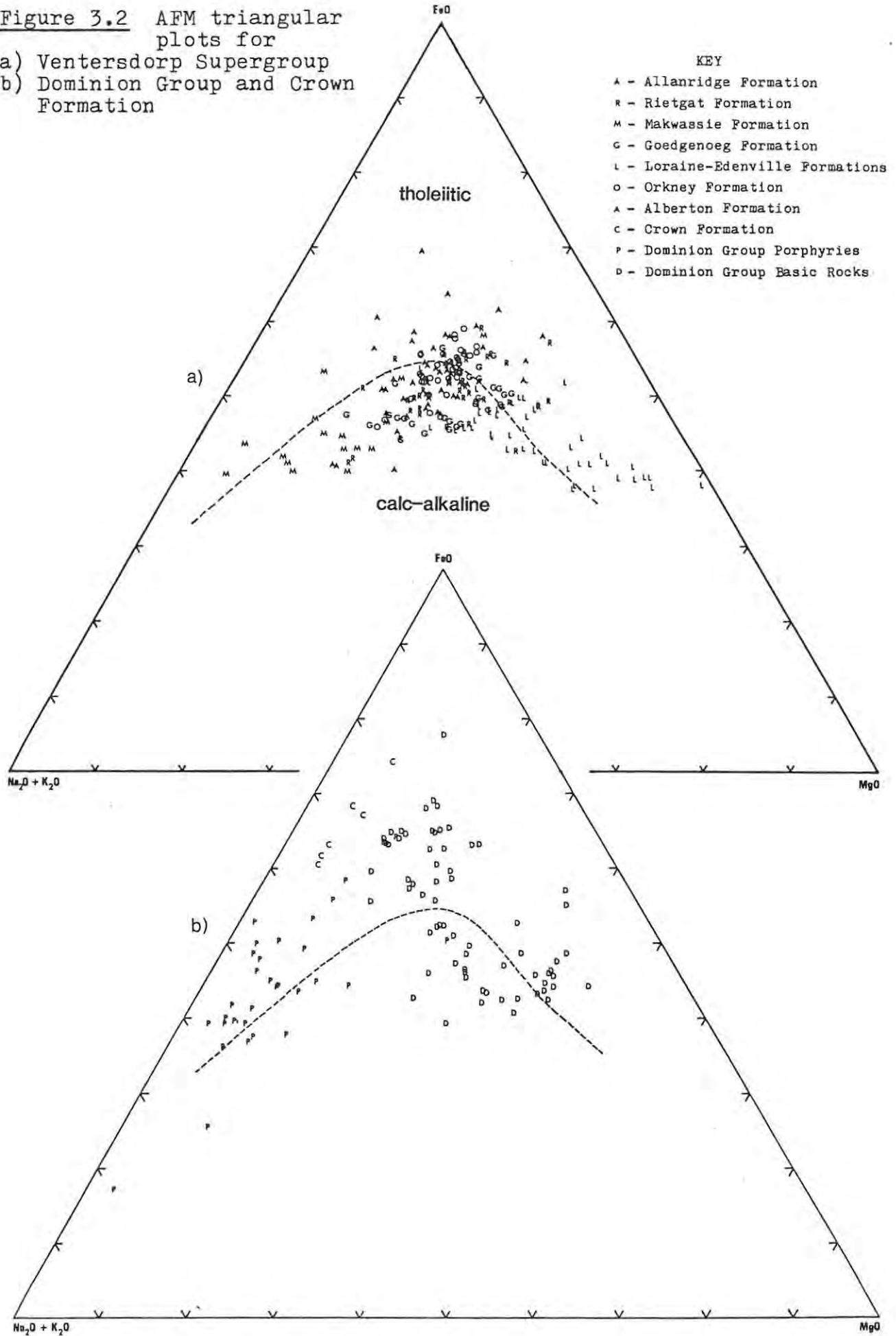
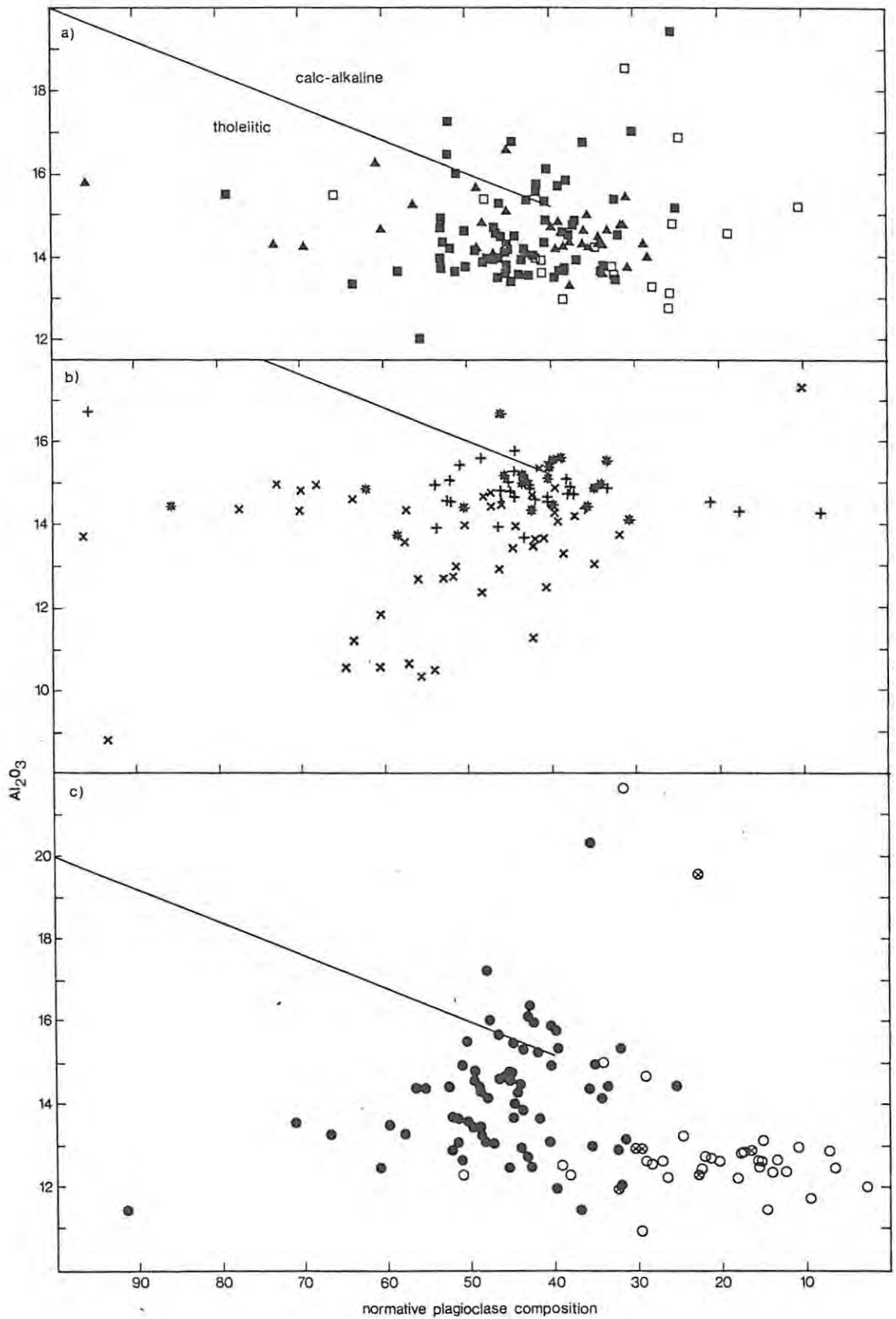


Figure 3.3 Al_2O_3 vs normative plagioclase composition (after Irvine and Baragar, 1971) for
a) Platberg Group and Allanridge Formation
b) Klipriviersberg Group
c) Dominion Group and Crown Formation



The most prominent chemical difference between the more basic members of typical calc-alkali and tholeiitic series of Cenozoic age is their alumina content (Irvine and Baragar, 1971): the calc-alkali basalts and andesites are generally high-alumina types, containing 16 to 20% Al_2O_3 , whereas their tholeiitic counterparts have only 12 to 16%. They suggested that a plot of Al_2O_3 against normative plagioclase composition ($100 An/(An + Ab + 5/3 Ne)$) provides a more significant separation of the two series than the AFM plot. Figures 3.3a (Platberg Group and Allanridge Formation), 3.3b (Klipriviersberg Group) and 3.3c (Dominion Group and Crown Formation) show that the following proportions of samples plot below the dividing line (or its projection) in the tholeiitic field:

Allanridge	- 86%
Rietgat & Goedgenoeg	- 79%
Makwassie	- 71%
Lorraine	- 96%
Orkney	- 77%
Alberton	- <u>63%</u>
Total Ventersdorp average	- <u>81%</u>
Crown	- 83%
Dominion porphyries	- 84%
Dominion basic lavas	- 83%

The increase in the proportion of tholeiitic samples towards the top of the Klipriviersberg Group is again evident, and in this plot 81% (as opposed to 40% in Figure 3.2) of the Ventersdorp samples plot in the tholeiitic field. The three underlying units are again conclusively tholeiitic.

Figures 3.4a, b and c are plots of normative colour index ($O1 + Opx + Cpx + Mt + Il + Hm$) against normative plagioclase composition, in which rock names are defined. The results of these plots are summarised as follows, with b = basalt, ta = tholeiitic andesite, a = andesite, d = dacite and r = rhyolite:

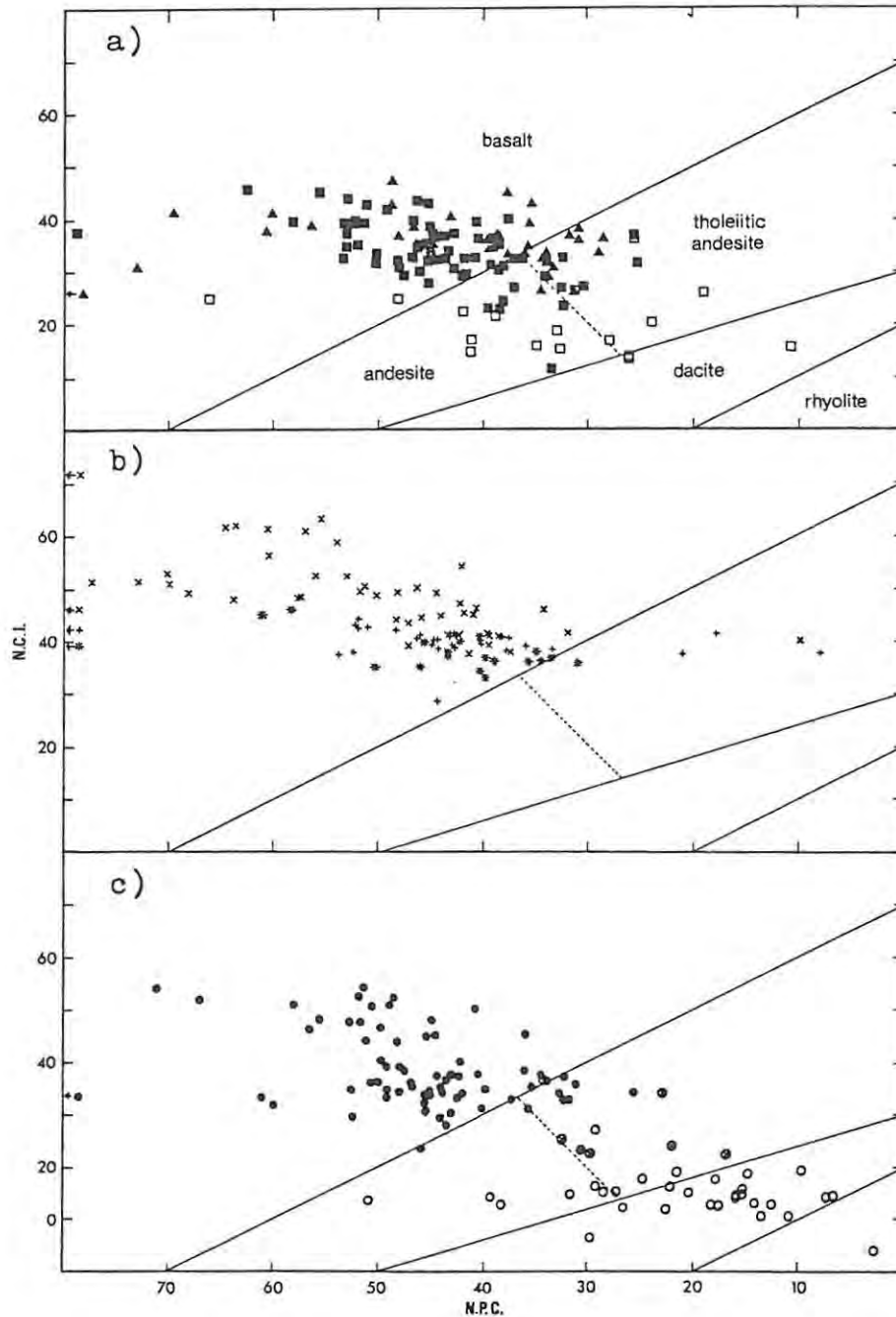
	<u>b</u>	<u>ta</u>	<u>a</u>	<u>d</u>	<u>r</u>	<u>name</u>
Allanridge	66%	31%	3%			b-ta
Rietgat & Goedgenoeg	75%	12%	13%			b
Makwassie	12%	29%	41%	18%		ta-a
Lorraine & Edenville	98%	2%				b
Orkney	91%	9%				b
Alberton	<u>95%</u>	<u>5%</u>	<u>—</u>	<u>—</u>	<u>—</u>	b
	<u>77%</u>	<u>13%</u>	<u>8%</u>	<u>2%</u>	<u>0%</u>	
Crown		100%				ta
Dominion porphyries	13%	13%	19%	63%	3%	d
Dominion basic lavas	88%	9%	3%			b

It must be stressed that names like dacite and andesite are used here without tectonomagmatic connotations. Using these definitions, the Allanridge, Rietgat, Goedgenoeg, Lorraine, Orkney and Alberton Formations, as well as the Dominion Group basic lavas, are all overwhelmingly basaltic. The Dominion Group porphyries are predominantly dacitic, with an andesitic component, the Crown lavas are all tholeiitic andesites, while the Makwassie samples spread over almost the entire range, from basalts through to dacites, with most of the samples being andesitic.

The (Na₂O + K₂O) versus SiO₂ classification diagram of Cox et al., (1979) has also been used, as shown in Figure 3.5, with the following results (abbreviations as above, but also ba = basaltic andesite, h = hawaiite, tb = trachybasalt, ta = trachyandesite and u = unclassified):

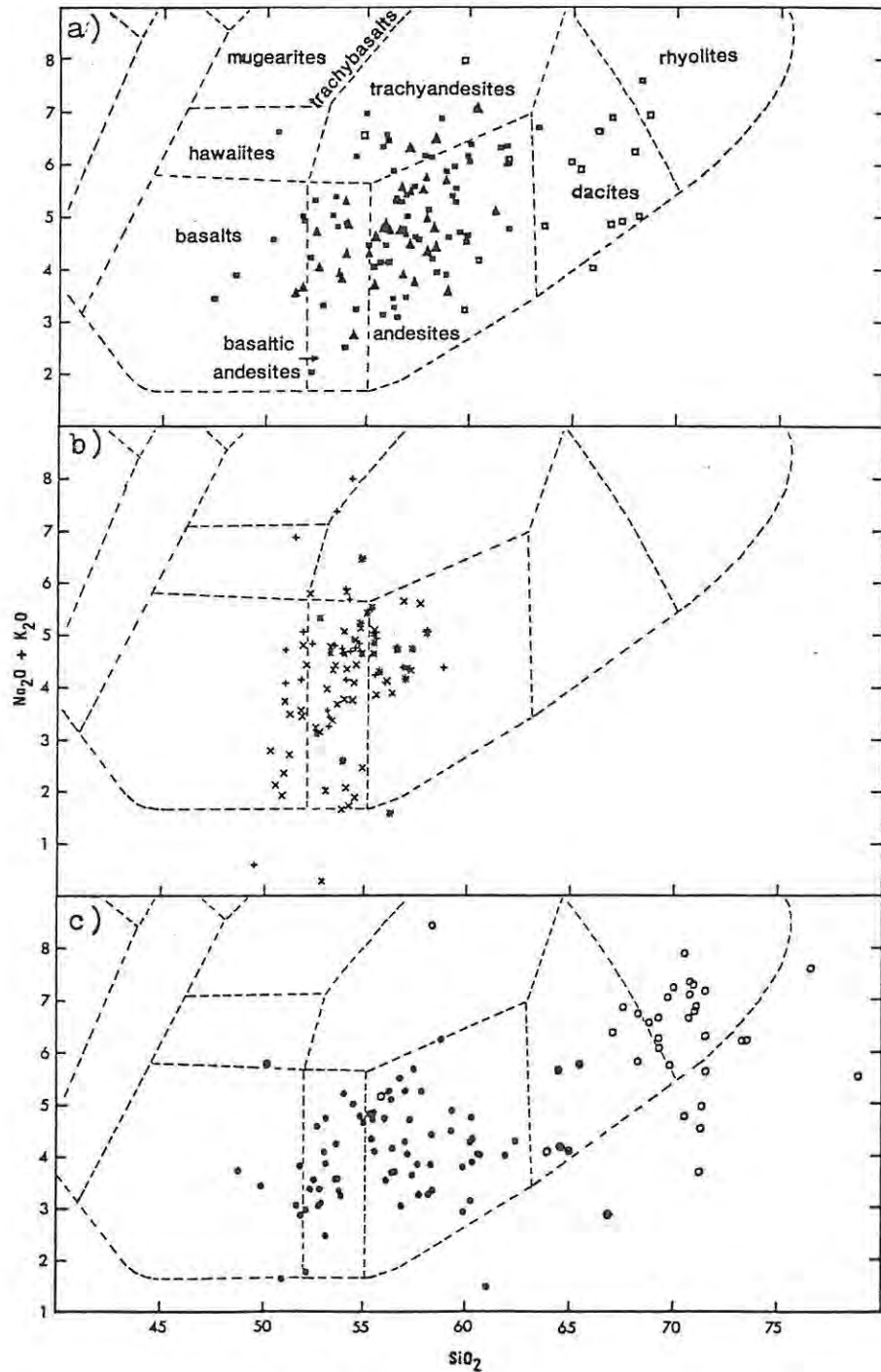
	<u>b</u>	<u>ba</u>	<u>a</u>	<u>d</u>	<u>r</u>	<u>h</u>	<u>tb</u>	<u>ta</u>	<u>u</u>	<u>name</u>
Allanridge	6%	23%	63%					9%		a
Rietgat & Goedgenoeg	7%	15%	63%	1%		1%		12%		a
Makwassie			18%	53%	12%			12%	5%	d
Lorraine	26%	52%	15%					4%	2%	b-a
Orkney	13%	47%	22%			3%	6%	6%	3%	ba-a
Alberton	<u>—</u>	<u>42%</u>	<u>47%</u>	<u>—</u>	<u>—</u>	<u>—</u>	<u>—</u>	<u>5%</u>	<u>5%</u>	ba-a
	<u>11%</u>	<u>30%</u>	<u>42%</u>	<u>5%</u>	<u>1%</u>	<u>1%</u>	<u>1%</u>	<u>8%</u>	<u>2%</u>	
Crown				67%		17			17%	d
Dominion porphyries			3%	16%	50%			3%	28%	r
Dominion basic lavas	9%	30%	55%						1%	ba-a

Figure 3.4 Normative colour index versus normative plagioclase composition (after Irvine and Baragar, 1971) for
a) Platberg Group and Allanridge Formation
b) Klipriviersberg Group
c) Dominion Group and Crown Formation



Under the Cox et al., (1979) classification, therefore, the names given to the individual units in all cases imply a slightly more acid rock type than those defined for the same unit by the Irvine and Baragar (1971) scheme. For example, the Dominion porphyries are dacites by Irvine and Baragar's definition, but rhyolites according to Cox et al..

Figure 3.5 Nomenclature of volcanic rocks - alkalis versus SiO_2 plot (after Cox et al., 1979) for
a) Platberg Group and Allanridge Formation
b) Klipriviersberg Group
c) Dominion Group and Crown Formation



The final classification phase of Irvine and Baragar is that for distinguishing between potassium-rich, "average", or potassium-poor varieties of rock. A series of triangular An-Ab-Or plots is illustrated in Figure 3.6. The percentage of samples from each formation plotting within each of these three fields is as follows:

	<u>K-poor</u>	<u>average</u>	<u>K-rich</u>
Allanridge	26%	46%	28%
Rietgat & Goedgenoeg	15	37	48
Makwassie	24	29	47
Lorraine & Edenville	9	39	52
Orkney	22	52	26
Alberton	37	37	26
Crown	83	17	0
Dominion porphyries	16	53	31
Dominion basic lavas	<u>22</u>	<u>56</u>	<u>22</u>
Total average	28%	41%	31%

For the Dominion Group (see Figure 3.6c), over half the samples of both porphyries and basic rocks lie within the "average" field. The remainder of the basic rocks are spread evenly throughout the K-poor and K-rich fields, while the rest of the porphyries lie mostly in the K-rich field, but with a fair number in the K-poor field. The Crown Formation samples are almost all K-poor, with one sample in the "average" field.

Figure 3.6b shows that the Klipriviersberg Group becomes progressively more potassic towards the top. The majority of the Alberton samples are K-poor to "average", most of the Orkney samples are "average", while K-rich samples predominate in the Lorraine and Edenville Formations.

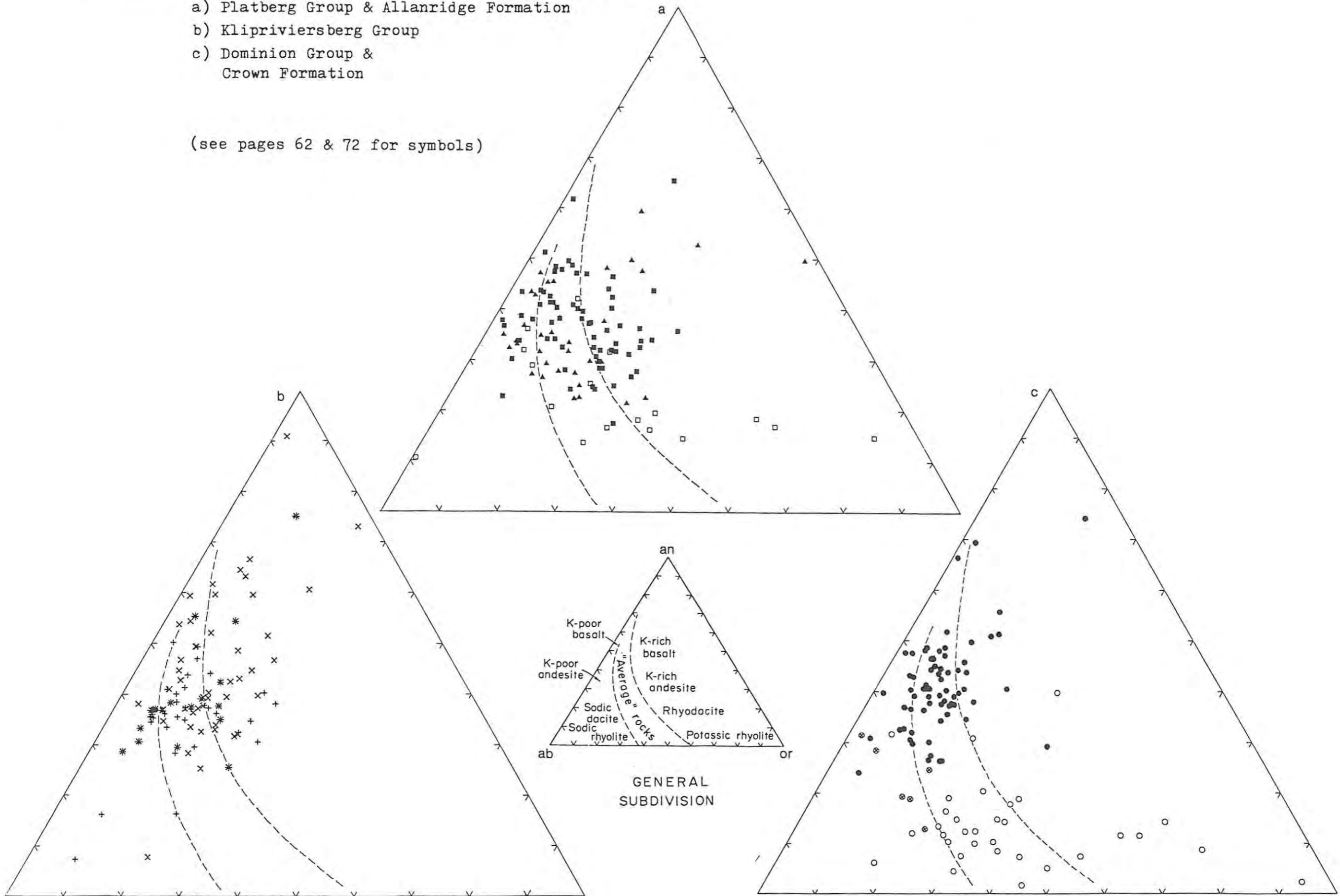
The Platberg Group is represented mostly by K-rich rocks, but the other two fields still accommodate many of them. The Allanridge samples also show a complete spread, most of the samples being "average".

Because, as has been suggested earlier in section 3.1, K appears to have been mobile during low-grade metamorphism, the spread of samples across the three fields in Figure 3.6 should be studied carefully. The equal K-poor and K-rich scatter about the K-average field for Allanridge, Orkney and Dominion basic samples may indicate that these samples were originally K-average, but alteration has spread them out. Similarly, the smearing out of samples in other formations may be the result of alteration.

Figure 3.6 An-Ab-Or projection (after Irvine & Baragar, 1971)

- a) Platberg Group & Allanridge Formation
- b) Klipriviersberg Group
- c) Dominion Group & Crown Formation

(see pages 62 & 72 for symbols)



In conclusion, the Witwatersrand triad volcanic rocks examined are virtually exclusively subalkaline. Both AFM and Al_2O_3 vs normative plagioclase composition (NPC) plots show that the Dominion Group basic lavas and porphyries are predominantly tholeiitic, as is the Crown Formation. The Klipriviersberg Group samples plot in both the tholeiitic and calc-alkaline fields on an AFM diagram. The calc-alkaline nature predominates marginally in the upper formations and more strongly in the lower formations. Irvine and Baragar (1971), however, recommended the use of the Al_2O_3 vs NPC plot for basaltic rocks, and on this diagram the majority of the Klipriviersberg samples plot in the tholeiitic field. They are therefore classified as tholeiitic with calc-alkaline tendencies, the latter increasing towards the base.

Similar reasoning requires that the Rietgat-Goedgenoeg and Makwassie rocks also be classified as tholeiitic with calc-alkaline tendencies, while the Allanridge samples are mostly tholeiitic, according to both plots.

Irvine and Baragar's (1971) diagram of normative colour index vs normative plagioclase composition classifies as basaltic the Dominion Group basic lavas, the Klipriviersberg Group formations, and the Rietgat-Goedgenoeg Formations. The Allanridge Formation samples are basalts to tholeiitic andesites, those of the Crown Formation are tholeiitic andesites, the Dominion Group porphyries are dacitic with an andesitic to tholeiitic andesitic tendency, while the Makwassie Formation samples are tholeiitic andesites to andesites, with tendencies toward both the basaltic and dacitic extremities.

The $Na_2O + K_2O$ vs SiO_2 plot of Cox et al. (1970) names the majority of the Ventersdorp and Dominion basic rocks as andesites and basaltic andesites. This is based almost entirely on their SiO_2 content and does not consider Al_2O_3 values at all, which, in most of these rocks, is below 16%. Wyatt (1976) reviewed the common usage of the name "andesite" for the Ventersdorp basic rocks and concluded that the Al_2O_3 content of the Klipriviersberg lavas he examined was too low to justify the name "andesite", although their SiO_2 content was acceptable. For similar reasons, the Irvine and Baragar (1971) nomenclature is preferred for the current study.

Finally, the potassium content of all the formations (except the Crown Formation) ranges from K-poor, through "average" to K-rich, with the "average" field dominating in the Dominion Group rocks, the Alberton, Orkney and Allanridge Formations. The Loraine and Edenville, Makwassie, Rietgat and Goedgenoeg Formations tend to be more K-rich, while the Crown Formation is K-poor.

3.2 A Comparison Between Rocks From the Klipriviersberg Group and Komatiitic Lavas

The Westonia Formation at the base of the Klipriviersberg Group contains a succession of volcanic rocks, termed the "Meredale Member" by Wyatt (1976), which are considerably more basic than those constituting the upper formations. McIver, Cawthorn and Wyatt (1982) suggested that this unit might represent the youngest komatiitic sequence in South Africa, and compared these high-MgO rocks with typical analyses from komatiitic suites having a comparable MgO content. The only differences they found were higher TiO₂, Zr and Sr values in the Meredale rocks.

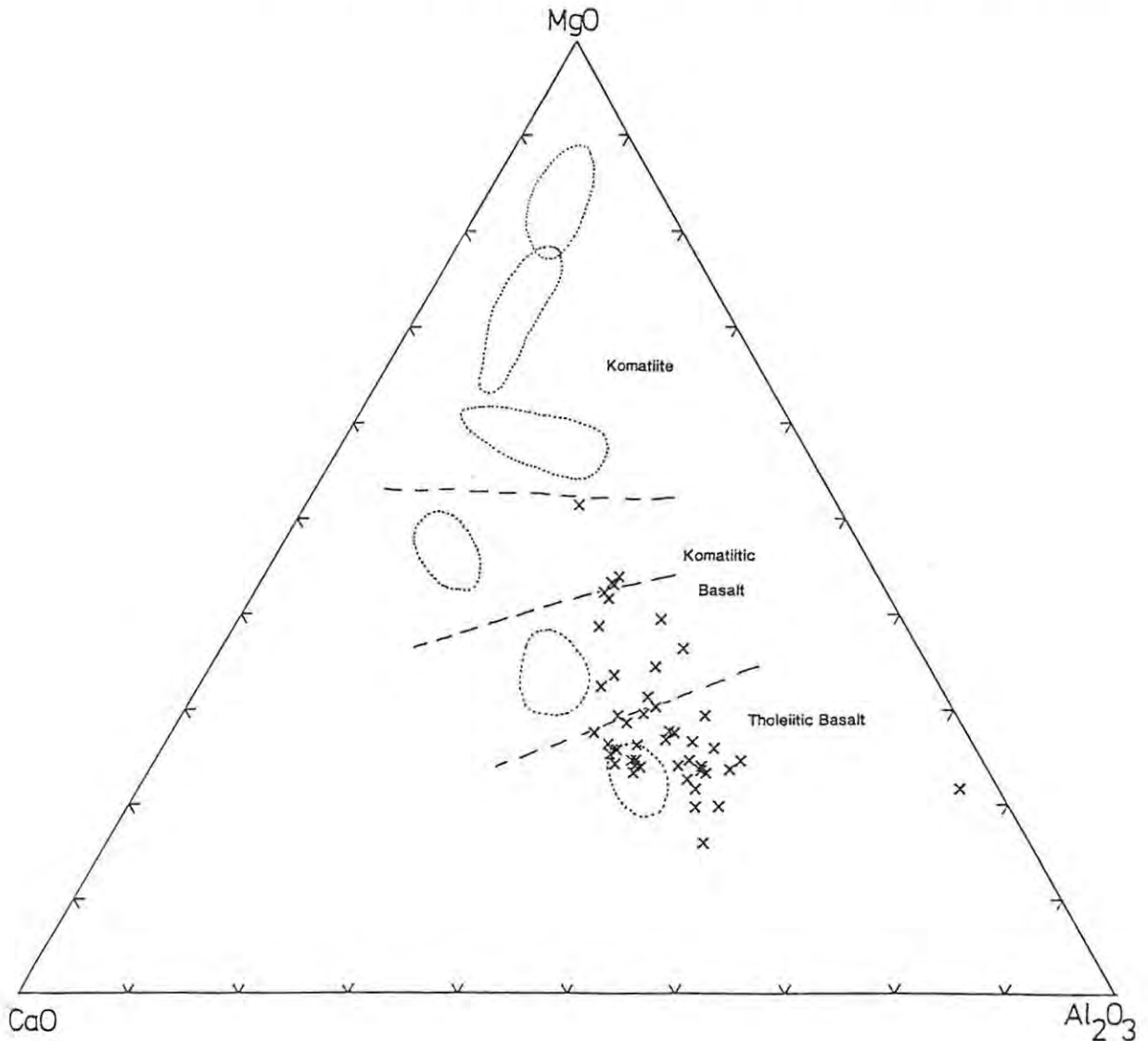
A number of high-MgO rocks has been analysed in the current study, but these samples occur near the top of the Klipriviersberg Group, in the Loraine-Edenville Formations. Figure 3.11a, which illustrates MgO variation with height in borehole WS-5, shows the existence of three sharp spikes defining narrow (<50m) high-MgO horizons, culminating in a steady increase in MgO in the top 180 metres of the unit. There are 21 samples with MgO content greater than 8.5%, which is about the value above which the high MgO peaks begin to manifest themselves, and values range up to 17.46% MgO.

Figure 3.7, a triangular plot of MgO-CaO-Al₂O₃, shows the original fields of komatiites and associated rocks as defined by Viljoen and Viljoen (1969) (finely dotted lines), as well as data breaks delineated by original and newer data from the Barberton and other South African greenstone belts (dashed lines).

Samples from the Loraine-Edenville Formations have been plotted on this diagram, which shows that 14 samples lie within the field of what Viljoen et al. (1982) now term komatiitic basalts, while the remaining

Figure 3.7 MgO-CaO-Al₂O₃ triangular plot (after Viljoen and Viljoen, 1982)

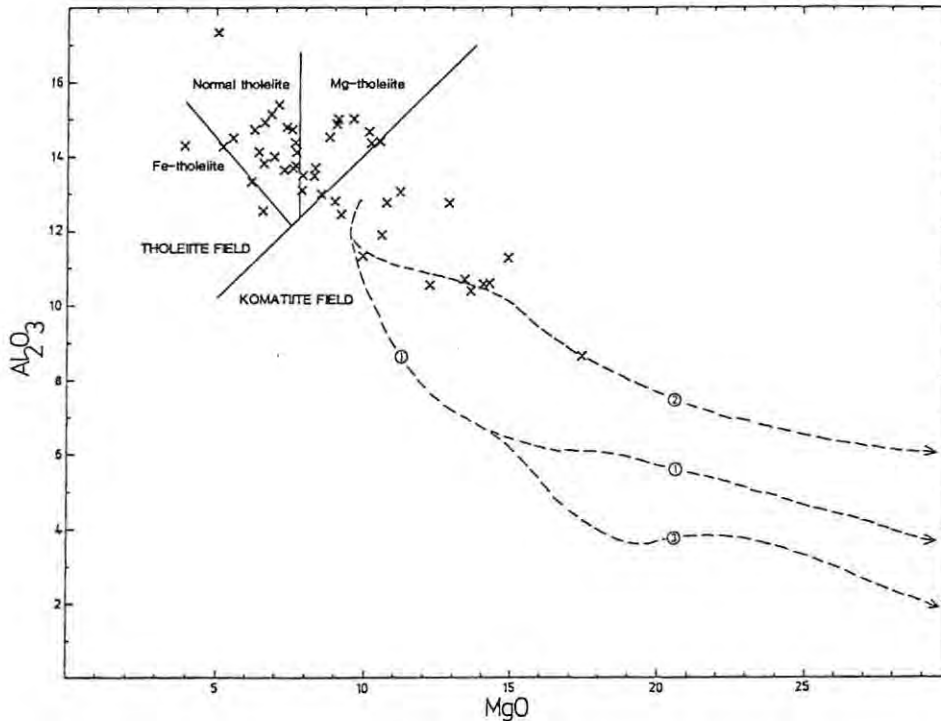
- ... = original fields of komatiites and associated rocks, as defined by Viljoen and Viljoen (1969)
- = data breaks delineated by original and new data from the Barberton and other South African greenstone belts



32 samples plot in the tholeiitic basalt field. CaC:Al₂O₃ ratios in the latter area compare well with those of the Viljoen et al. samples, while those in the komatiitic basalt field have slightly higher ratios. Some of the more recently analysed samples of Viljoen et al., however, have similar ratios.

The Loraine-Edenville samples were also plotted on the Al₂O₃ vs MgO diagram of Viljoen et al. (1982) (see Figure 3.8). The tholeiitic field is divided into three parts, delineating Fe-rich, normal and Mg-rich tholeiites. Of the Klipriviersberg samples, four plot in the

Figure 3.8 Al_2O_3 vs MgO plot with tholeiite and komatiite fields of Viljoen and Viljoen (1982)
1 = Main trend of aluminium depletion
2 = Secondary trend of aluminium undepletion
3 = Possible minor trend of pronounced aluminium depletion



Fe-tholeiite field, sixteen amongst the normal tholeiites, 11 with the Mg-tholeiites and 15 in the komatiitic field. The majority of the komatiitic samples lie directly on the "trend of aluminium undepletion" as defined by Viljoen *et al.* (1982).

Having shown that a number of Loraine-Edenville samples plots in komatiitic fields, it is advisable at this point to examine a more formal definition of a komatiite and a komatiitic basalt. Arndt and Nisbet (1982) define komatiites as ultramafic volcanic rocks with more than 18% MgO (anhydrous basis), and which form the ultramafic portion of a magmatic suite - the komatiitic suite or series - which also includes mafic volcanic rocks called komatiitic basalts. The chemistry of the latter is characterised by high MgO, Ni and Cr, low alkalis, TiO_2 and Fe:Mg ratio, high SiO_2 at the given MgO value, and CaO: Al_2O_3 ratios greater than 0.8-1.0. They emphasise that these

features are characteristic, but not necessarily diagnostic, as it is only their link with true komatiites which permits their being called komatiitic basalts. Under other circumstances, they would simply be called Mg-rich tholeiites, or some such term. Cameron and Nisbet (1982), however, suggested that the term "komatiitic basalt" could be used if mineralogical traits are sufficiently similar, without implying that the komatiitic basalts have fractionated from komatiites.

Of the 46 Loraine-Edenville samples, 17 lie within the komatiitic basalt field in either Figure 3.7 or 3.8, or both. The MgO values of these 17 are all greater than 9.00%, Ni ranges between 186 and 516 ppm, Cr from 584 to 2618 ppm, Na₂O from 0.19 to 2.39%, K₂O from 0.11 to 1.95% and TiO₂ from 0.400 to 0.533%. CaO:Al₂O₃ ratios are generally lower than strictly required, with only 8 samples having a value greater than 0.80, while 7 others lie between 0.70 and 0.79, and the remainder below this. Fe:Mg ratios range from 0.81 to 1.40, while SiO₂ values lie between 50.28 and 55.60%.

In the light of the foregoing discussion, if it is accepted that the Meredale Member contains komatiites, it can be concluded that the following samples (in decreasing order of MgO) may be classed as komatiitic basalts: KL-468, 158, 157, 159, 135, 156, 146, 205, 155, 154, 141, 43, 44, 90, 140, 204 and 382. However, until the case of McIver et al. (1982) is conclusively proved, it is preferable to designate these rocks Mg-rich tholeiites. Further sampling in the vicinity of these samples is recommended to ascertain whether samples with higher MgO values, comparable to true komatiites, are present.

3.3 Statistics

Table 3.1a contains the means, standard deviations and coefficients of variation (CV's) for all major element oxides for each formation. The oxide values are expressed in weight percent. Table 3.1b lists the means, standard deviations and CV's for the trace elements, values for which are given in parts per million. The following formulae were used in the calculations :

$$\text{Mean} = \bar{X} = \frac{\sum x}{N}, \text{ and}$$

$$\text{Standard deviation} = s = \sqrt{\frac{\sum (x - \bar{X})^2}{N - 1}}$$

where x = element values, and
 N = number of samples.

In Table 3.1b, N represents the maximum number of samples used in the calculations for that particular formation. The actual N used may be less than that given, because certain samples have trace element values below the lower limit of determination (LLD). In such a situation, the sample has been omitted from the calculations for that particular element, and the resulting mean will therefore be slightly higher than the true mean. The affected elements, together with the number of samples below the LLD, in brackets, are Nb (9), Cr (14), La (29), Ce(19) and Nd (28), giving a total of 99 unusable analyses, or 2.02 percent of all the trace element determinations.

In an attempt to reduce the number of unusable analyses, it was decided to distinguish between values which lay below the LLD and those below the detection limit (DL). The DL is half the LLD and is not usually used as the counting errors at the DL are large. For the purpose of the present calculations, however, it was decided that it would be expedient to use those values (marked in Table 2.4 with an *) which were below the LLD but above the DL, in order to obtain more representative statistics. This resulted in a reduction in the number of unusable analyses, to the following: Nb (2), Cr (10), La (6), Ce (7) and Nd (11), a total of 36, or 0.74 percent of the trace element analyses.

The coefficients of variation were included for the following reason. The actual variation or dispersion of an element as determined from the standard deviation is called the absolute dispersion. However, a variation or dispersion of 2 weight percent relative to an average of, say, 70 weight percent is quite different from the same variation of 2 weight percent relative to 5 weight percent. A measure of this effect is provided by the relative dispersion, which is defined as

$$\text{Relative Dispersion} = \frac{\text{Absolute Dispersion}}{\text{Average}}$$

If the absolute dispersion is the standard deviation, s , and the average is the mean, \bar{X} , the relative dispersion is called the coefficient of variation, given by :

$$\text{Coefficient of Variation} = CV = \frac{s}{\bar{X}}$$

and is usually expressed as a percentage. The lower the CV, the less scatter about the mean is displayed by that particular element.

Scatter may be the result of a number of causes. Firstly, a skewed distribution will result in a high CV, as shown by Cr within the Dominion Group basic lavas, where it in fact exceeds 100 percent. A skewed distribution will also be produced by differentiation, and is more readily displayed by the trace elements than the major elements.

Alternatively, scatter may be the result of the mobility of elements during secondary processes. This is demonstrated by Ba in all formations, for which CV's range from 38.4 percent for the Dominion Group porphyries, to 79.5 percent for the basic lavas in the same group.

The Witwatersrand triad rocks have been subjected to regional, low-grade greenschist facies metamorphism, resulting in alteration of the original mineralogy to a moderate extent. Figures 3.9a, 3.9b, 3.11a and 3.11b show that certain elements appear to have been unaffected by this process, while others show erratic variation profiles.

Bearing in mind the aforementioned circumstances, the coefficients of variation can generally be used to facilitate comparison between elements, especially major elements at different levels of concentration.

Tables 3.2a and 3.2b list the ranges of major and trace elements respectively, for each formation.

Table 3.1a Means, standard deviations and coefficients of variation for major element oxides

ALLA = Allanridge ALBE = Alberton
 RIET = Rietgat-Goedgenoeg JEPP = Jeppetown Amygdaloid
 MAKW = Makwassie DR-P = Dominion porphyries
 LORA = Loraine-Edenville DR-B = Dominion basic rocks
 ORKN = Orkney

MAJOR ELEMENTS

MEANS

	ALLA	RIET	MAKW	LORA	ORKN	ALBE	JEPP	DR-P	DR-B
N	35	67	17	46	31	19	6	32	69
SI02	56.37	56.54	64.56	53.59	54.14	55.40	62.82	69.78	56.05
TI02	1.197	1.524	1.222	0.574	0.981	1.067	1.133	0.812	1.085
AL203	14.68	14.58	14.46	13.39	14.78	14.95	13.76	12.92	14.18
FE203	11.59	10.66	7.62	11.11	12.76	11.44	12.63	6.70	11.82
MNU	0.15	0.14	0.09	0.17	0.15	0.13	0.18	0.08	0.16
MGO	4.18	5.14	2.43	9.01	5.09	5.12	1.02	1.15	5.25
CAO	6.87	6.22	3.34	8.41	7.06	7.19	3.00	2.01	7.16
NA2O	3.25	3.01	3.03	2.48	3.56	3.40	3.70	3.16	3.02
K2O	1.48	1.47	2.73	1.20	1.35	1.13	1.19	3.15	0.97
P205	0.231	0.715	0.521	0.072	0.125	0.156	0.567	0.253	0.308

STANDARD DEVIATIONS

	ALLA	RIET	MAKW	LORA	ORKN	ALBE	JEPP	DR-P	DR-B
SI02	2.56	3.19	3.91	1.87	1.72	1.61	6.23	4.20	3.15
TI02	0.098	0.195	0.181	0.121	0.062	0.047	0.264	0.151	0.399
AL203	0.69	1.19	1.51	1.68	0.47	0.66	2.88	1.76	1.43
FE203	2.21	1.62	1.98	0.84	0.86	0.67	2.08	2.08	1.91
MNU	0.03	0.03	0.03	0.02	0.02	0.02	0.04	0.03	0.03
MGO	1.03	1.37	1.27	2.90	0.55	0.78	0.30	0.93	2.47
CAO	2.27	1.64	1.55	1.98	1.36	1.45	0.54	1.19	1.93
NA2O	0.93	0.61	1.30	0.92	1.00	0.94	0.55	1.11	0.74
K2O	0.96	0.85	1.96	0.61	0.65	0.60	0.69	1.33	0.56
P205	0.029	0.132	0.074	0.022	0.010	0.014	0.148	0.062	0.179

COEFFICIENTS OF VARIATION

	ALLA	RIET	MAKW	LORA	ORKN	ALBE	JEPP	DR-P	DR-B
SI02	4.5	5.6	6.1	3.5	3.2	2.9	9.9	6.0	5.6
TI02	8.2	12.8	14.8	21.1	6.3	4.4	23.3	18.5	36.8
AL203	4.7	8.1	10.4	12.5	3.2	4.4	20.9	13.6	10.1
FE203	19.0	15.2	26.0	7.5	6.7	5.9	16.4	31.1	16.2
MNU	21.1	23.4	36.7	13.9	10.8	13.0	24.7	39.3	18.3
MGO	24.7	26.6	52.3	32.2	10.8	15.2	29.4	81.8	47.1
CAO	33.1	26.3	46.4	23.5	19.3	20.2	17.8	59.1	27.0
NA2O	28.6	20.3	43.1	37.3	28.0	27.7	14.7	35.0	24.5
K2O	64.9	57.6	71.7	51.2	48.0	53.4	58.2	42.2	57.8
P205	12.6	18.4	14.2	31.2	8.0	8.9	26.1	24.6	58.2

Table 3.1b Means, standard deviations and coefficients of variation for trace elements (see Table 3.1a for abbreviations)

TRACE ELEMENTS									
MEANS									
	ALLA	RIET	MAKW	LORA	ORKN	ALBE	JEPP	DR-P	DR-B
N	35	67	17	46	31	19	6	32	69
BA	417.	768.	1152.	349.	431.	438.	622.	768.	453.
NB	8.9	18.5	32.4	2.0	4.8	5.5	16.5	14.2	7.0
ZR	180.	336.	535.	59.	100.	117.	266.	282.	159.
Y	27.2	53.5	71.6	16.6	22.5	22.2	54.8	30.2	28.3
SR	439.	579.	622.	205.	347.	394.	343.	187.	430.
RB	36.7	34.7	89.2	43.4	51.1	39.5	36.9	85.5	28.0
ZN	106.	121.	112.	84.	93.	85.	173.	83.	110.
CU	89.	31.	14.	68.	95.	92.	22.	13.	52.
NI	145.	120.	25.	236.	142.	158.	14.	10.	139.
CO	63.	46.	22.	66.	68.	61.	29.	18.	61.
CR	59.	311.	23.	745.	50.	125.	7.	4.	328.
V	198.	202.	108.	201.	225.	184.	19.	54.	256.
LA	23.	59.	93.	6.	11.	12.	45.	48.	21.
CE	51.	130.	194.	13.	25.	30.	95.	92.	46.
ND	25.	66.	93.	6.	12.	14.	48.	40.	26.
STANDARD DEVIATIONS									
	ALLA	RIET	MAKW	LORA	ORKN	ALBE	JEPP	DR-P	DR-B
BA	221.	345.	807.	262.	256.	278.	378.	295.	362.
NB	1.7	3.9	5.1	1.3	1.0	1.2	3.5	2.9	3.7
ZR	14.	88.	79.	19.	5.	16.	58.	54.	73.
Y	2.5	8.7	8.2	3.2	1.5	1.5	8.9	6.5	9.1
SR	276.	242.	688.	126.	152.	184.	167.	115.	247.
RB	31.4	22.9	64.9	24.0	26.6	23.2	24.7	37.0	22.4
ZN	20.	21.	32.	15.	11.	8.	29.	28.	30.
CU	45.	12.	5.	22.	22.	17.	6.	18.	19.
NI	47.	40.	6.	93.	16.	30.	2.	2.	135.
CO	8.	8.	5.	7.	7.	4.	4.	4.	11.
CR	56.	148.	6.	637.	22.	105.	2.	3.	469.
V	28.	31.	19.	25.	15.	12.	4.	21.	70.
LA	4.	12.	24.	3.	2.	3.	7.	15.	10.
CE	7.	20.	43.	8.	4.	5.	15.	23.	22.
ND	4.	9.	28.	5.	3.	3.	5.	9.	13.
COEFFICIENTS OF VARIATION									
	ALLA	RIET	MAKW	LORA	ORKN	ALBE	JEPP	DR-P	DR-B
BA	53.0	44.9	70.1	75.0	59.5	63.4	60.8	38.4	79.9
NB	18.6	21.0	15.7	64.8	20.7	21.7	21.0	20.1	52.2
ZR	7.6	26.0	14.7	31.8	4.9	13.3	21.8	19.3	45.6
Y	9.1	16.2	11.5	19.5	6.7	6.9	16.3	21.4	32.2
SR	62.8	41.8	110.7	61.6	43.8	46.7	48.6	61.4	57.3
RB	85.7	66.0	72.7	55.3	52.0	58.7	66.9	43.2	80.0
ZN	18.6	17.0	28.7	18.4	11.9	9.7	16.9	33.6	27.8
CU	50.7	38.9	35.8	32.0	23.6	18.5	25.3	137.4	36.3
NI	32.6	33.3	24.1	39.5	11.0	19.2	18.0	22.9	97.7
CO	13.0	17.1	22.4	10.8	10.8	6.9	15.3	22.6	18.3
CR	93.7	47.6	28.2	85.5	44.0	84.1	33.9	77.4	142.9
V	14.4	15.5	18.0	12.5	6.6	6.6	22.6	39.8	27.3
LA	17.8	20.4	26.0	55.2	21.7	24.9	15.7	32.1	50.4
CE	12.8	15.5	22.4	56.0	15.4	17.6	16.3	25.3	49.0
ND	14.0	13.8	30.0	69.3	27.6	19.8	11.0	21.3	49.8

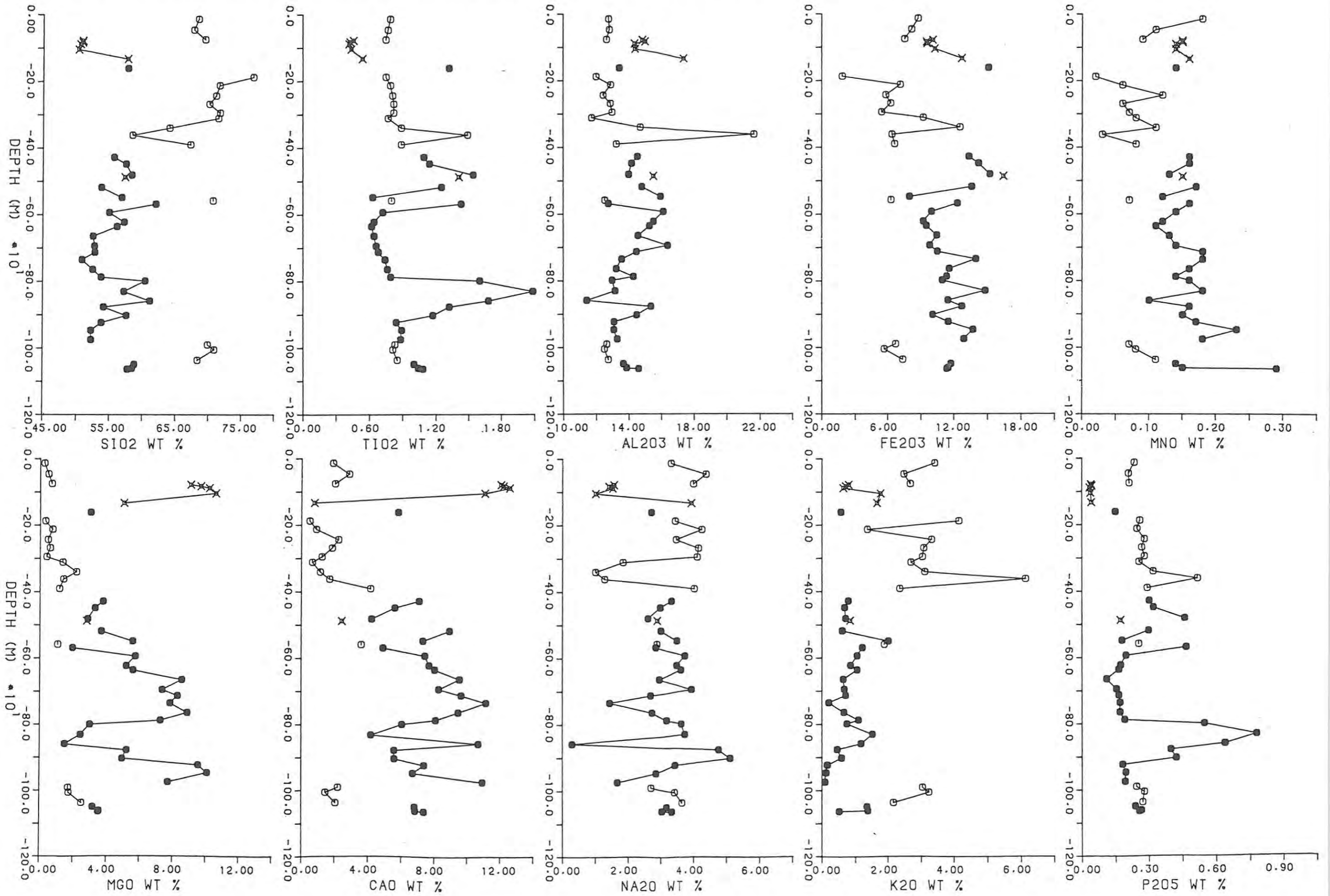
Table 3.2a Ranges of major elements (see Table 3.1a for abbreviations)

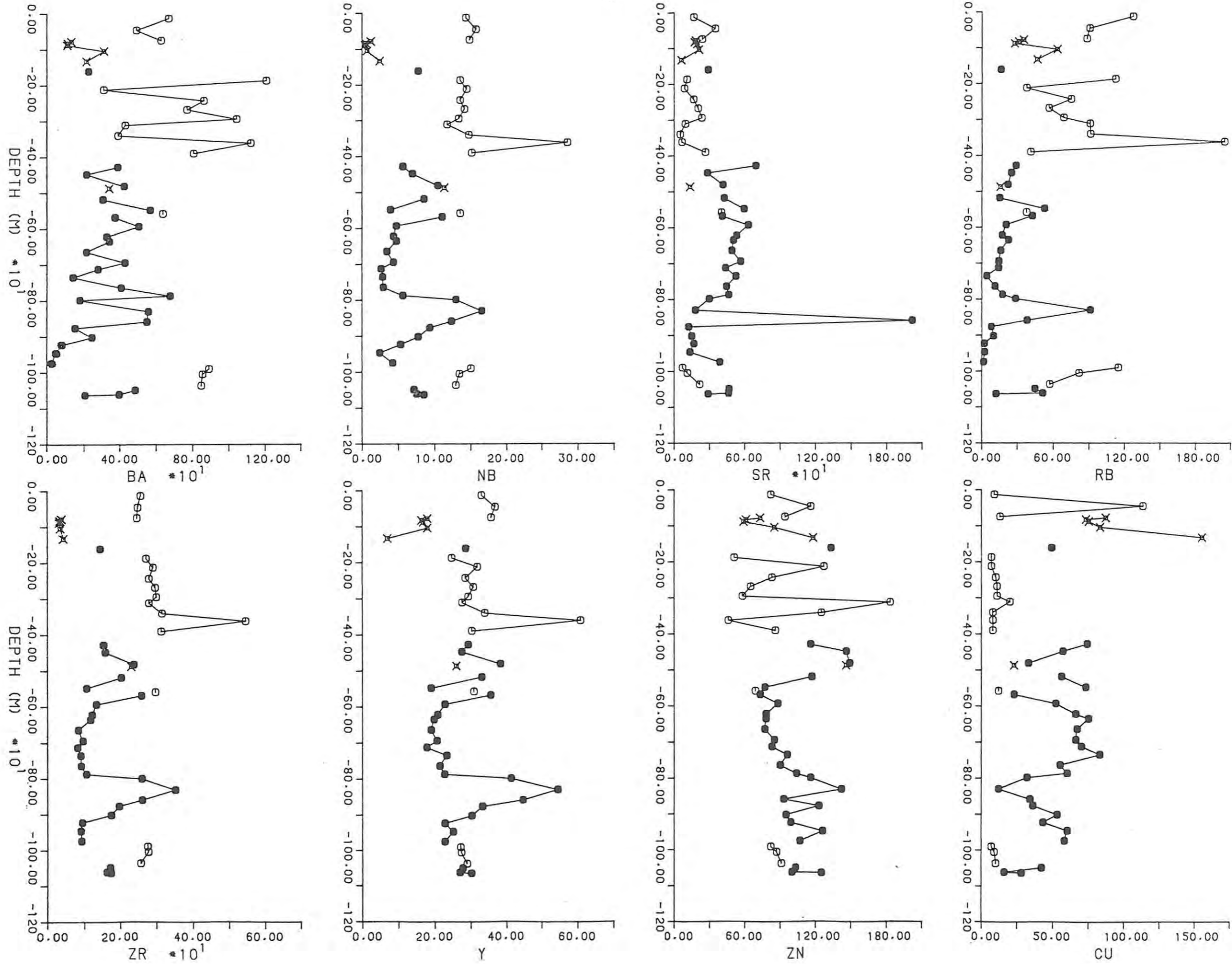
SIU2	min	max	MGU	min	max
ALLA	51.43	61.89	ALLA	2.02	5.78
RIET	47.50	63.42	RIET	2.35	8.20
MAKW	54.93	68.75	MAKW	0.76	5.22
LORA	50.28	57.72	LORA	4.18	17.46
ORKN	51.03	58.88	ORKN	3.57	6.21
ALBE	52.56	58.13	ALBE	4.14	7.29
JEPP	50.22	66.82	JEPP	0.84	1.61
DR-P	55.93	78.93	DR-P	0.26	5.34
DR-B	48.79	62.50	DR-B	1.48	11.36
TI02			CAU		
ALLA	0.891	1.413	ALLA	2.43	14.66
RIET	1.198	2.202	RIET	2.76	12.68
MAKW	1.064	1.729	MAKW	1.84	7.93
LORA	0.400	0.878	LORA	0.79	12.57
ORKN	0.879	1.097	ORKN	5.17	10.66
ALBE	0.973	1.174	ALBE	5.16	9.77
JEPP	0.998	1.670	JEPP	1.96	3.37
DR-P	0.522	1.489	DR-P	0.42	5.32
DR-B	0.603	2.076	DR-B	3.09	11.13
AL203			NA2O		
ALLA	13.39	16.66	ALLA	0.16	4.26
RIET	12.02	19.44	RIET	1.05	4.46
MAKW	12.75	18.55	MAKW	0.51	6.56
LORA	8.59	17.38	LORA	0.15	3.95
ORKN	13.69	15.77	ORKN	2.21	7.08
ALBE	13.76	16.66	ALBE	0.67	4.57
JEPP	11.94	19.57	JEPP	2.72	4.20
DR-P	10.93	21.63	DR-P	0.20	4.40
DR-B	11.41	20.33	DR-B	0.29	5.11
FE203			K2O		
ALLA	5.09	16.43	ALLA	0.30	4.59
RIET	6.29	14.44	RIET	0.11	3.10
MAKW	5.75	13.11	MAKW	0.07	7.47
LORA	9.11	12.72	LORA	0.11	2.49
ORKN	11.07	14.46	ORKN	0.34	2.90
ALBE	10.10	12.51	ALBE	0.38	2.77
JEPP	11.28	16.75	JEPP	0.15	1.90
DR-P	1.81	12.54	DR-P	0.76	7.12
DR-B	7.95	17.15	DR-B	0.08	3.69
MNO			P2O5		
ALLA	0.09	0.25	ALLA	0.164	0.282
RIET	0.07	0.26	RIET	0.520	1.168
MAKW	0.02	0.15	MAKW	0.436	0.687
LORA	0.12	0.24	LORA	0.033	0.131
ORKN	0.11	0.18	ORKN	0.106	0.149
ALBE	0.11	0.17	ALBE	0.137	0.183
JEPP	0.13	0.26	JEPP	0.496	0.869
DR-P	0.02	0.18	DR-P	0.125	0.517
DR-B	0.10	0.29	DR-B	0.109	0.778

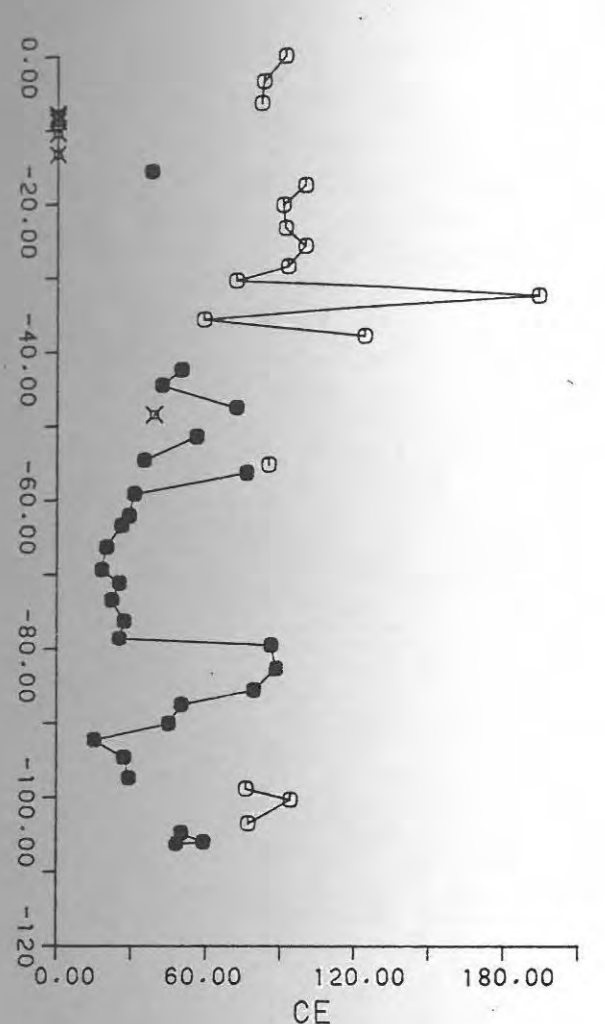
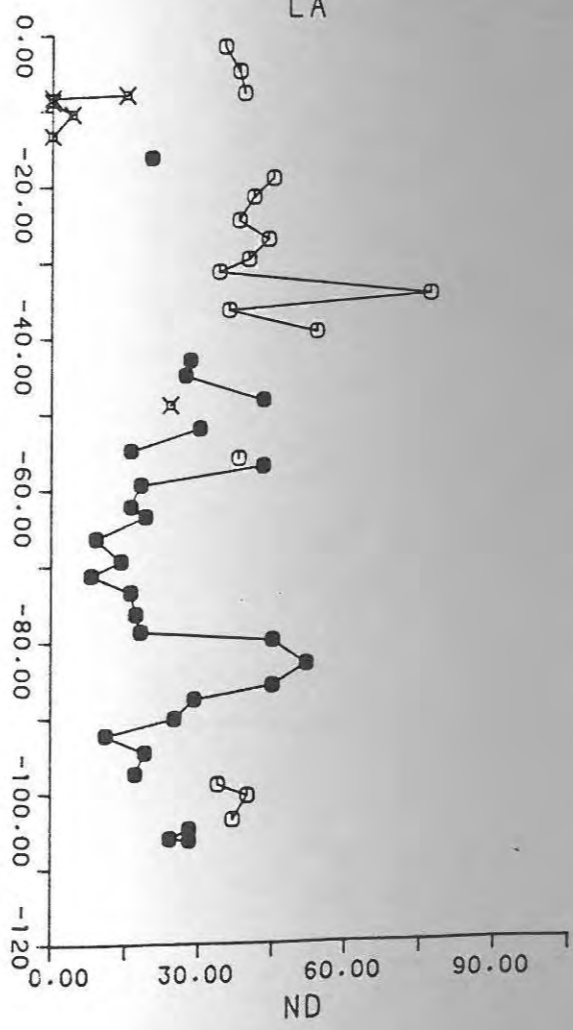
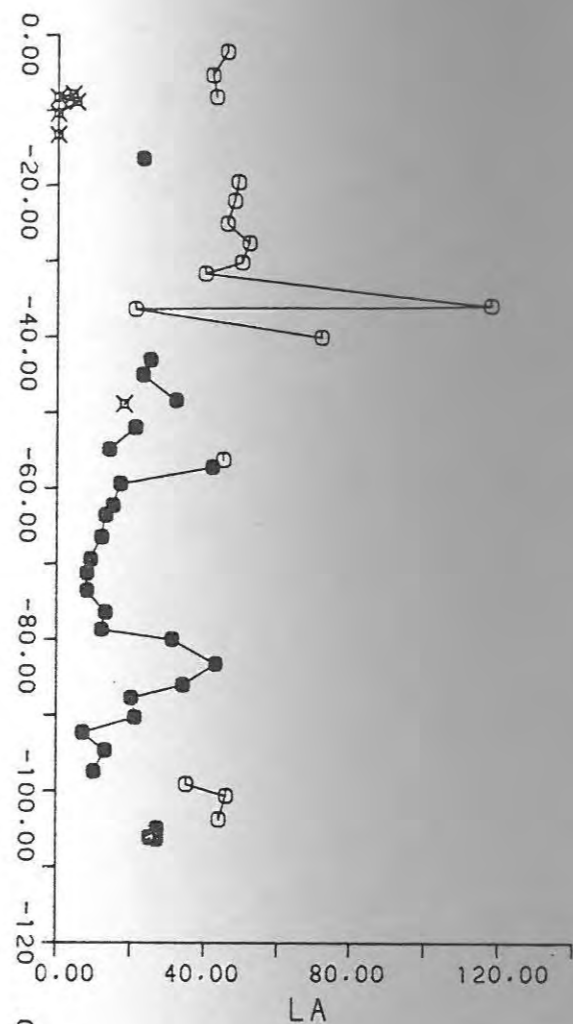
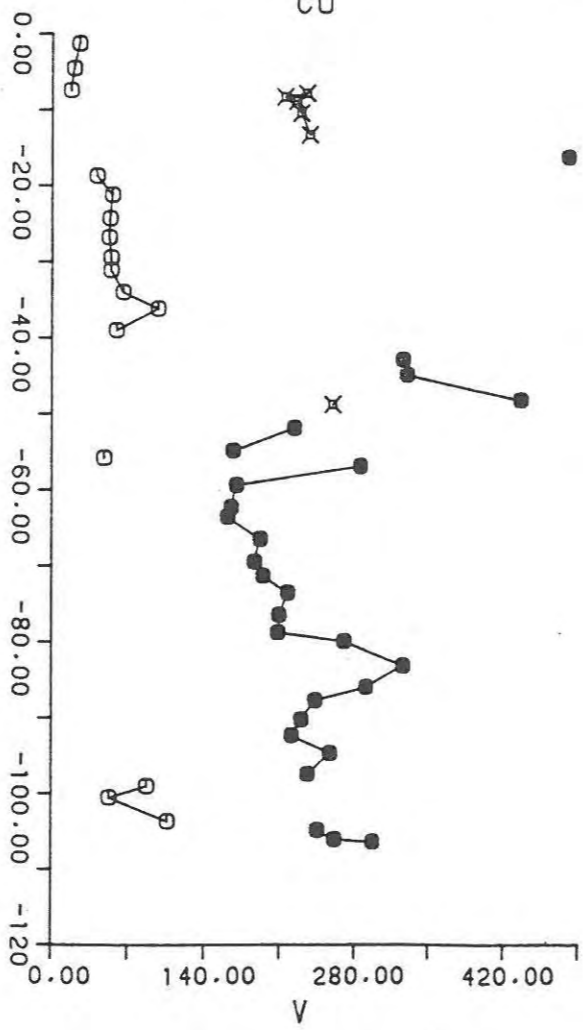
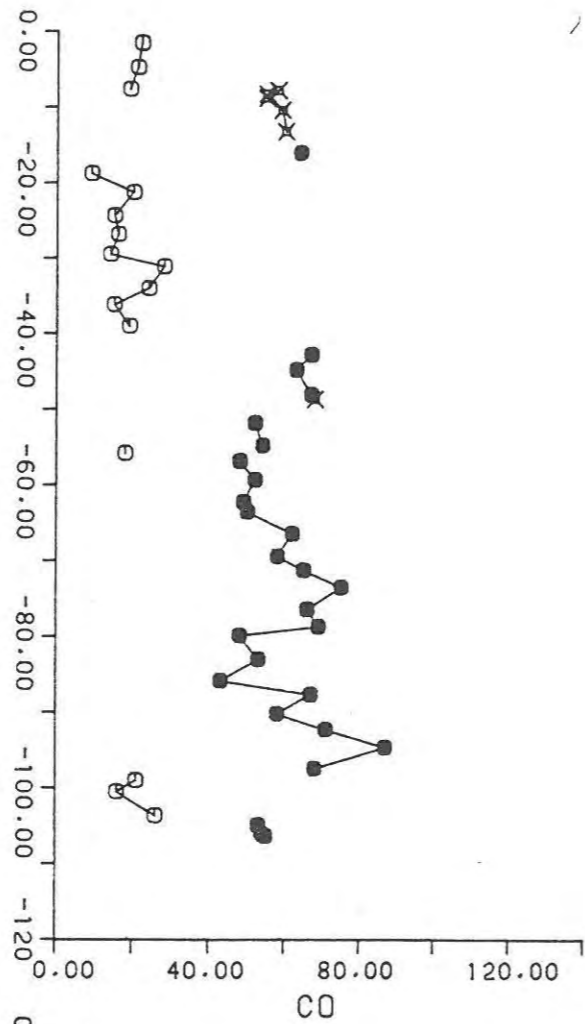
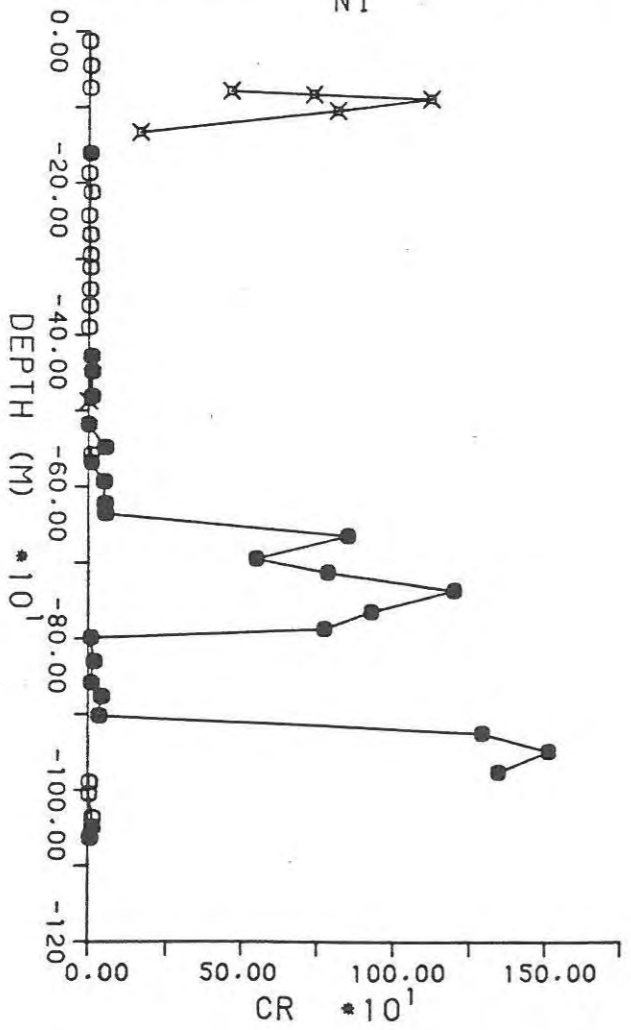
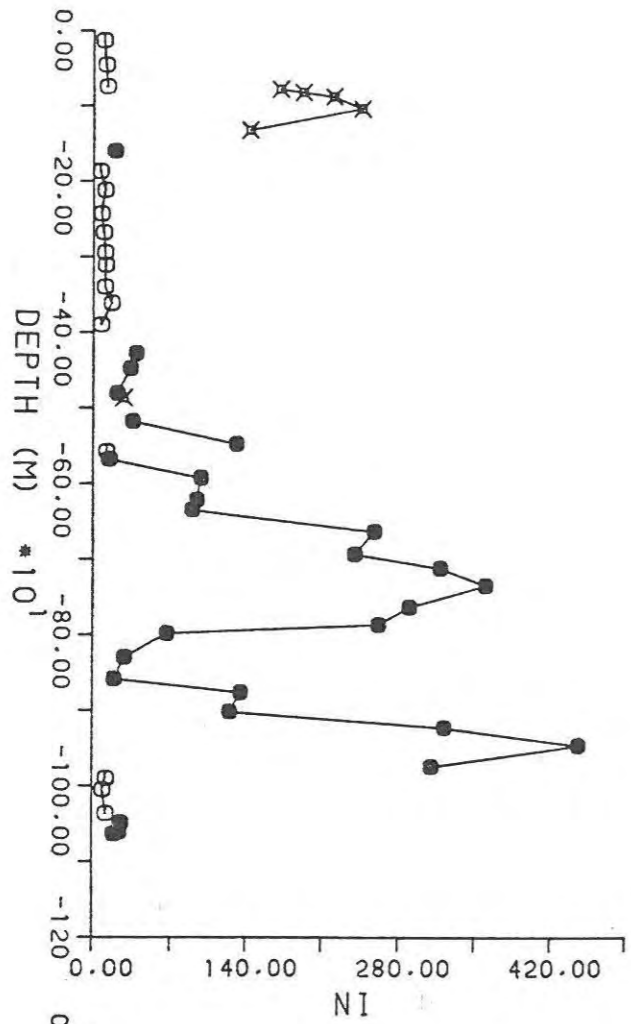
Table 3.2b Ranges of trace elements (see Table 3.1a for abbreviations)

BA	min	max	RB	min	max	CR	min	max
ALLA	97	998	ALLA	5.7	172.4	ALLA	U	213
RIET	85	1599	RIET	5.0	103.0	RIET	17	936
MAKW	25	2784	MAKW	2.0	234.4	MAKW	11	39
LORA	16	1470	LORA	4.1	101.0	LORA	21	2618
ORKN	128	1208	ORKN	10.9	111.9	ORKN	16	97
ALBE	155	1189	ALBE	10.1	87.7	ALBE	64	415
JEPP	122	1210	JEPP	4.3	66.8	JEPP	5	9
DR-P	311	1695	DR-P	55.6	204.4	DR-P	U	12
DR-B	25	2752	DR-B	1.3	106.2	DR-B	U	1510
NB			ZN			V		
ALLA	5.5	12.2	ALLA	52	146	ALLA	147	265
RIET	15.3	33.3	RIET	75	174	RIET	151	314
MAKW	25.6	45.8	MAKW	74	204	MAKW	89	157
LORA	0.0	6.1	LORA	59	118	LORA	153	260
ORKN	2.9	6.8	ORKN	67	126	ORKN	168	258
ALBE	2.0	8.1	ALBE	67	103	ALBE	168	211
JEPP	14.2	23.4	JEPP	155	232	JEPP	15	27
DR-P	11.3	28.5	DR-P	40	183	DR-P	17	106
DR-B	2.0	16.5	DR-B	75	240	DR-B	161	479
ZR			CU			LA		
ALLA	141	230	ALLA	5	182	ALLA	14	31
RIET	244	750	RIET	5	62	RIET	44	113
MAKW	452	748	MAKW	8	28	MAKW	45	140
LORA	28	113	LORA	19	155	LORA	U	15
ORKN	90	110	ORKN	16	137	ORKN	8	18
ALBE	60	137	ALBE	59	122	ALBE	5	19
JEPP	233	383	JEPP	13	30	JEPP	40	59
DR-P	220	545	DR-P	3	113	DR-P	21	118
DR-B	75	351	DR-B	6	86	DR-B	6	43
Y			NI			CE		
ALLA	22.2	32.1	ALLA	27	226	ALLA	57	61
RIET	39.6	95.5	RIET	25	220	RIET	98	199
MAKW	62.1	92.9	MAKW	17	38	MAKW	112	282
LORA	6.9	23.4	LORA	108	516	LORA	U	32
ORKN	19.1	25.5	ORKN	111	163	ORKN	15	31
ALBE	17.7	24.4	ALBE	101	226	ALBE	17	38
JEPP	46.8	72.3	JEPP	11	18	JEPP	81	124
DR-P	25.4	60.7	DR-P	6	18	DR-P	51	194
DR-B	17.3	54.4	DR-B	14	601	DR-B	15	88
SR			CO			ND		
ALLA	55	1472	ALLA	47	82	ALLA	17	31
RIET	216	1395	RIET	26	71	RIET	51	96
MAKW	113	2943	MAKW	16	33	MAKW	61	186
LORA	20	561	LORA	55	88	LORA	U	17
ORKN	174	781	ORKN	47	84	ORKN	U	17
ALBE	147	929	ALBE	51	69	ALBE	8	19
JEPP	194	594	JEPP	24	37	JEPP	41	57
DR-P	43	565	DR-P	9	28	DR-P	22	77
DR-B	121	2015	DR-B	40	110	DR-B	7	52

Figure 3.9a Major element variations with height in the Dominion Group - borehole DSF-7







3.4 Dominion Group

The description and discussion of the chemical data for the Dominion Group porphyries and basic lavas below is in reference to Figures 2.2a, 3.9a and 3.9b. Borehole DSF-7, the representative hole for the Dominion Group, is shown in Figures 3.9a and 3.9b with major and trace element abundances respectively plotted against vertical depth down the borehole. The vertical scale is the same as that in Figure 2.2a in order to facilitate comparison. The zero mark represents the borehole collar. The solid circles represent Dominion Group basic lavas, the open circles, the Dominion Group porphyries, while the "sputniks" indicate intrusive samples of other formations. The discussion which follows encompasses the Dominion Group stratigraphy encountered in all boreholes examined, but space does not permit the illustration of each hole.

SiO₂

The Dominion Group porphyries are the most silicic of the Witwatersrand triad volcanic rocks by virtue of their high SiO₂ content, with a mean of 69.78%. Although their SiO₂ values range between 55.93% and 78.93%, the majority lie above 67%, with only three samples below this value: DP-34 (64.02%), DP-33 (58.44%) and DP-263 (55.93%).

SiO₂ abundances are distinctively higher in the porphyries than in the basic lavas (see Fig. 3.9a). The thickest porphyritic unit (see Figure 2.2a), containing samples DP-32 to DP-40, shows a definite increase in SiO₂ content upwards, while the basic lavas display the same trend, to a lesser degree. The basic lavas cluster about a mean of 56.06% and none of them exceeds 62.5% SiO₂. Some of the most basic lavas of the Witwatersrand triad are found amongst this group, with SiO₂ values as low as 48.79% (in DB-78).

TiO₂ and P₂O₅

TiO₂ and P₂O₅ behave sympathetically and the highly constant values shown by the low-TiO₂ and -P₂O₅ samples are clearly

evident in Figure 3.1a. The porphyries, with the exception of DP-33, show a very tight clustering around .790% TiO_2 . A slight decrease upwards is evident in the thick porphyritic unit in borehole DSF-7, with TiO_2 decreasing systematically from .891% in DP-32 to .747% in DP-40. The basic lavas also show a very smooth linear trend decreasing from 1.084% TiO_2 in DB-1 at the base of DSF-7, to .625% in DB-26, halfway up the hole. This smooth trend is, however punctuated twice by sharp increases, the first peak occurring about 70m above the lowermost porphyry, with a maximum at DB-13 of 2.076%, and the second just below the thick porphyritic unit. This pattern is echoed in borehole DSF-10, which was also sampled at fairly close intervals, and seems to be associated with an increase in the SiO_2 content of the lava.

The pattern of P_2O_5 variation in borehole DSF-7 is virtually identical to that of TiO_2 . The two large spikes shown by TiO_2 variation in the basic lavas are also evident in the P_2O_5 data, reaching a maximum at DB-13 of .778% P_2O_5 , whereas the majority of basic rocks have P_2O_5 varying from .255% at the base to .177% in DB-26. The lowest value is .109% in DB-20. In the porphyries, P_2O_5 values are again very constant, with the exception of DP-33, but show a slight decrease upwards within the thick unit from .292% in DP-32 to .259% in DP-40, decreasing even further near the top of the hole to .208% in DP-48.

Nb, Zr and Y

The elements Nb, Zr and Y, usually regarded as immobile during low grade metamorphism and alteration of igneous rocks, and the rare earths La, Ce and Nd display virtually identical trends in borehole DSF-7. In the basic lavas high concentrations of these elements occur in the SiO_2 -rich samples. Amongst the other samples values decrease from the bottom of the hole until just below the first SiO_2 -rich horizon, after which they increase towards the top of the hole. Average values for Nb, Zr and Y are respectively, 7.0 ppm, 159 ppm and 28.3 ppm. The maximum values for all 6 trace elements are found in sample DB-13, in the middle of the lower SiO_2 -rich horizon in borehole DSF-7.

The trends displayed by the porphyries are identical for Nb and Zr, but less so for Y and the light rare earth elements. With the exception of

sample DP-33, which is highly enriched in these six elements, the following patterns emerge: Nb has a limited range from 11.3 ppm to 16.1 ppm, with an average of 14.2 ppm. No increase or decrease with depth is evident, although a very slight depletion towards the top of the hole may be observed in the thick porphyritic unit of DSF-7, from 15.1 ppm in DP-32 to 13.5 ppm in DP-40.

A more discernible trend in the same unit is shown by Zr, which decreases from 313 ppm in DP-32 to 271 ppm in DP-40. The range of Zr values overall is again fairly narrow, from 220 ppm in DP-95 to 353 ppm in DP-263, with a mean of 282 ppm.

Y and the light rare earth elements show slightly more erratic variations and no particular trend can be discerned in the thick porphyry. With the exclusion of DP-33, the range for Y is from 23.4 ppm to 36.7 ppm.

La, Ce and Nd

In the basic lavas La, Ce and Nd average 21 ppm, 46 ppm and 26 ppm respectively. La ranges from 6 ppm (DB-281) to 43 ppm (DB-13), Ce from 15 ppm (DB-9) to 88 ppm (DB-13) and Nd from 7 ppm (DB-257) to 52 ppm (DB-13). The profiles produced by the variations of these light rare earth elements with depth are almost identical to one another and to those of Nb, Zr and Y.

Within the porphyries the average values for La, Ce and Nd are 48 ppm, 92 ppm and 40 ppm respectively, roughly double those for the basic rocks. However, the variation patterns depart from those shown by Nb, Zr and Y, which decrease upwards within the thick porphyry, as the rare earth element concentrations behave highly erratically in the lower part of the porphyry unit, thus obscuring any possible trends present. Ranges are : La, from 21 ppm (DP-33) to 72 ppm (DP-32), except DP-34 (118 ppm); Ce, from 51 ppm (DP-265) to 124 ppm (DP-32), except DP-34 (154 ppm); Nd, from 22 ppm (DP-265) to 54 ppm (DP-32), except DP-34 (77 ppm).

MgO, Cr and Ni

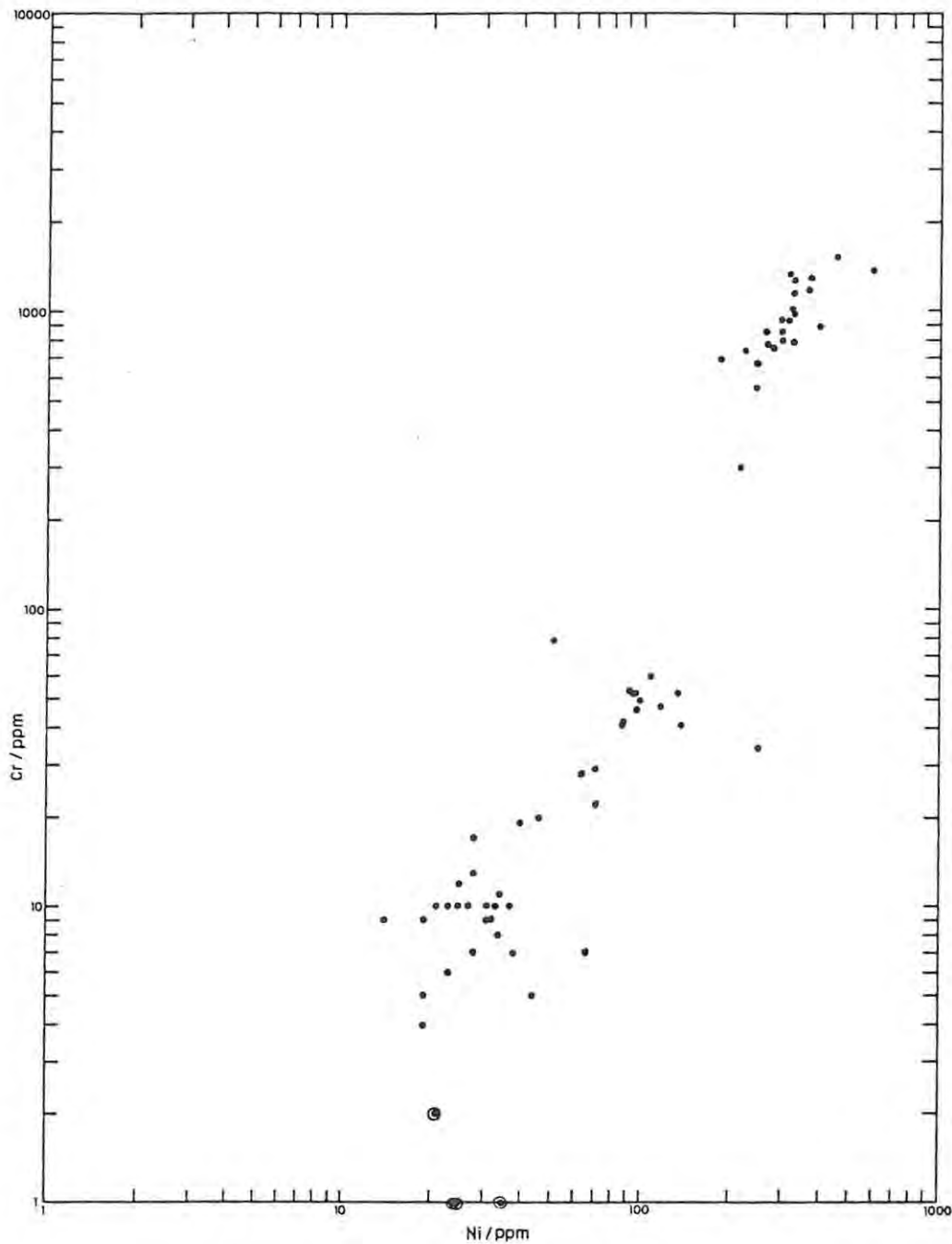
The average MgO content of the porphyries - 1.13% - is, with the exception of that for the Jeppestown amygdaloid, the lowest of all the Witwatersrand triad volcanic rocks. The porphyries display a progressive decrease in MgO content from 2.50% in DP-4 near the base of the Dominion Group, to 0.26% in DB-49 at the top of borehole DSF-7. The latter is the minimum MgO value obtained throughout all the porphyry samples, while the maximum - 5.34% - occurs in DP-263, the sample with the lowest SiO₂ value. The second highest MgO value is the aforementioned 2.50% in DB-4. MgO values in the porphyries at the top and bottom of the hole differ, therefore, by a factor of almost ten.

The basic lavas also show a marked decrease in MgO content from the bottom to the top of DSF-7. The three lowermost samples show relatively low values around 3.5%, but a marked increase is evident just above the basal porphyry, with DB-8 containing 10.07% MgO. This is very close to the maximum of 11.36% shown by DB-52, which occurs in the same stratigraphic position in borehole DSF-10. A prominent negative peak in the upward variation is delineated by samples DB-10 through to DB-14, and again by DB-24. These correspond with the SiO₂-rich samples which also show the very high TiO₂ and P₂O₅ values mentioned earlier.

The trends displayed by Cr and Ni are very similar in borehole DSF-7, and follow closely that of MgO. Both Cr and Ni concentrations are very low in the porphyries, with Cr averaging 4 ppm and Ni averaging 10 ppm, the lowest of all the formations. Cr values lie between 0 ppm (DP-32, 38, 40, 48 and 288) and 12 ppm (DP-4, 89 and 99), while Ni ranges from 6 ppm (DP-40) up to 18 ppm (DP-98), with all but three samples lying below 12 ppm, viz. DP-265 (13 ppm), DP-33 (16 ppm) and DP-89 (18 ppm).

The basic lavas, on the other hand, have fairly high average Cr and Ni contents, 328 ppm and 139 ppm respectively. These averages need to be examined more closely, however, as they are misleading. Within the complete range of values displayed by Cr, 0 to 1510 ppm, there are in fact two distinct populations. Forty six out of the sixty nine basic rock samples, i.e. exactly two thirds, have relatively low concentrations between 0 (DB-27, 273 and 286) and 79 ppm (DB-97).

Figure 3.10 Plot of Cr vs Ni values for the Dominion Group basic lavas showing separation into two distinct groups



Their average is 21 ppm Cr. The remaining twenty three samples have very much higher values, lying between 548 ppm (DB-19) and 1510 ppm (DB-8), with one sample at 299 ppm (DB-251). The average for this group is 942 ppm, almost 45 times that of the low-Cr group. Within the high-Cr group in borehole DSF-7, the highest values are to be found near the bottom, and a general upward decrease is evident.

A similar bimodal grouping emerges when Ni concentrations are examined. Concentrations in the low-Ni group range between 14 ppm (DB-24) and 135 ppm (DB-11) and between 182 ppm (DB-56) and 601 ppm (DB-78) in the high-Ni group. Average values for the two groups are 52 ppm and 311 ppm respectively, with the high-Ni values decreasing towards the top of the borehole. Figure 3.10 is a plot of Cr vs Ni for the basic lavas and the separation of the rocks into two distinct groups is clear.

Fe₂O₃, V and Zn

Within the basic lavas there is a gradual decrease in Fe₂O₃ from roughly between 11.5% and 13.5% at the base, to 7.95% in DB-26, followed by an increase in the overlying silicic horizon to between 13.33% and 16.43% Fe₂O₃. DB-26 reflects the minimum within the basic lavas, while the maximum occurs at 17.15% in sample DB-259 from borehole DRH-15. This sample has low SiO₂ (49.95%). DB-259 corresponds stratigraphically to the upper silicic horizon which shows similar high values. Interestingly, the lower SiO₂-rich horizon is not demarcated by Fe₂O₃ values.

The average Fe₂O₃ value for the porphyries is 6.70%, the lowest of the nine formations of the Witwatersrand triad. Maximum and minimum values are displayed by DP-34 (12.54%) and DP-40 (1.81%) respectively. Within a vertical sequence there appears to be a decrease in Fe₂O₃ content upwards - see DP-32 to DP-40 - although some values are rather erratic. This is reflected in a fairly large CV of 31.3%.

V in the porphyries has an average concentration of 54 ppm, the second lowest after that of the Jeppestown Amygdaloid. The limiting values are 17 ppm (DP-47) and 106 ppm (DP-4). Within the porphyries, a decrease in V content from 106 to 25 ppm is evident towards the top of borehole DSF-7.

The basic lavas have an average of 256 ppm V, the highest of all the formations. Values range between 161 ppm in DB-21 and 479 ppm in DB-41. Again, there is a general decrease towards the top of DSF-7, from 297 ppm at the bottom to 166 ppm in DB-26. Concentrations then show a sudden increase, culminating in 479 ppm in DB-41.

The variation of Zn concentrations in the basic lavas produces a pattern in borehole DSF-7 which is similar to that of V, decreasing towards the top of the hole, but with a few high values near the top. The lower and upper limits are 73 ppm (DB-24) and 168 ppm (DB-52), with the exception of a high 240 ppm in DB-67. The average is 110 ppm.

In the porphyries, Zn values average 83 ppm, the lowest of all the formations, and range between 40 ppm (DP-265) and 183 ppm (DP-35). Unlike V, the Zn concentration in the porphyries has a rather erratic variation and no trend is evident. Both V and Zn follow the general pattern shown by Fe_2O_3 .

MnO

MnO in the basic lavas shows a relatively smooth decrease from .29% in DB-1 at the base of DSF-7, to .16% in DB-31, just below the thick porphyritic unit. Extreme values are the aforementioned .29% and .10% in DB-12. In the porphyries MnO values are variable and range from .02% in DP-40 up to .18% in DP-49, averaging .08%.

CaO and Sr

The basic lavas have an average of 7.09% CaO. Values are highly variable, but there is a slight decrease upwards. The maximum value of 11.13% occurs in DB-17 from the lower half of borehole DSF-7, while the minimum of 3.09% is found in DB-78 from borehole DRS-6.

CaO concentrations in the porphyries are very low, in fact, the lowest of all the Witwatersrand triad volcanics, averaging 2.01%. The range is fairly small, from 0.42% in DP-265 to 5.32% in DP-263, with no particular trends emerging.

The porphyries have an average Sr content of 187 ppm, the lowest of all the volcanics. Values lie between 43 ppm (DP-265) and 565 ppm (DP-81) and are clearly widely scattered about the mean. There is no discernible trend in the variation of concentration with depth.

In the basic lavas Sr ranges between 121 ppm (DB-11) and 750 ppm (DB-61), with the exception of DB-12 which has 2015 ppm Sr. The

average Sr value is 430 ppm, and the lower values appear to be near the bottom of the hole.

Al₂O₃

Al₂O₃ content in the porphyries is extremely constant, the majority of samples ranging between 12.00 and 13.00%, and only six samples fall outside these limits. These are DP-265 (10.93%), DP-35 (11.73%), DP-32 (13.25%), DP-34 (14.70%), DP-263 (15.11%) and DP-33 (21.63%). There appears to be no systematic increase or decrease with height. The average Al₂O₃ value of 12.92% is the lowest of all the formations of the Witwatersrand triad volcanics.

The basic lavas have an average of 14.23% Al₂O₃ and range between 11.41% in DB-12 and 16.37% in DB-19, with the exception of two very high values occurring in DB-259 (17.25%) and DB-78 (20.33%). Concentrations are slightly lower in the bottom half of borehole DSF-7, but decrease again near the top.

Na₂O

Na₂O values in both the porphyries and the basic lavas are highly variable, ranging in the basic lavas from 0.29% (DB-12) up to 5.11% (DB-10), with an average of 3.04%. In the porphyries the range is slightly smaller, from 0.20% (DP-265) to 4.40% (DP-51). The average is 3.16% and again there is no discernible trend with depth.

K₂O, Rb and Ba

K₂O concentrations in the porphyries average 3.15%, the highest of all the formations under scrutiny. Values lie between 0.76% (DP-263) and 5.70% (DP-80), except for DP-33 which has 7.12% K₂O. Within the thick porphyritic unit in borehole DSF-7, concentrations increase upwards from 2.35% (DP-32) to 4.14% (DP-40).

Within the basic lavas, a general increase in K_2O towards the top is also evident, producing a trend very similar to that of SiO_2 . With an average of 1.00%, the lowest of the nine formations, values range between 0.08% (DB-7) and 2.11% (DB-72), with the exception of DB-85, which has 3.65% K_2O .

The average Rb concentration in the porphyries is 85 ppm, the range being from 34 (DP-81) to 150 ppm (DP-265), with the exception of DP-33, which has 204 ppm Rb. The variation with height is highly erratic with no discernible trend with depth.

In contrast, the basic lavas, show a systematic increase in Rb content with height in the thickest unit, the trend being punctuated by two large peaks delineated by samples which also have high Nb, Zr, Y and light rare earth elements. The average Rb content is 28.0 ppm, the lowest in all the volcanics. Values range from 1.3 ppm (DB-7) up to 106 ppm (DB-72), with all except four values lying below 70 ppm.

In the basic lavas Ba concentrations show a large degree of scatter. Despite the scatter, there is a general upward increase in Ba, similar to that of K_2O . The average is 453 ppm and values range from 25 ppm in DB-7 to 1090 ppm in DB-79, with an excessive amount of 2752 ppm in DB-85. This very high value corresponds with a high K_2O content of 3.65% in a group of lavas whose average K_2O content is only 1.00%.

Amongst the porphyries the spread of Ba values is very large and erratic, but higher on average than in the basic lavas. Values range from 311 ppm in DP-39 to 1695 ppm in DP-59, with an average of 768 ppm. Ba seems to be amongst the most mobile of the trace elements determined in the Dominion Group and as such should not be utilised for any petrogenetic interpretation.

Cu

The Cu profile down borehole DSF-7 is much the same as that of Ni, but without the strong polarisation into high- and low-concentration groups within the basic lavas. The average for the basic lavas is 52 ppm, extreme values being 6 ppm in DB-6, and 86 ppm in DB-61.

The porphyries have a rather restricted Cu content, averaging 13 ppm, the lowest of all the formations. Values range from 3 ppm (DP-265) up to 20 ppm (DP-35), with the exception of DP-48 (113 ppm). There is no particular trend with depth in either the porphyries or the basic lavas.

Co

The range of Co concentrations in the porphyries extends from 9 ppm (DP-40) up to 28 ppm (DP-35), with an average of 18 ppm, the lowest for all the volcanics. There is very little spread about the mean, and no discernible systematic variation with depth.

Concentrations of Co in the basic lavas display a distinct decrease with height in the thickest unit. The average value for Co is 61 ppm, with limiting values at 40 ppm (DB-50) and 87 ppm (DB-8), except for a high 110 ppm in DB-78. The tight clustering about the mean is reflected in a low CV of 18.3.

3.5 Crown Formation

Only six samples were taken from the Crown Formation, three from borehole R-1, immediately south of Klerksdorp, and three from borehole JY-8, south of Orkney, about 14.5 km away. The small number of samples did not warrant the plotting of variation-with-depth diagrams, so only a very cursory description of the major and trace element concentrations follows.

The first five samples, JC-366 to JC-370, are very similar in major element composition, while the last sample, JC-371, differs sharply in all respects except CaO, Na₂O and K₂O values (see Table 2.4). It must be borne in mind, therefore, when examining mean values, that they will have been influenced by some extreme values from JC-371.

Of all the formations studied, the Crown Formation has the highest average MnO and Na₂O concentrations, 0.18% and 3.70% respectively, and the lowest average MgO (1.02%). For such low-MgO and relatively high-SiO₂ rocks, the Fe₂O₃ average of 12.63% is very high.

Trace element determinations show that sample JC-371 again has very different concentrations from those of the other five samples, values being higher for all trace elements except Sr, Rb and Cr. The average value of V - 19 ppm - is the lowest of all the formations, while the Zn average of 173 ppm is the highest.

3.6 Ventersdorp Supergroup

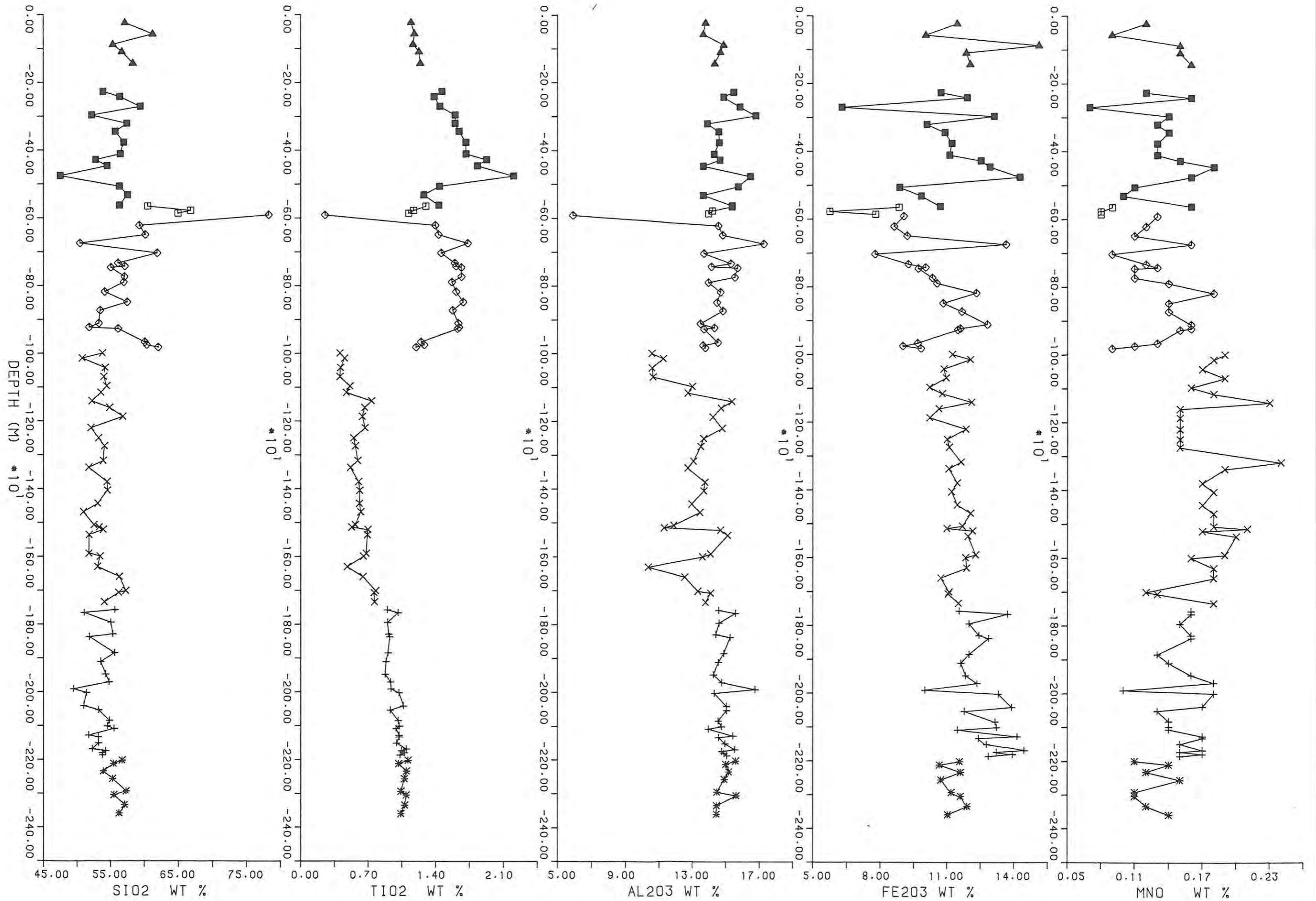
Figures 3.11a and 3.11b show major and trace element concentrations respectively plotted against vertical depth down borehole WS-5, the representative borehole for the Ventersdorp Supergroup. The lithostratigraphic log of borehole WS-5, which also shows the positions of samples (see Figure 2.2b), has been drawn at the same vertical scale as the down-the-hole plots in order to facilitate comparison between the two. Symbols used are as follows:

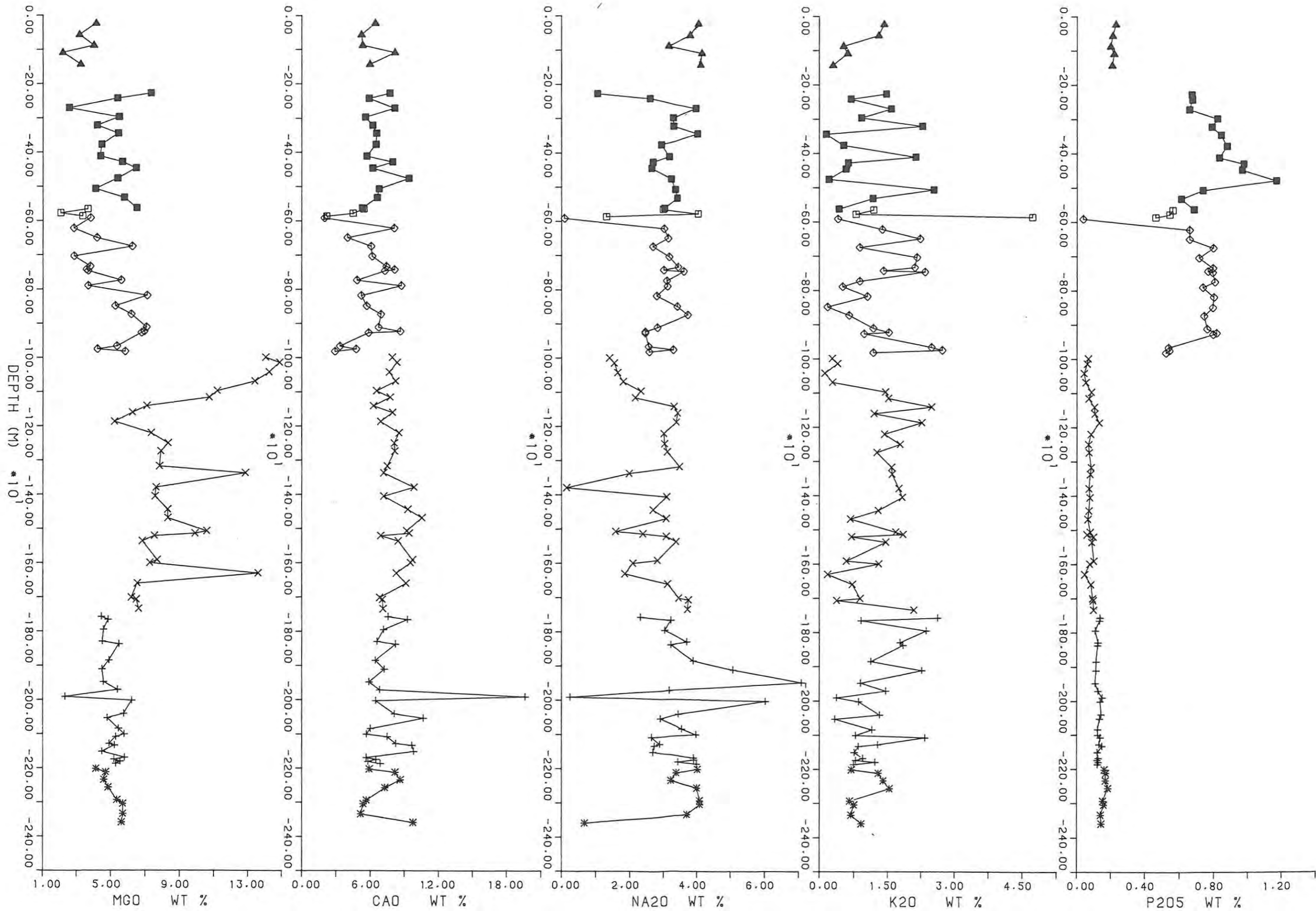
- | | |
|----------------|-----------------------|
| ▲ - Allanridge | x - Loraine-Edenville |
| ■ - Rietgat | + - Orkney |
| □ - Makwassie | * - Loraine |
| ◇ - Goedgenoeg | |

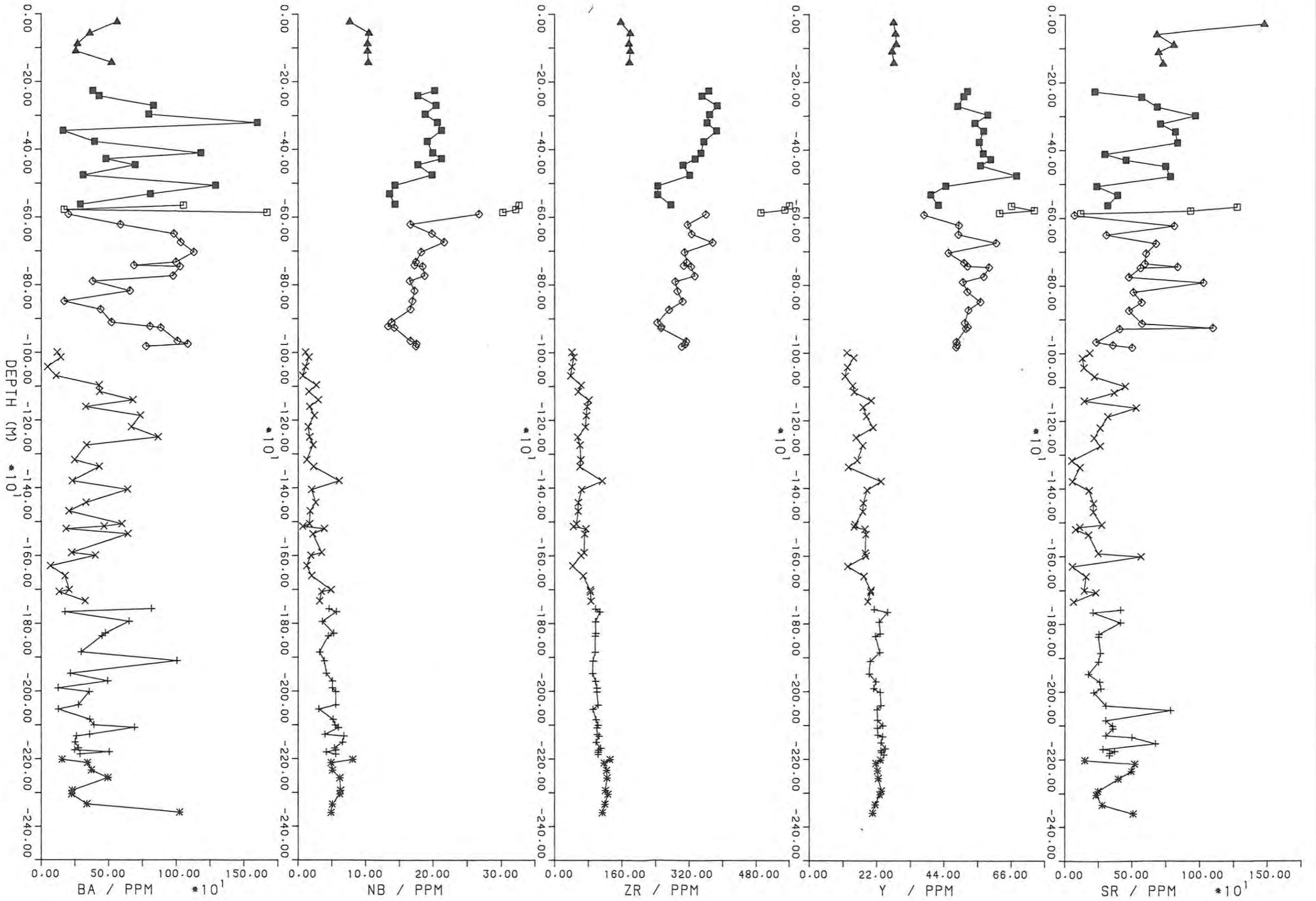
An examination of chemical analyses has revealed a number of interesting points: the three lowermost samples in the Goedgenoeg formation, PG-160, 161 and 162, overlie the Klipriviersberg Group directly and are separated from the rest of the Goedgenoeg formation by a 36-metre conglomerate intercalation of the Kameeldoorns Formation. In borehole WS-4, which is nearest to WS-5, a similar situation occurs, where sample PG-206 is overlain by 55 metres of conglomerate, which in turn is overlain by the rest of the Goedgenoeg formation. In borehole SH-1, sample PR-395, which was collected from a volcanic horizon within the Kameeldoorns conglomerates, has chemical characteristics identical to those of other samples of the Goedgenoeg-Rietgat volcanic rocks.

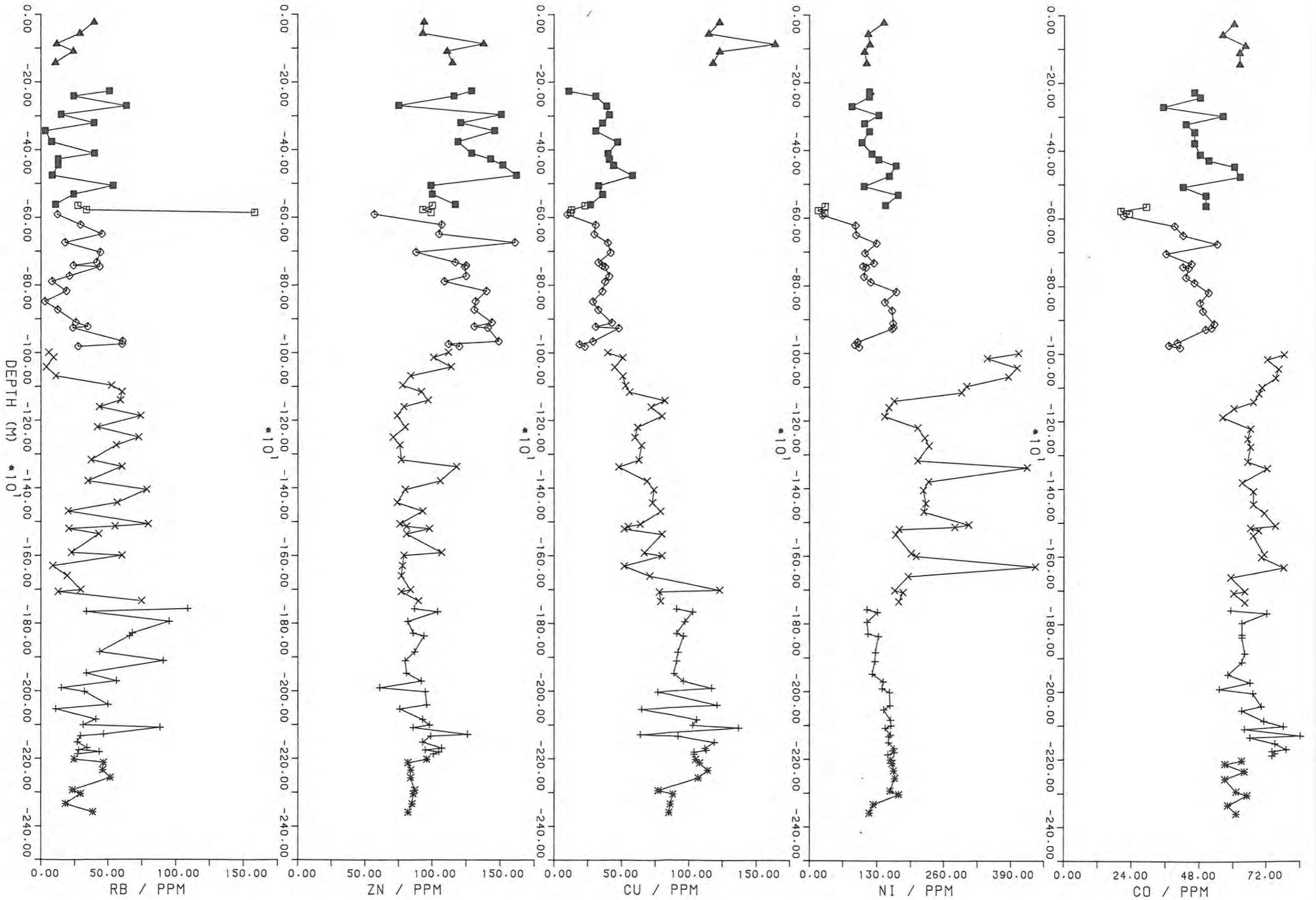
As this study progressed, several changes were proposed, on the basis of chemistry, to the positioning of formational boundaries which had been visually demarcated by company geologists, in boreholes W-20, PK-10 and WS-5. In W-20, the interval from 192m to 267m had been logged as part of the Goedgenoeg formation, followed by 4.5m of

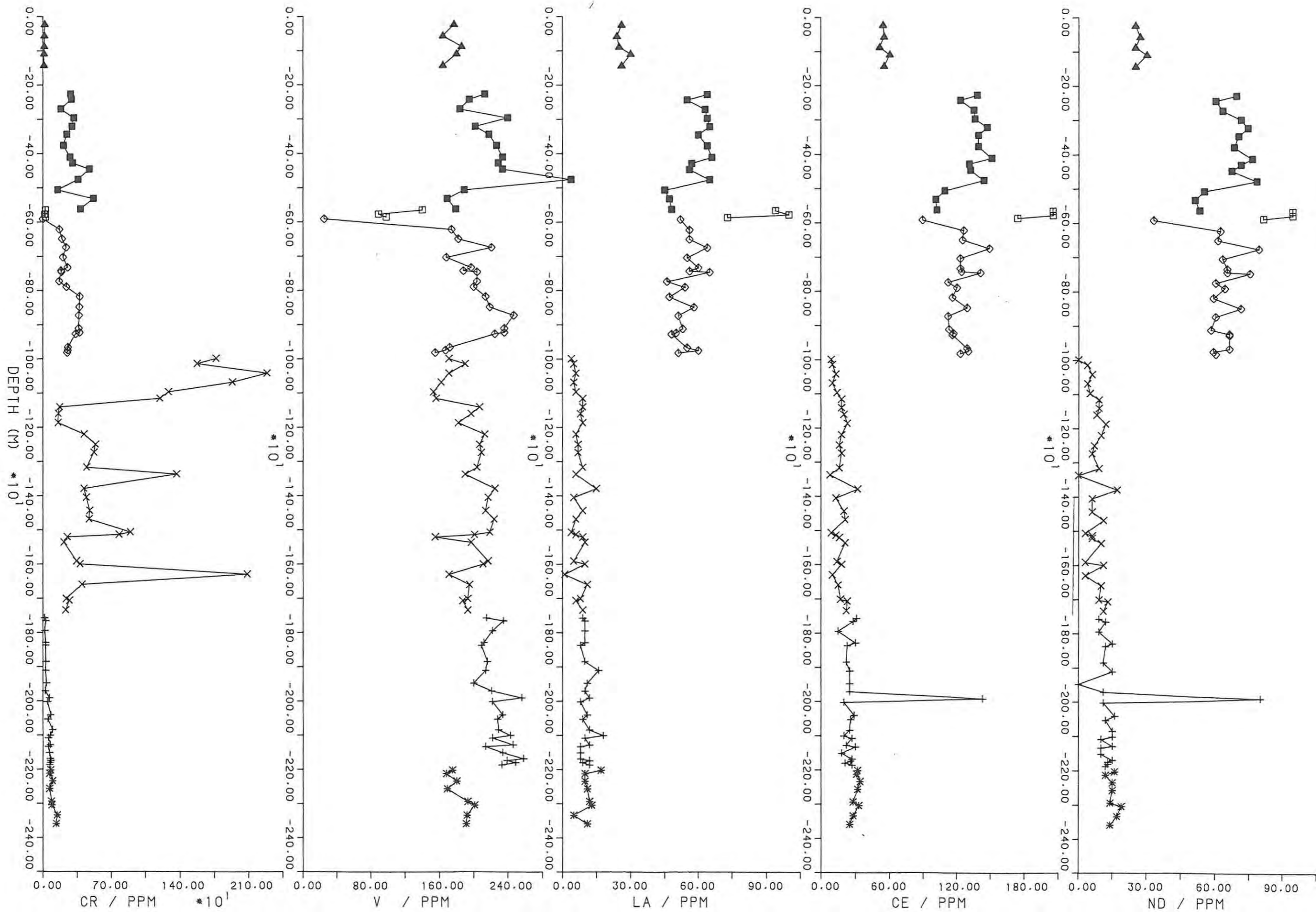
Figure 3.11a Major element variations with height in the Ventersdorp Supergroup - borehole WS-5 - 73 -











Kameeldoorns conglomerate, with the underlying Loraine Formation occurring to 960m. Chemical analyses suggested, however, that the upper limit of the Orkney Formation was not at 960m, but higher. Unfortunately the sampling interval here is large, but the suggested boundary must be between 626m (depth of altered zone) and 550m (base of flow containing sample AR-350) from surface. Above this, all samples are classified as Allanridge lavas.

In borehole PK-10 the geochemical boundary between the Loraine and the overlying Allanridge Formation has been moved from 573m down to somewhere below 769m, the depth of sample AR-300. Finally, in borehole WS-5, the Loraine-Orkney boundary was shifted from 1873m up to somewhere between 1734m and 1738m (the boundaries of the flows from which samples KL-131 and KO-130 were taken). The Orkney-Alberton boundary was also moved slightly, from 2207m up to somewhere between 2202m and 2186m (between samples KO-108 and KA-107). It must be stressed that these boundary changes were made purely on a geochemical basis. It appears superficially as if the original visual logging was incorrect, but, it may well transpire that visual and chemical criteria for boundaries do not coincide. This cannot be resolved until the borehole core is re-examined visually in the light of geochemical results.

SiO₂

SiO₂ content in the Klipriviersberg Group (KRBG) decreases upwards, the Alberton Formation (AF) averaging 55.40% SiO₂; the Orkney (OF), 54.14% and the Loraine-Edenville (LEF), 53.59%, the lowest of all the Witwatersrand triad lavas. There is very little scatter, shown by low CV's of 2.9, 3.2 and 3.5 respectively.

The Goedgenoeg and Rietgat Formations (GRF) are more silicic, averaging 56.54% SiO₂, while the intervening Makwassie Formation (MF) has an average of 64.56% SiO₂. Sample PG-177, a vitric tuff, has anomalous chemistry and was not included in the calculation of averages. There is more scatter in the Platberg Group (PG) than in the KRBG, shown by larger CV's, but no systematic trends are evident.

The Allanridge Formation (ARF) has a mean SiO_2 content of 56.37%, similar to that of the GRF, and exhibits no systematic trend.

MgO, Cr, Ni and Co

MgO in the AF decreases consistently towards the top, averaging 5.12%. A similar trend is evident in the OF, for which the average is a slightly lower 5.09%. A sudden increase occurs, however, at the base of the LEF, and is continued towards the top, where MgO values up to 17.46% (KL-468) occur. In addition, three sharp spikes delineating narrow MgO-rich zones are present.

Cr and Ni values follow almost exactly the same trends as MgO. Concentrations of Cr in the AF and OF are fairly low, averaging 125 and 50 ppm respectively, while in the overlying LEF, the mean increases to 745 ppm, with some values as high as 2618 ppm. The AF Ni average of 158 ppm and the OF average of 142 ppm are again lower than that of the LEF, which is 236 ppm, with some values as high as 516 ppm.

Overlying the high-MgO KRBG lavas, the PG has relatively low MgO, averaging 5.14% in the GRF and 2.43% in the MF. There appears to be a decrease in MgO upwards in Goedgenoeg samples, and possibly the same in the Rietgat unit, but the trend in the latter is rather obscure. Cr and Ni concentrations, however, do decrease upwards in both the Goedgenoeg and Rietgat units. No systematic variations of Cr and Ni with height are observed in the MF, the averages of which are both 6 ppm. Concentrations of MgO, Cr and Ni, which average 4.18%, 59 ppm and 145 ppm respectively, increase upwards in the ARF.

The variation in Co is broadly similar to that of MgO, although Co content does not decrease upwards in the AF, nor does it increase upwards in the ARF. In addition, the OF has a higher Co average (68 ppm) than the LEF (66 ppm), although Co concentrations do increase in the same horizons as do MgO, Ni and Cr, but not to such a large degree.

TiO₂ and P₂O₅

Sympathetic variations are shown by TiO₂ and P₂O₅. Although concentrations increase slightly upwards in the AF, within the KRBG as

a whole there is a general decrease upwards, the mean TiO_2 values for the AF, OF and LEF being 1.06%, 0.981% and 0.574% respectively, the latter the lowest average TiO_2 of all the Witwatersrand triad lavas. P_2O_5 averages 0.156%, 0.125% and 0.075% respectively, the latter again being the lowest overall mean.

Compared with the KRBG, TiO_2 and P_2O_5 contents in the PG are much higher, and are in fact the highest of all the Witwatersrand triad lavas. Within the GRF, TiO_2 ($\bar{X} = 1.524\%$) and P_2O_5 ($\bar{X} = 0.715\%$) decrease upwards, while within the MF these elements display no systematic trend with height and average 1.222% and 0.521% respectively. The ARF averages 1.197% TiO_2 and 0.231% P_2O_5 , with TiO_2 decreasing upwards and P_2O_5 showing no systematic variation.

Nb, Zr and Y

These immobile trace elements exhibit vertical variations similar to those of TiO_2 and P_2O_5 . Concentrations of these elements in the AF increase slightly upwards, but an overall decrease upwards in the KRBG as a whole is shown by the means for the AF, OF and LEF, which are, for Nb, 5.5, 4.8 and 2.0 ppm respectively; for Zr, 117, 100 and 59 ppm respectively, and for Y, 22.2, 22.5 and 16.6 ppm respectively. The means for the LEF are the lowest of all the Witwatersrand triad lavas for these three elements.

The various members of the PG have the highest Nb, Zr and Y contents of all the Witwatersrand triad lavas. The means for Nb, Zr and Y for the MF are 30, 535 and 72 ppm respectively, while those for the GRF are 18.5, 336 and 54 ppm respectively. In both the Goedgenoeg formation (GF) and the Rietgat Formation (RF) an upward increase in these elements is evident, while in the MF there is no systematic trend.

The averages of Nb, Zr and Y for the ARF are respectively 8.9, 180 and 27.2ppm, and no systematic variation in the concentration of these elements with height is discernible.

La, Ce and Nd

These light rare earth elements (LREE) form a further group which varies sympathetically with TiO_2 and P_2O_5 . In the KRBG, a

general decrease upward is evident, with the averages of La, Ce, Nd in the AF, OF and LEF being respectively 12, 11 and 6 ppm for La, 30, 25 and 13 ppm for Ce, and 14, 12 and 6 ppm for Nd. The sharp spike showing anomalously high Ce and Nd represents sample KO-121, which has highly aberrant chemistry and was therefore not included in the various statistical calculations.

The base of the PG is demarcated by a sudden increase in the LREE concentrations. The averages for the GRF are 23, 51 and 25 ppm for La, Ce and Nd respectively, all of which show slight increases upwards. The MF has the highest average LREE concentrations, namely 93, 194 and 93 ppm for La, Ce and Nd respectively, but these elements show no systematic trend with depth.

The concentrations of LREE in the ARF lie between those of the KRBG and the PG, with La, Ce and Nd averaging 23, 51 and 25 ppm respectively. No particular systematic trends with depth are evident.

Fe₂O₃, V and Zn

Within the AF Fe₂O₃ decreases slightly upwards, averaging 11.44%. At the base of the overlying OF Fe₂O₃ increases sharply, thereafter decreasing towards the top, with an average of 12.76%. Values decrease still further towards the top of the LEF, which averages 11.11% Fe₂O₃.

The GRF has a more variable Fe₂O₃ content, but it appears to decrease upwards in each unit, the average being 10.66%. The MF has a still lower average Fe₂O₃ content - 7.62% - with no particular variation with height being evident.

In the overlying ARF the mean Fe₂O₃ content is 11.59%, with no systematic increase or decrease evident.

V behaves sympathetically with Fe₂O₃, averaging 184 ppm in the AF, 225 ppm in the OF, and 201 ppm in the LEF. Within each individual formation Fe₂O₃ contents decrease upwards. Within the GF and RF Fe₂O₃ contents also decrease upwards, averaging 202 ppm, about the same as that of the underlying KRBG. The MF, however, has much lower V contents, the average being 108 ppm, while that of the ARF is 198 ppm.

Zn averages 85 ppm in the AF, 93 ppm in the OF, and 84 ppm in the LEF. Zn decreases upwards in the OF, and then increases towards the top of the LEF. Within the GF Zn concentrations decrease upwards, the mean being 121 ppm. In the MF Zn averages 112 ppm, and in the ARF 106 ppm, with an upwards decrease in Zn concentration being evident in the ARF.

K₂O, Rb and Ba

K₂O concentrations are highly variable, but there appears to be an upward-increasing trend in the OF and the LEF, followed by a decrease in the uppermost samples of the latter. Averages for the AF, OF and LEF are, respectively, 1.13%, 1.35% and 1.20%. No vertical trends are evident in the GRF, or in the MF, the averages for which are 1.47% and 2.73% respectively. K₂O averages 1.48% in the ARF and increases upwards in the sequence.

The variation shown by Rb is almost identical to that of K₂O. Averages for the AF, OF and LEF are 40, 51 and 43 ppm respectively, and the same increase upwards in Rb concentration in the OF and LEF is evident. However the uppermost samples in the LEF reflect a sharp decline in Rb. The GRF averages 35 ppm, and the MF, 89 ppm, the highest of all the Witwatersrand triad lavas. Rb increases upwards in the ARF and averages 37 ppm Rb.

Ba concentrations are also highly variable, but similar trends to those shown by Rb and K₂O are apparent. The average Ba values for the AF, OF and LEF are 438, 431 and 349 ppm respectively, the latter being the lowest of all the Witwatersrand triad lavas. The average Ba content for the GRF is 768 ppm, while that of the MF is 1152 ppm, the highest of all the formations. The ARF has an average of 417 ppm Ba, and the expected increasing-upwards trend has begun to be obscured by the variation.

CaO and Sr

CaO increases slightly from the base to the top of the KRBG, with the average for the AF being 7.19%, that of the OF marginally lower at 7.06%, and that of the LEF 8.41%. The large spike amongst the Orkney

samples is produced by sample K0-121, which has anomalously high CaO (19.61%) accompanied by a dilution of most of the other major element constituents. Interestingly, TiO_2 and P_2O_5 concentrations, as well as those of the immobile trace elements, Nb, Zr and Y, have been unaffected. Also of note, is the fact that the Sr concentration - 266 ppm - is not sympathetically high.

The average Sr content in the AF is 394 ppm, 347 ppm in the OF, and 205 ppm in the LEF. There is thus an overall decrease in Sr concentration towards the top of the KRBG. In the GRF no systematic trend is evident, and CaO averages 6.22%. The average Sr value is 579 ppm, but concentrations are highly variable. There are no systematic CaO or Sr trends in the MF, and averages are 3.34% and 622 ppm respectively, the latter being the highest Sr average of the Witwatersrand triad lavas. This is unusual for a formation which has the lowest CaO average after the Dominion Group porphyries.

Neither CaO nor Sr show any systematic trend in the ARF, and averages are 6.87% and 439 ppm respectively.

Al_2O_3 and Cu

Al_2O_3 in the KRBG averages 14.95%, 14.78% and 13.39% in the AF, OF and LEF respectively, producing an upward-decreasing profile. The GRF average is marginally higher at 14.58%, and there appears to be a slight increase upwards in Al_2O_3 content. No systematic trend is evident in the MF, the average of which is 14.46%. The ARF has the highest Al_2O_3 average of all, 14.68%.

Cu behaves sympathetically with Al_2O_3 in the KRBG, averaging 92, 95 and 68 ppm in the AF, OF and LEF respectively, indicating a general decrease in Cu content towards the top of the group. In the overlying PG, however, the Cu- Al_2O_3 relationship appears to be antipathetic in the GF, and Cu averages are lower in the GRF and MF, being 31 and 14 ppm respectively. Cu values are higher again in the ARF, averaging 89 ppm.

Na₂O

Na₂O variation is very erratic in the KRBG, especially in the OF, but values appear to decrease upwards overall. The averages for the AF, OF and LEF are 2.40%, 2.56% and 2.48% respectively, the latter being the lowest of all the Witwatersrand triad lavas. Three prominent negative peaks in the LEF samples correspond with the high-MgO peaks described earlier. Concentrations of Na₂O in the GRF show a slight increase with height, the average being 3.01%, while that for the MF is similar, at 3.03%. The average Na₂O content of the ARF is a little higher, 3.25%, and the profile shows no systematic trend.

MnO

All the samples were crushed in Mn-steel containers, and although Mn contamination of samples can be expected this has been shown (J.S. Marsh, pers.comm.) to be slight. Nevertheless Mn variations described here should be treated with circumspection. There appears to be a general increase in MnO towards the top of the KRBG, with averages being 0.13%, 0.15% and 0.17% for the AF, OF and LEF respectively. MnO concentrations in the PG show no systematic trends, but averages for the GRF and MF are 0.14% and 0.09% respectively, while that of the ARF, in which there seems to be a decrease with height, is 0.15%.

To conclude, despite the effects of low grade metamorphism to which these rocks have been subjected, the variation patterns with depth for most elements are still reasonably clear. Those showing the greatest amount of scatter, namely Na, K, Mn, Ba and Rb, should be avoided in making any petrogenetic deductions, while the immobile elements Ti, P, Nb, Zr and Y, and also the light rare earth elements La, Ce and Nd, appear to have been little affected by any post-extrusive processes.

4. DISCRIMINATION TECHNIQUES

4.1. Analysis of Multivariate Data

Most mineralogical and petrological observations are multivariate. Composition data, whether obtained by chemical, modal or other analysis, are multivariate. Multivariate data also occur because the science of geology often involves the dimensions of space and time (Koch and Link, 1971).

Univariate observations can be plotted on a straight line (Fig. 4.1a). If the straight line is divided into class intervals, the number of observations in each interval can be counted, and a histogram can be formed, the plotting of which requires two dimensions.

Bivariate observations can be plotted on a scatter diagram in two dimensions (Fig. 4.1b). If the diagram is divided into cells, the number of observations in each cell can be counted and plotted on a histogram. This histogram takes three dimensions to plot and may be represented by a contour map.

Trivariate observations can be plotted on a scatter diagram in three dimensions (Fig. 4.1c). The most usual configuration of points defines an ellipsoid. If the three-dimensional space is divided into cubes or other solid figures, the number of observations in a histogram would require four dimensions to plot. Trivariate observations of constant sum can also be plotted on triangular coordinates.

Observations with four or more variables cannot be represented on ordinary plots. Discriminant analysis is designed to cater for this type of situation.

4.2. Theory of Discriminant Analysis

Discriminant analysis is a set of methods which yields a function for classifying multivariate observations into two or more groups through developing one or more linear combinations of the variables that

Figure 4.1 Plots of observations in various dimensions
(from Koch and Link, 1971)

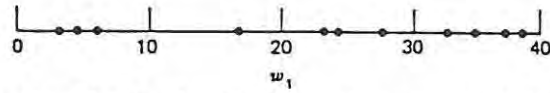


Fig 4.1a Plot of univariate observations on a straight line.

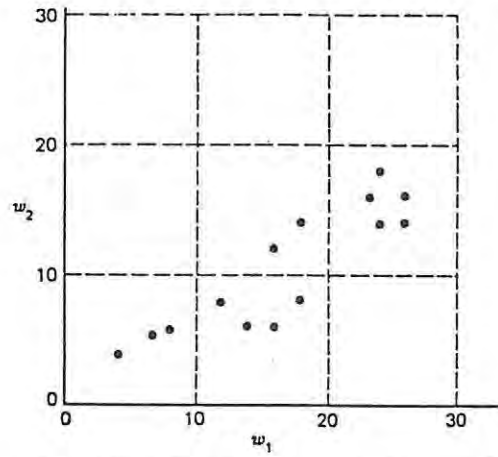


Fig 4.1b Plot of bivariate observations on a scatter diagram in two dimensions.

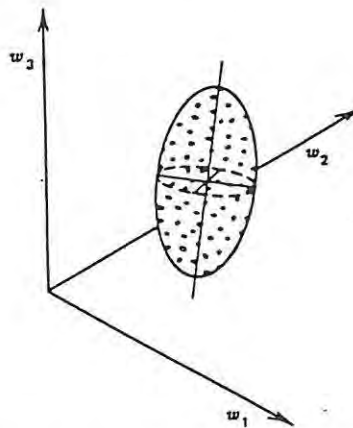


Fig 4.1c Plot of trivariate observations on a scatter diagram in three dimensions.

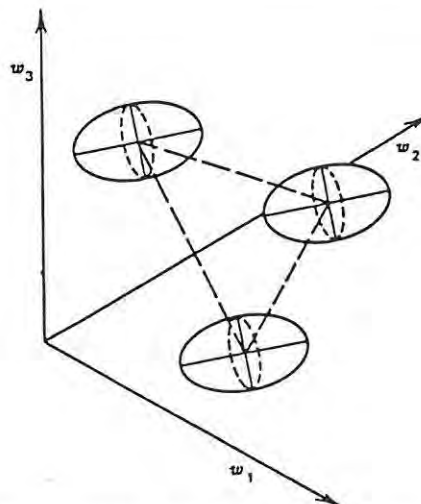


Fig 4.1d Three ellipsoids representing clouds of observations that are classified by discriminant analysis.

comprise the observations. In principle, ellipsoid centres are located and the directions and lengths of the distances connecting paired ellipsoid centres are examined. Figure 4.1d shows three ellipsoids and three dashed lines denoting paired distances.

Discriminant analysis is one of the statistical options forming part of the SPSS (Statistical Package for the Social Sciences) program (Release 3), developed at the University of Nottingham, England. Klecka (1975), who designed and programmed the discriminant analysis package used in this study, stated that "the statistical theory of discriminant analysis assumes that the discriminating variables (which in this study are major and trace element abundances, together with selected ratios of these) have a multivariate normal distribution and that they have equal variance-covariance matrices within each group. In practice, the technique is very robust and these assumptions need not be strongly adhered to." In other words, the data used need not necessarily be strictly normally distributed, and the technique can therefore be applied perfectly satisfactorily to geological data. (The distribution of the data used in this study is discussed in chapter 5). In addition, Koch and Link (1971) explained that they omitted several multivariate methods from their book because they require data more nearly normal than those usually found in geology. The obvious implication is that discriminant analysis, which is included in the book, does not require strictly normal data.

The SPSS program was run on the Rhodes University ICL 1904S computer. The data used are shown in Figure 4.2. Line 3 lists the discriminating variables, which are read in according to the format in line 5. Lines 8 to 10 define the ratio calculations which need to be performed. The normalised, volatile-free data for each sample fill two lines - the first contains the sample name and number (S and SANC), followed by a number (See Figure 4.3) denoting the formation to which the sample belongs (FMTS), and the major element analyses (in weight percent), while the second line lists trace element concentrations (in parts per million). At the end of the sample data, discriminant analysis is requested (lines 661 to 663) giving the number of groups into which the data is divided, and the variables which must be used to calculate the discriminant functions. The "OPTIONS" and "STATISTICS" lists allow the user to choose which tables and statistics are to be printed out.

Figure 4.2 Data for SPSS program.

```

0001 RIJN NAME      TEST
0002 FILE NAME     FILE
0003 VARIABLE LIST S,SANO,FMTS,SIO2,TIO2,AL203,FE203,MNO,MGO,CAO,NA2O,K2O,P2O5,
0004                GAP,BA,NB,ZR,V,SR,RB,ZN,CU,HI,CO,CR,V,LA,CE,M
0005 INPUT FORMAT   FIXFO(A3,F3.0,F2.0,10F6.0/F7.0,F5.0,3F4.0,2F5.0,4F4.0,F5.7,4F4.0)
0006 INPUT MEDIUM  CARD
0007 NO. OF CASES  324
0008 COMPUTE        TI:TI=(TIO2*5495)/7R
0009 COMPUTE        P:TI=(P205*4363)/(TIO2*5495)
0010 COMPUTE        ZR:P=7R/(P205*4363)
0011 TRANSFORM
0012 READ INPUT DATA
0013 DR- 110 57.76 1.084 14.58 11.55 0.29 3.51 7.35 3.33 0.50 0.255 0.70
0014      204. 8.5173.30.1 286. 11.9125. 2R. 19. 55. 5.297. 27. 48. 28.
0015 DR- 210 58.44 1.041 13.85 11.48 0.15 3.54 6.82 3.03 1.37 0.255 0.70
0016      395. 7.5167.26.9 460. 51.0100. 16. 23. 54. 6.261. 25. 59. 24.
0017 DR- 310 58.78 1.004 13.66 11.72 0.14 3.17 6.79 3.17 1.34 0.257 0.70
0018      484. 7.1171.27.7 465. 44.5103. 42. 25. 53. 12.245. 27. 50. 28.
0019 DR- 4 9 68.32 0.849 12.74 7.35 0.11 2.57 2.05 3.65 2.16 0.271 0.70
0020      846.12.9256.28.9 214. 56.9 91. 10. 12. 26. 12.106. 44. 77. 37.
0021 DR- 5 9 70.84 0.809 12.50 5.66 0.08 1.73 1.45 3.42 3.72 0.278 0.70
0022      853.13.4277.27.3 111. 81.5 87. 9. 9. 16. 2. 53. 46. 94. 40.
:
:
0653 AR-496 1 53.74 1.201 15.75 12.23 0.13 4.07 9.17 2.46 1.38 0.264 0.70
0654      175. 7.8182.28.7 196. 48.2 97.129.157. 63. 51.258. 27. 57. 29.
0655 AR-497 1 51.76 1.210 15.75 13.67 0.16 5.78 7.76 3.34 0.34 0.229 0.70
0656      97. 7.4181.24.2 242. 10.9106.137.213. 64. 213.272. 16. 48. 22.
0657 AR-498 1 56.77 1.146 15.16 11.41 0.14 5.61 5.64 3.52 0.37 0.233 0.70
0658      202. 8.1189.32.0 553. 5.7105. 95.166. 69. 56.208. 25. 50. 27.
0659 AR-500 1 56.74 1.018 14.29 11.16 0.18 4.96 5.86 3.29 2.29 0.220 0.70
0660      705. 8.1159.28.0 539. 51.8 93. 88. 97. 56. 83.188. 23. 52. 27.
0661 DISCRIMINANT  GROUPS=FMTS(1,10)
0662                VARIABLES=SIO2 TO P205,BA TO ND,TI:ZR,P:TI,ZR:P/
0663                ANALYSIS=SIO2 TO P205,BA TO ND,TI:ZR,P:TI,ZR:P/
0664                METHOD=MAHAL/
0665                PRIORS=EQUAL/
0666 OPTIONS        3,5,6,7,11,12,17,18,19
0667 STATISTICS     1,2
0668 FINISH
    
```

There are two aspects to discriminant analysis: (1) classifying groups of observations, and (2) classifying a single observation into one of a number of predefined groups.

Discriminant analysis begins with the desire to statistically distinguish between two or more groups of cases (geological units, in the current study). The mathematical objective of discriminant analysis is to weight and linearly combine the discriminating variables in some fashion so that the groups are forced to be as statistically distinct as possible. These "discriminant functions" are of the form

$$D_i = d_{i1}Z_1 + d_{i2}Z_2 + \dots + d_{ip}Z_p \dots\dots\dots(1)$$

where D_i is the score on discriminant function i ,
 the d 's are weighting coefficients, and
 the Z 's are the standardised values of the p discriminating variables used in the analysis.

The maximum number of functions which can be derived is either one less than the number of groups, or equal to the number of discriminating variables, if there are more groups than variables. Ideally, the D's for the cases within a particular group will be fairly similar. The functions are formed in such a way as to maximise the separation of the groups. Once the discriminant functions have been derived, the two research objectives of this technique, analysis and classification, may be pursued.

During the analysis phase, if there are more than two groups, it may be possible to obtain satisfactory discrimination with fewer than the maximum number of functions. Since the discriminant functions (i's) can be thought of as the axes of a geometric space, they can be used to study the spatial relationships among the groups. The weighting coefficients (d's) serve to identify the variables which contribute most to differentiation along the respective dimension or function.

4.2.1 Analytical features of discriminant analysis

4.2.1.1 Discriminant functions

Two measures are provided for judging the importance of the discriminant functions. One of these is the relative percentage of the eigenvalue associated with the function. It is computed in the process of deriving the discriminant function and is a measure of the relative importance of the function. A portion of the program output in Table 4.1 shows that the first discriminant function, with an eigenvalue of 72.36 percent, is very much more important than the remaining functions. The relative percentages of the last three functions show that they are fairly insignificant compared with the others.

The second measure for judging the importance of a discriminant function is its associated canonical correlation. This is a measure of association between the discriminant function and the set of variables which define the group memberships, i.e. it measures the function's ability to discriminate among the groups. In Table 4.1 it can be seen that the first six discriminant functions are each well correlated with the groups, while the next two are moderately correlated.

Table 4.1 Portion of discriminant analysis output from the SPSS program.

DISCRIMINANT FUNCTION	EIGENVALUE	RELATIVE PERCENTAGE	CANONICAL CORRELATION	:	FUNCTIONS DERIVED	WILKS' LAMBDA	CHI-SQUARE	DF	SIGNIFICANCE
1	50.03636	72.36	0.990	:	0	0.0000	3559.226	208	0.000
2	6.65657	9.63	0.932	:	1	0.0004	2365.700	175	0.000
3	4.60026	6.65	0.906	:	2	0.0032	1747.906	144	0.000
4	3.69825	5.35	0.887	:	3	0.0177	1225.033	115	0.000
5	2.31371	3.35	0.836	:	4	0.0830	755.460	88	0.000
6	1.34362	1.94	0.757	:	5	0.2750	391.847	63	0.000
7	0.32912	0.48	0.498	:	6	0.6444	133.357	40	0.000
8	0.16751	0.24	0.379	:	7	0.8565	47.005	19	0.000

It has already been stated that the maximum number of discriminant functions which can be derived is either one less than the number of groups, or equal to the number of discriminating variables, whichever is the smaller. This stems from basic geometric principles, i.e. two points define a line, three define a plane, etc. An exception to this geometric rule that x points define $(x-1)$ dimensions is the "degenerate" situation of, for example, three points falling on the same line, four on the same plane, etc. That is, the last point is situated such that it does not add a new dimension. Thus, in discriminant analysis, two functions may be adequate for describing four groups, for example.

A second criterion for eliminating discriminant functions is to test for the statistical significance of discriminating information not already accounted for by the earlier functions. Thus Wilks' lambda is computed as each function is derived, starting with no functions. Lambda is an inverse measure of the discriminating power in the original variables which has not yet been removed by the discriminant functions, i.e. the larger lambda is, the less information remains. Lambda can be transformed into a chi-square statistic for an easy test of statistical significance.

Table 4.1 shows that Wilks' lambda was zero before any functions were removed, indicating that immense discriminating power exists in the variables being used. After some of this discriminating power has been removed by placing it into the first discriminant function, lambda increases very slightly to 0.0004, and the chi-square shows that a statistically significant amount of discriminating information still exists. It is only after the seventh function has been derived that a fairly large (0.8565) but still significant lambda is found. A significance of zero means that a lambda of this magnitude or smaller has no probability of occurring due to the chances of sampling, even if there was no further information to be accounted for by an eighth function in the population. Clearly, an eighth function is statistically significant in this case.

4.2.1.2 Discriminant function coefficients

The standardised discriminant function coefficients are shown in Table 4.2. These coefficients correspond to the values of the d_{ij} 's given in equation (1). They are used to compute the discriminant score for a case in which the original discriminating variables are in standard form, i.e. Z scores. The discriminant score (D) is computed by multiplying each discriminating variable by its corresponding coefficient and adding together these products. Each case has a separate score on each function.

Table 4.2 Standardised discriminant function coefficients.

STANDARDIZED DISCRIMINANT FUNCTION COEFFICIENTS

	FUNC 1	FUNC 2	FUNC 3	FUNC 4	FUNC 5	FUNC 6	FUNC 7	FUNC 8
SI02	0.01420	0.47917	0.05585	-0.64850	-1.59610	-2.05176	-4.55288	5.83264
TI02	-0.05306	0.48913	-2.59056	1.28369	-0.27572	0.50511	1.89024	1.61214
AL203	-0.02857	0.18200	-0.29379	-0.34104	-0.45153	-0.45999	1.50384	1.80944
FE203	-0.07083	0.05096	-0.02194	-0.62449	-1.26712	-0.90223	2.26101	1.78941
MGO	0.04858	0.70829	0.31024	0.08662	-0.24321	-0.51184	2.52677	3.38678
CA0	-0.01952	0.35838	0.00336	-0.39002	-0.76979	-1.07839	1.54588	2.32507
NA20	-0.00516	0.15776	0.06557	-0.16927	-0.30754	-0.56690	0.91262	0.35363
P205	0.82454	-1.27344	2.85151	-1.88591	1.86502	2.01634	-2.77942	-0.56699
BA	-0.02000	0.14329	-0.02743	-0.05875	-0.20776	-0.09444	0.36333	0.37877
NR	-0.06470	0.47402	0.81368	1.25941	0.00331	-1.86725	-0.33422	-0.77026
ZR	0.61879	-0.45065	0.06809	-0.55609	1.20833	0.08972	-0.06953	0.25478
Y	-0.52675	0.35944	0.61153	0.22145	-2.18198	0.52812	-0.59835	0.73190
SR	-0.01427	0.06170	0.07040	0.16577	0.06920	-0.09702	0.11568	-0.11587
RB	0.01861	-0.06160	0.13862	-0.13160	-0.14711	-0.29581	0.69696	-0.06266
ZN	-0.03683	-0.12085	-0.00979	-0.05531	-0.07869	-0.19551	-0.63014	0.21964
CU	0.04806	0.08034	-0.16758	0.18286	0.11342	0.15855	-0.08956	-0.13039
NI	-0.07725	-0.03402	-0.05027	0.30152	0.36824	-0.29235	0.40721	0.35942
CO	0.19367	0.25804	0.45292	0.64371	-0.05640	0.20750	-1.48416	-1.81835
CR	-0.09352	-0.07113	-0.51375	-0.72429	-0.30963	-0.25200	-0.06467	0.08923
V	-0.05557	0.58136	0.26195	-0.51014	0.65303	-1.10041	0.73050	-0.47199
LA	0.04262	-0.17579	-0.65132	-0.15067	0.42185	-0.22494	-0.39450	-0.11770
CE	-0.13492	0.25026	0.55236	0.71804	0.41094	1.04089	0.90160	0.28076
ND	0.02295	-0.07142	0.15589	-0.12987	-0.20543	-0.92546	-0.23079	0.05327
TI:ZR	-0.21314	-0.26445	0.45245	0.53248	-0.35447	0.06273	1.61407	-0.56624
P:TI	-1.44448	0.05414	-1.87460	0.42604	-1.20211	-0.56442	3.27993	-2.72300
ZR:P	-0.24841	-0.24336	0.02965	0.14924	0.26997	0.46076	0.51714	-1.44628

The coefficients have been derived in such a way that the discriminant scores produced are in standard form, i.e. over all cases in the analysis, the score from one function will have a mean of zero and a standard deviation of one. Thus, any single score represents the number of standard deviations that case is away from the mean for all cases on the given discriminant function.

Averaging the scores for all cases within a particular group will give the group mean on the respective function. For a single group, the means on all the functions are referred to as the group centroid, which is the most typical location of a case from that group in the discriminant function space. A comparison of the group means on each function reveals how far apart the groups are along that dimension (see Table 4.3).

Table 4.3 Centroids of groups in reduced space.

CENTROIDS OF GROUPS IN REDUCED SPACE									
		FUNC 1	FUNC 2	FUNC 3	FUNC 4	FUNC 5	FUNC 6	FUNC 7	FUNC 8
GROUP	1	0.75759	0.08470	-1.09620	1.20585	0.39884	-0.07252	-0.99785	-0.04265
GROUP	2	-1.09770	0.88147	-0.14300	-0.23024	0.37493	0.87255	0.04938	-0.07578
GROUP	3	-2.18379	-0.07184	1.69889	1.98341	-0.05452	-1.55854	0.04538	0.19110
GROUP	5	0.95347	-0.30109	1.69197	-0.58730	0.12164	0.39064	-0.36792	0.16622
GROUP	6	1.00930	-0.04622	0.26384	0.87191	-0.74439	0.20837	0.63293	-0.82804
GROUP	7	0.97197	-0.16909	-0.68555	1.15109	-0.44052	0.72004	0.93037	1.06922
GROUP	8	-1.98780	-1.65217	-0.79485	-0.94412	-5.15565	0.71662	-1.03364	0.14768
GROUP	9	-0.59037	-2.36177	-0.60938	-0.50115	0.90956	-0.07180	0.22786	-0.11255
GROUP	10	0.30953	0.62584	-0.42967	-0.87941	-0.15180	-1.00781	0.13596	0.05428

The standardised discriminant function coefficients themselves are of great analytical importance. When the sign is ignored, each coefficient represents the relative contribution of its associated variable to that function. The sign merely denotes whether the variable is making a positive or negative contribution. In Table 4.2 it can be seen that for the first function, which is by far the most important, the greatest contribution is made by the ratio P:Ti, followed by P_2O_5 , Nb and Zr. Y is almost half as important as Zr, and this is followed by two more ratios, Zr:P and Ti:Zr, etc.

On the second function the most important variables, in order, are P_2O_5 , MgO, P:Ti, V, Zr, TiO_2 and SiO_2 . If all the functions are considered together, with values on each function being weighted in accordance with their eigenvalues, the order of importance of the discriminating variables is as follows: P_2O_5 , P:Ti, SiO_2 , TiO_2 , V, Ce, Y, Ti:Zr, Nb, Fe_2O_3 , CaO, Al_2O_3 , MgO, Co, Zr, Zr:P, Na_2O , Cr, La, Ni, Nd, Rb, Zn, Cu, Ba, Sr.

4.2.2 Classification in discriminant analysis

The use of discriminant analysis as a classification technique comes after the initial computation. By classification is meant the process of identifying the likely group membership of a case when the only information known is the case's values on the discriminating variables. Another use of classification is in testing the adequacy of the derived discriminant functions. By classifying the cases used to derive the functions in the first place, and comparing predicted group membership with actual group membership, one can empirically measure the success in discrimination by observing the proportion of correct classifications.

Classification is achieved through the use of a series of classification functions, one for each group. They are derived from the pooled within-groups covariance matrix and the centroids for the discriminating variables. The resulting classification coefficients are multiplied by the raw variable values, summed together and added to a constant as follows:

$$C_i = c_{i1}V_1 + c_{i2}V_2 + \dots + c_{ip}V_p + C_{i0} \dots \dots \dots (2)$$

where C_i is the classification score for group i ,
the c_{ij} 's are the classification coefficients,
 c_{i0} is the constant, and
the V 's are the raw scores on the discriminating variables.

There is always a separate equation for each group. The case would be classified into the group with the highest score, i.e. for which it has the greatest probability of membership. If required, adjustments are

made on the basis of prior probabilities being proportional to the number of cases in each group, i.e. cases will more likely be assigned to a larger group. For the current study the SPSS program was run using both equal and proportional probabilities. The result was the same in both cases.

Table 4.4 shows an excerpt from the SPSS program output in which the discriminant scores for each case are printed together with the predicted group membership. The largest probability $P(G/X)$ is also printed and if the value for the second highest probability is greater than .0005, it too is printed. Another useful probability $P(X/G)$ which appears on the output is the probability that a member of the predicted group would be as far from the centroid as the case being considered.

Table 4.4 Portion of output showing probabilities of group membership and discriminant scores.

ACTUAL SEQUENCE	GROUP	HIGHEST PROBABILITY		2ND HIGHEST		DISCRIMINANT SCORES								
		GROUP	$P(X/G)$	$P(G/X)$	GROUP	$P(G/X)$	FUNC 1	FUNC 2	FUNC 3	FUNC 4	FUNC 5	FUNC 6	FUNC 7	FUNC 8
1.	10	10	1.000	1.000			0.355	0.636	-0.320	-1.222	-0.111			
2.	10	10	1.000	1.000			-2.315 0.347	-0.819 0.535	0.220 -0.012	-1.035	-0.149			
3.	10	10	1.000	1.000			-1.308 0.491	0.365 0.043	-0.466 -0.255	-1.168	-0.283			
4.	9	9	0.995	1.000			-1.551 -0.410	-0.446 -1.391	-0.550 -0.511	-0.820	0.760			
5.	9	9	1.000	1.000			-0.836 -0.598	0.414 -2.151	0.586 -0.762	-0.620	1			

4.3 Application of Discriminant Analysis to the Witwatersrand Triad Samples

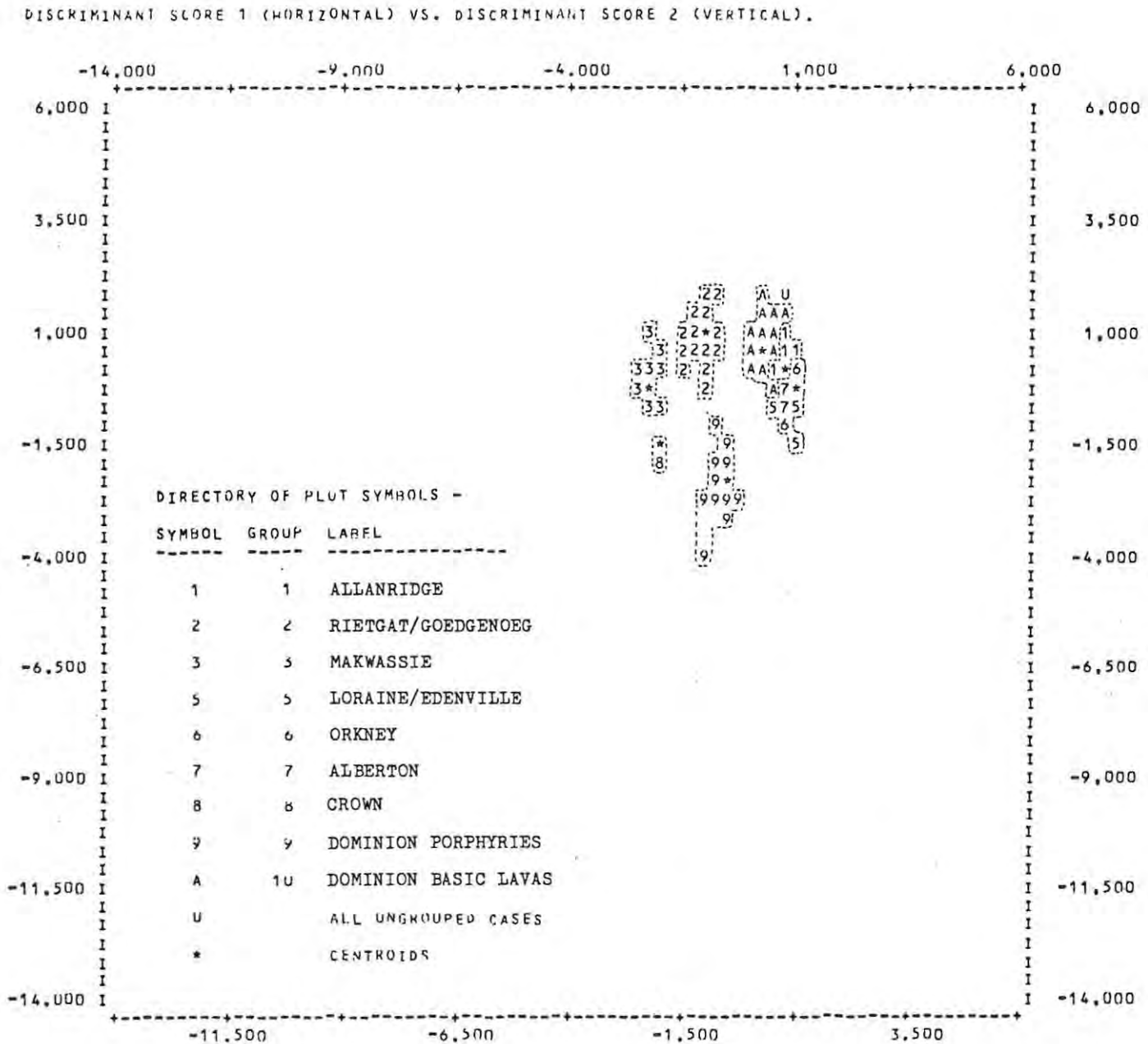
4.3.1 Evaluation of discriminant plots and grouping of samples

The SPSS program produces plots of the cases along the first two discriminant function continuums. Because the functions are arranged in order of decreasing importance, the graph will represent the two most important functions.

It is essential in applying discriminant analysis to a set of data, that a preliminary classification of some or all of the data into groups must exist. The discriminant functions are derived such that the first function separates the groups as much as possible. The second function separates them as much as possible in an orthogonal direction given the first separation. The third function provides maximal separation in another orthogonal direction, etc. The end result is that the groups are as distinct as possible given the original discriminating variables. In this way the discriminant functions can be considered as defining geometric axes in space in which each case and group centroid are points. In Figure 4.3 the first two discriminant functions are plotted. It can be seen that groups 3, 8 and 9, the Makwassie Formation, the Crown Formation and the Dominion Group porphyries respectively, are separated completely from one another and from the rest of the points by these two functions. Groups 2 and 4 (the Rietgat and Goedgenoeg Formations) together form a separate cluster, but in the two dimensions shown there is considerable overlap between the groups themselves. The original listings of group centroid positions confirmed that the centroids for groups 2 and 4 were very closely spaced along function 1 (-1.02745 and -1.05813 respectively) and were almost coincident along function 2 (0.88064 and 0.87991 respectively). On function 3, however, a more satisfactory separation was achieved (-0.28846 and 0.07967), and also on most of the other functions. Due to the limitations of the graph these were not evident graphically. It must be remembered, though, that because the functions are arranged in order of decreasing importance, a given difference between group centroids on the third or fourth function is not as meaningful as the same difference on the first function.

The separation of the two centroids in other dimensions was thus still not sufficient to eliminate all overlap between the two groups. This was shown in the original output, where four members (12.5%) of group 2 (the Rietgat Formation) were classified as members of group 4 (the Goedgenoeg Formation) and eight members (25%) of group 4 were classified as group 2 types. This relatively large amount of overlap suggests that the differences between these two groups may not be as significant as those between the other groups, and the possibility of classifying them as a single group was implicitly suggested by these results.

Figure 4.3 Plot of discriminant score 1 (horizontal) vs discriminant score 2 (vertical)



Petrogenetic studies undertaken by M.P. Bowen (in prep.) on the same samples have also led to the conclusion that the Rietgat and Goedgenoeg Formations cannot be separated geochemically. He has therefore treated them as a single group in all subsequent work. Further investigation in the current study has supported this contention and Table 4.5 shows the classification of cases after these two groups had been merged. Group 2, which now contains both the Rietgat and the Goedgenoeg Formations, shows a 100% correct classification of cases.

The remaining five groups form a cluster on the right-hand side of Figure 4.3. Group 1, the Allanridge Formation, overlaps with a number of other groups, but Table 4.5 shows that total separation is achieved when taking the other dimensions into account. Group 10 (the Dominion Group basic lavas), represented by an 'A' in the diagram, falls on the left-hand side of the cluster, overlapping groups 1 and 5 slightly. Table 4.5 shows, however, that separation of group A from the others is

Table 4.5 Predicted group memberships of previously grouped cases.

PREDICTION RESULTS -

ACTUAL GROUP	NO. OF PREDICTED CASES	GROUP MEMBERSHIP									
		GP. 1	GP. 2	GP. 3	GP. 5	GP. 6	GP. 7	GP. 8	GP. 9	GP. 10	
GROUP 1	35.	35.	0.	0.	0.	0.	0.	0.	0.	0.	0.
		100.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%
GROUP 2	67.	0.	67.	0.	0.	0.	0.	0.	0.	0.	0.
		0.0%	100.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%
GROUP 3	17.	0.	0.	17.	0.	0.	0.	0.	0.	0.	0.
		0.0%	0.0%	100.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%
GROUP 5	46.	0.	0.	0.	45.	1.	0.	0.	0.	0.	0.
		0.0%	0.0%	0.0%	97.8%	2.2%	0.0%	0.0%	0.0%	0.0%	0.0%
GROUP 6	51.	0.	0.	0.	0.	30.	1.	0.	0.	0.	0.
		0.0%	0.0%	0.0%	0.0%	58.8%	3.2%	0.0%	0.0%	0.0%	0.0%
GROUP 7	19.	0.	0.	0.	0.	0.	19.	0.	0.	0.	0.
		0.0%	0.0%	0.0%	0.0%	0.0%	100.0%	0.0%	0.0%	0.0%	0.0%
GROUP 8	6.	0.	0.	0.	0.	0.	0.	6.	0.	0.	0.
		0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	100.0%	0.0%	0.0%	0.0%
GROUP 9	32.	0.	0.	0.	0.	0.	0.	0.	32.	0.	0.
		0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	0.0%	100.0%	0.0%	0.0%
GROUP 10	64.	0.	0.	0.	0.	1.	0.	0.	0.	63.	0.
		0.0%	0.0%	0.0%	0.0%	1.4%	0.0%	0.0%	0.0%	98.6%	0.0%
UNGROUPED CASES	2.	0.	0.	0.	0.	0.	1.	0.	0.	0.	1.
		0.0%	0.0%	0.0%	0.0%	0.0%	50.0%	0.0%	0.0%	0.0%	50.0%

PERCENT OF "GROUPED" CASES CORRECTLY CLASSIFIED: 99.07%

virtually perfect, and only two samples (2.8%) belonging to Group 10 were classified as members of other groups. DB-28 had a probability of .985 of being a member of the Allanridge Formation, with a probability of only .011 of belonging to the Dominion Group basic lavas. DB-42 was classified as a Loraine-Edenville member with a probability of .999, and a very weak probability of .001 of belonging to the Orkney Formation.

4.3.2 Aberrant and unclassified samples in the Dominion Group

An examination of the stratigraphic position (see Figure 2.2a) of DB-42 showed that it comes from borehole DSF-7 at a depth of 133m from the surface. Between this position and the surface a very clearly intrusive horizon was logged. The intrusion extends from 76.9m to 103.9m below the surface. It has a fairly coarse-grained texture with finer-grained chilled margins. Some alteration in colour of the overlying rock is evident near the upper contact. The intrusion was sampled in four places, at 103, 88, 82 and 78m below the surface, one sample at each margin and two in the centre. During the classification phase of the program these four samples, IN-43, IN-44, IN-45 and IN-46, were all unambiguously allocated to the Loraine-Edenville Formations. Probabilities of membership were 1.0 for all four samples, and they were therefore prefixed "KL".

DB-42, the single sample under consideration, was classified in exactly the same way. At the time of sampling, it was not recognised as an intrusive horizon due possibly to the absence of chilled margins in a thin unit, but an examination of its major and trace element analyses in relation to the nearby intrusive and the other basic lavas suggested strongly that this might be so. The major element values on their own are not conclusive, but the trace elements, especially Zr, Ni, Cr and the rare earths, provide strong evidence for the possible relationship between the five samples. As a result, DB-42 was reclassified and prefixed "KL" to indicate this relationship, but in the absence of unequivocal field relationships, final judgement must be reserved.

The second errant sample DB-28 occurs in the same borehole as the aforementioned intrusion, but at a depth of 487m. It shows no obvious differences in major and trace element contents from the surrounding Dominion Group basic lavas, but discriminant analysis has shown that it is more likely to belong to the Allanridge Formation lavas.

The cryptic differences between the two aforementioned sets of lavas come to light when the ratios Ti:Zr and Ti:P are plotted against one another. Figures 4.6 and 4.7 shows that DB-28 lies near the Allanridge samples and not with the rest of the Dominion Group basic lavas. It was therefore considered that this sample is either an intrusive one of Allanridge

origin, or that some Allanridge-like lavas occur in the Dominion Group. As with sample DB-42, further work is necessary to resolve this.

A second borehole, DRS-6, six kilometers to the south of DSF-7, also contains a number of intrusive horizons, from which samples IN-86, IN-87 and IN-90 were taken. Classification of these intrusives of unknown affinity indicated that all three had a probability of 1.0 of belonging to the Loraine-Edenville Formations, and their prefixes were therefore also changed accordingly.

Sample DR-91, taken near the top of the same hole, proved to be a somewhat enigmatic sample. The horizon was not regarded as an intrusive, but a flow, as evidenced by the presence of prominent amygdaloidal zones at both the top and the bottom of the unit. The geochemistry differs markedly from all the other Dominion Group basic lavas, but it was considered unjustified to omit the analysis from the rest of the data without a valid reason simply because it did not "fit" well. The explanation was provided almost simultaneously and independently from two different sources. During field studies carried out in the vicinity of borehole DRS-6, M.P. Bowen (pers. comm., 1982) discovered that lavas of the Platberg Group lapped onto Dominion Group lavas in this area, with the usual intervening rocks missing. At about the same time, sample DR-91 was classified by discriminant analysis as either a member of the Rietgat Formation of the Platberg Group with a probability of .962, or as part of the Goedgenoeg Formation, also of the Platberg Group, with a much lower probability of .038. With this conclusive evidence, the sample was renamed PR-91.

With these eliminations the Dominion Group basic lavas would therefore now be completely separable from all the other groups, as shown in Table 4.5, except that a new sample, DB-97, has been misclassified as a result of changing parameters. Investigation showed, however, that there was no justifiable reason for excluding it from the Dominion Group basic lavas.

4.3.3 Aberrant samples in the Klipriviersberg Group

Users of discriminant analysis should be wary of developing a blind acceptance of all reclassifications without carefully investigating the reasons for this. The lavas of the Klipriviersberg Group provide a

salutary lesson in this respect. Groups 5, 6 and 7 represent the following formations within the Klipriviersberg Group respectively: the Loraine-Edenville, Orkney and Alberton Formations. Table 4.5 shows that group 7, the Alberton Formation is separated completely from the others. However, sample KO-356 from the Orkney Formation in borehole W-20 was reclassified by the SPSS program as an Alberton sample with .987 probability against a probability of .013 that it was an Orkney sample. Examination of the stratigraphic column for that borehole (Figure 2.2b) shows that the sample lies well within the Orkney Formation.

Two possibilities may be considered to explain this. Firstly, it is conceivable that an interfingering of two different lava types might exist, i.e. extrusion of the underlying Alberton lavas may not have been entirely completed before the Orkney-type emissions began from a different source. Sample KO-356 may thus represent one of the final pulses of extrusion from a virtually exhausted Alberton source. Alternatively, the close similarity between the chemistry of the Alberton and Orkney lavas may be the reason for this classification. The sample has the highest SiO_2 value of all the Orkney samples analysed and in this respect it resembles the Alberton lavas. All Orkney samples, however, have Zr values between 90 and 110 ppm, with all Alberton samples (except one at 60 ppm) between 106 and 137 ppm Zr. Sample KO-356, with 96 ppm Zr, clearly has Orkney affinities, and this is corroborated by Zr vs TiO_2 (Figure 4.8) and Zr vs Fe_2O_3 plots, in which this sample lies well within the Orkney field. It was therefore decided that KO-356 should remain as an Orkney member and not be reclassified.

In group 5, the Loraine-Edenville Formations, two samples were misclassified. KL-133 had a probability of .940 of belonging to the Orkney Formation and only .059 of being part of its expected formation. This sample was taken in the lower altered zone of borehole WS-5, about 37m above the contact between the Orkney and Loraine Formations. The top of the Orkney Formation is defined as the top of the flow of the upper purple band (Winter, 1976 and 1965b, after Pienaar, 1956). This marker horizon is absent in WS-5 and the division was made on a geochemical basis. In some boreholes e.g. JBF-1 and WS-4 this horizon does coincide with the geochemical boundary within the limits of sampling intervals, but not in others, e.g. W-20. The problem of membership is resolved unequivocally by TiO_2 values, which lie between .400 and .878 weight

percent for Loraine-Edenville lavas, and between .879 and 1.097 for Orkney lavas. KL-133, with a TiO_2 content of .780, must therefore belong to the Loraine-Edenville Formations. Zr vs TiO_2 (Figure 4.8) and Zr vs Fe_2O_3 plots confirm this, as does the Cr content of 238 ppm, as no Orkney lava contains more than 97 ppm Cr. This misclassification must be again be ascribed to either the similar chemical characteristics shown by the different formations within the Klipriviersberg Group, or to interfingering of two lava types.

The second erroneous sample, KL-471 was classified unambiguously as a Dominion Group basic lava (probability = 1.0). This sample was taken from the uppermost flow of the Klipriviersberg Group lavas, which is overlain by Kameeldoorns sediments. A possible explanation for its presence at a much higher stratigraphic level than expected is offered. An examination of the stratigraphic log of JBF-1, the borehole from which KL-471 was collected, showed that the thickness of Kameeldoorns sediments in this hole (85m) was fairly substantial compared with that in some of the adjacent boreholes, such as SH-1 (20m). This suggests the proximity of a fault, as Winter (1970) stated that "the greatest depths at which the Kameeldoorns Formation is found are adjacent to major faults and this formation subsided as clastic wedges against these. The coarsest sediments adjoin the faults." In support of this is the fact that virtually the whole intersection of the Kameeldoorns sediments in this hole comprises conglomerate. Winter (1965b, Fig. 5) showed that there is indeed a fault in this area, the Jersey Fault, which strikes NE-SW and lies immediately to the north-west of JBF-1 (see Figure 2.1), and against which the Kameeldoorns sediments are abruptly truncated. The north-western side of the fault is upthrown. Evidence of this can be seen in borehole JY-8, which contains rocks of the Central Rand Group and the Jeppestown Subgroup of the West Rand Group. The stratigraphic log stops in the Rietkuil Formation, so it is not known at what depth the Dominion Group first occurs.

Movement along the Jersey Fault has therefore caused older lithologies to be brought closer to the surface in the north west. It is possible that Dominion Group basic lavas were exposed at the fault scarp and that fragments may have broken away and become incorporated with the lavas and sediments filling the graben-like depository adjacent to the fault. In this manner the position of KL-471, a lava showing Dominion Group

chemical characteristics, found in the uppermost flow of the last Klipriviersberg lavas extruded before the onset of sedimentation, may be explained. Major and trace element values and interelement plots are, however, inconclusive. There appears to be a bias towards the Dominion lavas, but some of the evidence is contradictory. As a result, it was decided to leave this sample out of the analytic phase of the discriminant analysis program, until further work is able to resolve its affinity.

4.3.4 Other aberrant samples

Four unknown samples still remain to be classified. IN-419 and IN-421 both occur in borehole JHA-1, in which a very unusual succession is present (see Figure 2.2b). From the surface down to 876m the lithologies are quite normal: rocks of the Karoo Supergroup occur down to 165m and are underlain by the Allanridge and Bothaville Formations. From 876m to 2658m, where pre-Ventersdorp rocks are reached, the sequence consists of Kameeldoorns-type sediments interbedded with flows of porphyritic and non-porphyritic lavas. A 4m intersection of lava with Makwassie-type phenocrysts occurs between 1041 and 1045m. The thickness proportion of lavas to sediments is about 3:1. The sediments include conglomerates, shales, dolomites and cherts. No particular sequence is evident - for example, one conglomerate horizon overlies a shale-dolomite-chert succession, suggesting a number of episodes of movement along the Jersey Fault (see Figure 2.1). Some of the volcanic rocks have been logged by company geologists as "possibly intrusive". It was from two such horizons that IN-419 and IN-421 were taken. Four other lava samples analysed from the intercalated succession (viz. PG-406, PG-417, PR-426 and PG-433) were classified with the Goedgenoeg-Rietgat group of lavas. IN-421 was also allocated to the Rietgat Formation, and its prefix was thus changed to accommodate this classification. IN-419 was named as an Alberton lava with a probability of 1.0, but the probability that any other member of the Alberton Formation would be as far from the centroid as IN-419, was 0. This implies that IN-419 lies outside the bounds of the Alberton Formation, but closer to this formation than to any other. Both the Transvaal and the Karoo Supergroups outcrop in the area, and the possibility must be entertained that this intrusive horizon might represent one of these. This sample was therefore excluded from the analysis part of the program on the grounds of its tenuous association with the lavas under consideration in this project.

KO-121 was the only sample analysed which contained a large number of amygdalites. The predominant filling is calcite, which also occurs together with epidote in the groundmass. The major element analysis for this sample shows a very high CaO value (19.61 percent), as expected, and a relatively high Al_2O_3 value (16.75 percent), resulting in a dilution of the other major elements. Trace element values remain comparable with those of other samples taken nearby, with the exception of Ce (143 ppm) and Nd (80 ppm), which are enriched roughly six and seven times respectively compared with the average values for the Orkney Formation (Ce - 25 ppm, Nd - 12 ppm). The La value (12 ppm) is, however, normal (formation average - 11 ppm).

Wendlandt and Harrison (1979) showed that in the presence of a CO_2 -rich vapour, light REE are fractionated preferentially into the vapour relative to silicate melts, with maximum REE solubility in CO_2 vapour occurring at near-surface pressures. This could explain the enrichment of Ce and Nd in this sample, but not the low La value, which should also have shown enrichment, unless the original La concentration was about 2 ppm, in which case equivalent enrichment would have increased the value to the present 12 ppm. Discriminant analysis grouped this sample with either the Dominion Group basic lavas ($P(G/X) = .933$) or with the Allanridge lavas ($P(G/X) = .062$). As this sample occurs in the middle of the Orkney Formation, this misclassification is clearly due to the unusual chemistry displayed by this sample and on this basis it was excluded from the data file.

PG-177 was taken 4.3m from the top of the uppermost flow of the Goedgenoeg Formation in borehole WS-5. The 7.8m flow is overlain by quartz-feldspar porphyries of the Makwassie Formation. During logging the unit was described as "atypical of usual top of flow" and "possibly intrusive". Microscope examination, however, showed that the rock is clearly a vitric tuff. Analysis of major elements (Table 2.4) showed an anomalously high SiO_2 content of 78.24 percent, very low TiO_2 , Al_2O_3 , CaO and Na_2O , extremely low P_2O_5 , and normal Fe_2O_3 , MgO and MnO. Most of the trace element values differ widely from those in the other Goedgenoeg rocks. On various interelement plots this sample lies extremely far away from all other plotted samples. Discriminant analysis did not group it into any of the nine formations under consideration, so this sample was also omitted from the data file.

4.3.5 Conclusions

The final outcome of the discriminant analysis exercise was thus a correct classification of 99.07% of all samples used (see Table 4.5). It is an essential prerequisite of discriminant analysis that, in order to apply the technique to a set of data, a preliminary classification of some or all of the data into groups must exist. Naturally, the greater the number of pre-classified cases, the better defined the groups will be. The outcome of any subsequent work may result in the need for reclassification of samples already on the file. In this way the file may be periodically updated and the discriminant functions will thereby become progressively refined. Users are cautioned, however, against reclassifying samples blindly on the basis of discriminant analysis recommendations, without fully investigating the reasons for this.

4.4 Orthogonal Discrimination

Discriminant analysis, as described in section 4.2, is a highly sophisticated method of discrimination, but it has a number of disadvantages. Firstly, the researcher must have access to a computer with adequate facilities, viz. a discriminant analysis package as part of the software. Secondly, assuming that the first condition is satisfied, graphical representation of the discrimination is awkward to plot, as only the scores on the two most important discriminant functions are plotted by the program. In the current exercise, there are nine chemically distinct sets of rocks giving eight discriminant functions. In order to plot each against the other, so that the situation might be examined in all its dimensions, it would require 36 graphs to be plotted. Thirdly, if the user wishes to find out which elements are the most important in discriminating between two particular formations, the program will have to be run using data from those two formations only. This is due to the fact that the discriminant functions are set up with the objective of separating as much as possible all formations under consideration simultaneously. To isolate the most favourable discriminating elements between each and every formation in this study, would therefore require 36 separate runs.

As an alternative, a second, perhaps less sophisticated but systematic approach has been adopted, which may prove to have greater and more rapid practical application than discriminant analysis.

All major element and trace element analyses were stored on a computer file, using the format shown in Figure 4.4. Each sample is represented by two lines of data, the first of which contains the sample name (6 columns), a number representing the formation, if known (2 columns) and the major element analysis. Line two contains trace element values.

Figure 4.4 Format for computer data file in which sample name, formation number, major and trace element analyses are stored

```

DB- 110 57.76 1.084 14.58 11.35 0.29 3.51 7.35 3.33 0.50 0.255 0.00
      204. 8.5173.30.1 286. 11.9125. 28. 19. 55. 5.297. 27. 48. 28.
DB- 210 58.44 1.041 13.85 11.48 0.15 3.54 6.82 3.03 1.37 0.265 0.00
      395. 7.5162.26.9 460. 51.0100. 16. 23. 54. 6.261. 25. 59. 24.
DB- 310 58.78 1.004 13.66 11.72 0.14 3.17 6.79 3.17 1.34 0.237 0.00
      484. 7.1171.27.7 465. 44.5103. 42. 25. 53. 12.245. 27. 50. 28.
DP- 4 9 68.32 0.849 12.74 7.35 0.11 2.50 2.05 3.65 2.16 0.271 0.00
      846.12.9256.28.9 214. 56.9 91. 10. 12. 26. 12.106. 44. 77. 37.
DP- 5 9 70.84 0.809 12.50 5.66 0.08 1.73 1.45 3.42 3.22 0.278 0.00
      853.13.4277.27.3 111. 81.5 87. 9. 9. 16. 2. 53. 46. 94. 40.
DP- 6 9 69.87 0.830 12.65 6.71 0.07 1.70 2.20 2.70 3.03 0.243 0.00
      889.15.0275.27.1 71.114.7 82. 7. 12. 21. 4. 87. 35. 76. 34.
DB- 710 52.21 0.883 13.29 12.87 0.18 7.71 10.92 1.67 0.08 0.191 0.00
      25. 4. 2 91.22.7 386. 1.3107. 58.310. 68.1344.236. 10. 29. 17.
DB- 810 52.19 0.889 13.08 13.68 0.23 10.07 6.68 2.85 0.11 0.194 0.00
      48. 2.4 89.25.0 132. 2.2126. 60.446. 87.1510.257. 13. 27. 19.
DB- 910 53.77 0.843 13.09 11.48 0.17 9.52 7.36 3.43 0.15 0.180 0.00
      77. 5.3 93.22.7 165. 1.9 99. 43.322. 71.1290.221. 7. 15. 11.
DB- 1010 57.55 1.171 14.46 10.05 0.15 4.94 5.57 5.11 0.57 0.421 0.00
      244. 7.7174.30.2 147. 9.8 95. 53.125. 58. 34.230. 21. 45. 25.
    
```

A computer program, ELEMENT (see Figure 4.5), was developed which reads in the above data, sorts the samples into their respective formations, and calculates the largest and smallest value for each major and trace element within each formation. These ranges are output as shown in Tables 3.2a and 3.2b. The program then compares each formation with every other formation and prints a summary of the elements whose ranges do not overlap for each pair. Table 4.6 has been constructed to show which pairs of formations can be separated from one another on the basis of discrete ranges of certain elements. Major element separators are located in the upper half of the table above the main diagonal, while trace element separators are listed below the diagonal.

Use of the table is best demonstrated by working through an actual example. One may wish, for instance, to determine which elements may be used to distinguish between the Allanridge and Loraine Formations. Following row 1 (Allanridge) and column 4 (Lorraine), a block is reached

Figure 4.5 Computer program ELEMENT (in FORTRAN)

```
001      LIST
002      PROGRAM(ELEM)
003      INPUT 5 = CR0
004      OUTPUT 6 = LPO
005      COMPRESS INTEGER AND LOGICAL
006      TRACE 2
007      END
008
009      MASTER ELEMENTS
010
011 C-----
012 C PROGRAM TO GIVE RANGES OF MAJOR OR TRACE ELEMENTS FOR EACH FORMATION !
013 C AND TO DISCRIMINATE BETWEEN FORMATIONS USING MAJOR OR TRACE ELEMENTS !
014 C-----
015
016      INTEGER EL,DISCRIM
017      DIMENSION ELEM(330,15), HEAD(15)
018      DIMENSION FORM(10), SMALL(10,15), BIG(10,15)
019      DIMENSION INDEX(10,100)
020      DIMENSION DIFF(10,10,15), DISC(10,10,15)
021
022 C READ PROGRAM PARAMETERS
023 C *****
024      READ (5,1010) EL,DISCRIM
025
026 C INITIALISE ARRAY 'INDEX'
027 C *****
028      DO 2 J = 1,10
029      DO 2 K = 1,100
030 2      INDEX(J,K) = 0
031
032 C READ ELEMENT HEADING & NAMES OF FORMATIONS
033 C *****
034      IF (EL.EQ.10) READ (5,10U2) (HEAD(J), J=1,EL)
035      IF (EL.EQ.15) READ (5,10L2) (HEAD(J), J=1,EL)
036      READ (5,10O3) (FORM(J), J=1,10)
037
038 C READ ELEMENT DATA & SET UP INDEX OF FMNS
039 C *****
040      READ (5,1000) ICOUNT
041      DO 5 I = 1,ICOUNT
042      IF(EL.EQ.10) READ (5,1001) SANO, IFMN, (ELEM(I,J), J=1,EL)
043      IF(EL.EQ.15) READ (5,1011) SANO, IFMN, (ELEM(I,J), J=1,EL)
044      IF (IFMN.EQ.0) GO TO 5
045      INDEX(IFMN,1) = INDEX(IFMN,1) + 1
046      J = INDEX(IFMN,1)
047      INDEX(IFMN,J+1) = I
048 5      CONTINUE
049
050 C INITIALISE ARRAYS 'BIG' & 'SMALL'
051 C *****
052      DO 1 I=1,10
053      DO 1 J=1,EL
054      DO 1 K=1,EL
055      SMALL(I,J) = 9999.999
056 1      BIG(I,J) = 0.0
057
058 C FIND LARGEST & SMALLEST ELEMENT VALUES FOR EACH FORMATION
059 C *****
060      DO 100 J = 1,EL
061      DO 100 L = 1,10
062      DO 100 M = 2,INDEX(L,1)+1
063      I = INDEX(L,M)
064      IF (I.EQ.0) GO TO 100
065      IF(ELEM(I,J).LT.SMALL(L,J)) SMALL(L,J) = ELEM(I,J)
066      IF(ELEM(I,J).GT.BIG(L,J)) BIG(L,J) = ELEM(I,J)
067 100      CONTINUE
068
069      IF (DISCRIM.EQ.0) CALL RANGES(EL,HEAD,FORM,BIG,SMALL)
070      IF(DISCRIM.EQ.1)CALL DISCRIMINATE(EL,HEAD,FORM,BIG,SMALL,DIFF,
071      *DISC)
072
073 10U0      FORMAT (I3)
074 10U1      FORMAT (A6,I2,F6.2,F6.3,/F6.2,F6.3,/)
075 1011      FORMAT (A6,I2,/,7X,F5.0,F4.1,F4.0,
076      * F4.1,F5.0,F5.1,4F4.0,F5.0,4F4.0)
077 10U2      FORMAT (10A5,/)
078 10L2      FORMAT (/15A2)
079 10U3      FORMAT (10A4)
080 1010      FORMAT (I2,I1)
081
082      STOP
083      END
084
```

Figure 4.5 continued

```
085      SUBROUTINE RANGES(EL,HEAD,FORM,BIG,SMALL)
086 C      #####
087
088      INTEGER EL
089      DIMENSION HEAD(EL), FORM(10)
090      DIMENSION SMALL(10,EL), BIG(10,EL)
091
092 101   DO 110 J=1,EL
093      IF(EL.EQ.10) WRITE(6,5003) HEAD(J)
094      IF(EL.EQ.15) WRITE (6,5000) HEAD(J)
095      DO 105 L=1,10
096      IF(BIG(L,J).EQ.0.AND.SMALL(L,J).EQ.9999.999) GO TO 105
097      IF(EL.EQ.10) WRITE (6,5001) FORM(L),SMALL(L,J),BIG(L,J)
098      IF(EL.FO.15) WRITE (6,5002) FORM(L),SMALL(L,J),BIG(L,J)
099 105   CONTINUE
100      WRITE (6,5012)
101 110   CONTINUE
102
103 5000  FORMAT (1X,A2)
104 5001  FORMAT (1X,A4,F7.3,F7.3)
105 5002  FORMAT (1X,A4,F6.1,F7.1)
106 5003  FORMAT (1X,A5)
107 5012  FORMAT (1H0)
108
109      RETURN
110     END
111
112
113      SUBROUTINE DISCRIMINATE(EL,HEAD,FORM,BIG,SMALL,DIFF,DISC)
114 C      #####
115
116      DIMENSION DIFF(10,10,EL),DISC(10,10,EL)
117      DIMENSION KEY(100), INFO(100)
118      DIMENSION BIG(10,EL), SMALL(10,EL), FORM(10), HEAD(EL)
119      REAL INFO,KEY
120      INTEGER EL,BMIN1
121
122 C INITIALISE ARRAYS 'DIFF' & 'DISC'
123 C *****
124 400   DO 3 L=1,10
125      DO 3 M=1,10
126      DO 3 J=1,EL
127      DIFF(L,M,J) = -1.0
128 3     DISC(L,M,J) = -1.0
129
130 C CALCULATE DIFF. BETWEEN LIMITS & CALCULATE DISCRIMINATION VALUE
131 C *****
132      DO 203 J=1,EL
133      DO 201 L=1,10
134      DO 200 M=1,10
135      IF(L.EQ.M) GO TO 200
136      IF(SMALL(M,J).LT.BIG(L,J)) GO TO 200
137      IF(SMALL(M,J).EQ.9999.999.AND.BIG(M,J).EQ.0) GO TO 200
138      DIFF(L,M,J) = SMALL(M,J) - BIG(L,J)
139      IF(SMALL(M,J).EQ.0.0) GO TO 200
140      DISC(L,M,J) = DIFF(L,M,J) / SMALL(M,J)
141 200   CONTINUE
142 201   CONTINUE
143 203   CONTINUE
144
145 C INITIALISE ARRAYS 'INFO' & 'KEY'
146 C *****
147      DO 300 I = 1,100
148      KEY(I) = -1.0
149 300   INFO(I) = X
150
151 C SORT & PRINT DISCRIMINATION VALUES
152 C *****
153      DO 700 L=1,10
154      DO 700 M=1,10
155      IF(L.EQ.M) GO TO 700
156      I=0
157      DO 600 J=1,EL
158      IF(DIFF(L,M,J).LT.0.0) GO TO 600
159      I=I+1
160      INFO(I)=HEAD(J)
161      KEY(I) = DISC(L,M,J)
162 600   CONTINUE
163      B=I
164 610   N = U
165      BMIN1 = B-1
166      IF(BMIN1.LT.1) GO TO 625
```

Figure 4.5 continued

```

167      DO 620 J=1,BMIN1
168      IF(KEY(J).GE.KEY(J+1)) GO TO 620
169      TEMP = KEY(J)
170      KEY(J) = KEY(J+1)
171      KEY(J+1) = TEMP
172      TEMP = INFO(J)
173      INFO(J) = INFO(J+1)
174      INFO(J+1) = TEMP
175      N=J
176 620   CONTINUE
177 625   IF(N.EQ.0) GO TO 630
178      B=N
179      GO TO 610
180 630   WRITE(6,4000) FORM(L), FORM(M)
181      IF(I.EQ.0) GO TO 700
182      DO 640 K=1,I
183      IF(EL.EQ.10) WRITE (6,4002) K, INFO(K), KEY(K)
184 640   IF(EL.EQ.15) WRITE (6,4001) K, INFO(K), KEY(K)
185 700   CONTINUE
186
187 4000  FORMAT(1X,/,1X,A4,1H:,A4,/)
188 4001  FORMAT (1X,I2,1X,A2,1X,F4.3)
189 4002  FORMAT (1X,I2,1X,A5,1X,F4.3)
190      RETURN
191      END
192
193
194
195      FINISH

```

which contains the major element separators for those two formations, P_2O_5 and TiO_2 . A similar exercise using row 4 (Lorraine) and column 1 (Allanridge) isolates the trace element separators, Zr, Ce and Nd. The figure next to each element name indicates on a scale between 0 and 1 the relative magnitude of the separation, and is calculated as follows: Looking at Zr values, for example, for the two formations under consideration, Table 3.2b shows that Zr values for the Lorraine Formation lie between 28 and 113 ppm, while those for the Allanridge Formation range between 141 and 230 ppm. The two sets are therefore separated by an amount of 28 ppm. A separation of 28 ppm at about the 100 ppm level, as in this case, is obviously of much greater significance than a similar separation at about the 1000 ppm level. In order to reflect this discrepancy, the number is divided by the lower limit of the higher range, i.e. $28/141 = .199$. This number will be referred to as the orthogonal discriminator or OD, and has limiting values of 0 and 1. The OD attains its lower limit when the two ranges under consideration are immediately adjacent to one another, as shown by Nd for the Lorraine Formation (0 to 17 ppm) and the Allanridge Formation (17 to 31 ppm)(see Table 3.2b). The OD will then be equal to $(17 - 17)/17$, which is zero. The upper limit of 1 will only be reached if a particular element is entirely absent from one of the formations, in which case the OD will be equal to $(x - \hat{o})/x$, giving 1. In practice this limit is only approached, the highest OD in Table 4.6 being .859 for Cr between the

Jeppestown amygdaloid and the Alberton Formation. Their ranges are, respectively, 3 to 9 ppm and 64 to 415 ppm, resulting in an OD of $(64 - 9)/64 = .859$.

The computer program ELEMENTS was expanded to extend the approach described above to include major:major element ratios, trace:trace element ratios and major:trace element ratios. The results are listed in Tables 4.7a, 4.7b and 4.7c respectively, and are read in the same manner as Table 4.6. The ranges from which these three tables were derived are shown in Tables 4.8a, 4.8b and 4.8c respectively. For Table 4.7c and 4.8c the major element oxide values have been converted to their respective cation values.

The OD for a particular ratio is the same, no matter which way around the ratio is expressed. For example, the OD for the ratio $P_2O_5:TiO_2$ between the Allanridge and Rietgat Formations is $(0.394-0.237)/0.394 = .398$, while that for the inverse ratio, $TiO_2:P_2O_5$, is $(4.213-2.537)/4.213 = .398$. Because of this, in Tables 4.7a, 4.7b and 4.7c, ratios have been expressed in only one way, but it is implied that the inverses for those ratios are also valid.

Five of the trace elements have values of zero in certain samples: Nb, in the Loraine Formation; Cr, in the Allanridge Formation, Dominion Group porphyries and basic lavas; La, Ce and Nd in the Loraine Formation; and Nd in the Orkney Formation. These cannot be used in the denominator of a ratio, as the resulting ratio would be infinity, so these particular ratios are omitted when calculating the ranges for the ratios. This only affects the upper limit of the range, which takes the value of the next highest ratio.

An examination of Tables 4.6, 4.7a, 4.7b and 4.7c reveals that, of the 36 possible pairs of formations, only 4 pairs are not able to be separated by orthogonal discrimination. All four pairs involve the Dominion Group basic rocks:

- Dominion Group basic rocks - Allanridge
- Dominion Group basic rocks - Loraine
- Dominion Group basic rocks - Orkney
- Dominion Group basic rocks - Alberton

Table 4.6 Separation of formations using major and trace elements (see Table 3.1a for abbreviations)

	ALLA	RIET	MAKW	LORA	ORKN	ALBE	JEPP	DR-P	DR-B
ALLA		P205.458	P205.353	P205.201 TI02.015	P205.091		P205.431 MGO .203		
RIET	ND .392 CE .378 LA .295 Y .189 NB .083 ZR .057			P205.748 TI02.267	P205.713 TI02.084	P205.648 TI02.020	MGO .315	P205.006	
MAKW	ND .492 ZR .491 Y .483 NB .483 CE .455 LA .311 CO .298			P205.700 TI02.175	P205.658	P205.580			
LORA	ZR .199 CE .135 ND .000	CE .673 ND .667 LA .659 NB .541 ZR .537 Y .409	ZR .750 NB .742 LA .721 CE .714 LA .667 NI .648 Y .623 CO .400		TI02.001	TI02.098 P205.044	P205.736 MGO .615 TI02.120		
ORKN	ZR .220 CE .162 ND .000	CE .684 ND .667 LA .591 ZR .549 NB .489 Y .356	ZR .757 CE .723 ND .721 NB .712 NI .653 LA .600 Y .589 CO .298 V .165				P205.700 MGO .549 CAO .348		
ALBE	ZR .028	ND .627 CE .612 LA .568 ZR .439 NB .391 Y .384	ZR .697 ND .689 CE .661 NB .657 NI .624 Y .607 LA .578 CU .525 CR .391 CO .353 V .065				P205.631 MGO .611 CAO .347		
JEPP	V .816 NI .333 Y .314 CE .247 ND .244 LA .225 CO .213 NB .141 ZN .058 ZR .013	V .821 CR .471 NI .280	V .697 CR .182 ZR .153 ND .066 NB .008	NI .833 V .824 LA .625 CE .605 ND .585 CR .571 NB .570 ZR .515 Y .500 CO .327 ZN .239	V .856 NI .838 CE .617 ND .585 LA .550 ZR .528 NB .521 Y .455 CR .438 CO .213 ZN .187	CR .859 V .839 NI .822 ND .537 CE .531 LA .525 CU .492 Y .479 NB .430 ZR .412 ZN .335 CO .275			
DR-P	CO .404 NI .333 V .279	V .298 CR .294 NI .280	Y .023	NI .833 CO .491 ZR .486 NB .460 CR .429 CE .373 V .307 LA .286 ND .227 Y .000	NI .838 ZR .500 V .436 CO .404 NB .398 CE .392 CR .250 ND .227 LA .143	NI .822 CR .813 CO .451 ZR .377 V .369 NB .283 CE .255 ND .136 LA .095			
DR-B		CE .102 LA .023	NB .301 ZR .223 CE .214 CO .175 ND .148 Y .124 LA .044 V .025				V .832 CO .075	V .342 CO .300	

TABLE 4.7a Separation of formations using major element ratios

ALLANRIDGE							
RIETGAT	P205 :AL203 .443 P205 :TI02 .398 P205 :SI02 .383 P205 :MNO .155 P205 :FE203 .058	RIETGAT					
MAKWASSIE	P205 :TI02 .403 P205 :AL203 .396 P205 :MNO .267 P205 :SI02 .191		MAKWASSIE				
LORAINÉ	MGO :P205 .227 SI02 :P205 .178 AL203:P205 .166 FE203:TI02 .050 MNO :P205 .042 AL203:TI02 .010	AL203:P205 .743 FE203:P205 .731 SI02 :P205 .725 MNO :P205 .695 MGO :P205 .630 TI02 :P205 .482 FE203:TI02 .343 AL203:TI02 .257 SI02 :TI02 .181 MNO :TI02 .057	MNO :P205 .735 MGO :P205 .722 AL203:P205 .721 FE203:P205 .684 SI02 :P205 .639 TI02 :P205 .486 AL203:TI02 .228 FE203:TI02 .191 MNO :TI02 .165 MGO :TI02 .076 CAO :P205 .047		LORAINÉ		
ORKNEY	MGO :P205 .126 AL203:P205 .073 SI02 :P205 .001	FE203:P205 .747 AL203:P205 .715 SI02 :P205 .666 MNO :P205 .638 TI02 :P205 .617 CAO :P205 .585 MGO :P205 .581 NA20 :P205 .555 FE203:TI02 .225 AL203:TI02 .109	FE203:P205 .703 AL203:P205 .690 MGO :P205 .686 MNO :P205 .686 TI02 :P205 .621 CAO :P205 .607 SI02 :P205 .561 NA20 :P205 .294 AL203:TI02 .074 FE203:TI02 .046		ORKNEY		
ALBERTON		AL203:P205 .657 FE203:P205 .641 SI02 :P205 .605 TI02 :P205 .568 MGO :P205 .487 MNO :P205 .486 CAO :P205 .456 FE203:TI02 .100 AL203:TI02 .080	AL203:P205 .628 MGO :P205 .615 FE203:P205 .579 TI02 :P205 .572 MNO :P205 .554 CAO :P205 .484 SI02 :P205 .482 AL203:TI02 .044	TI02 :FE203 .138 TI02 :SI02 .085 P205 :SI02 .022 TI02 :AL203 .020		ALBERTON	
JEPPESTOWN	P205 :MGO .760 P205 :TI02 .507 P205 :CAO .488 P205 :AL203 .482 FE203:MGO .436 P205 :SI02 .317 TI02 :MGO .315 AL203:MGO .298 NA20 :MGO .247 MNO :MGO .200 P205 :MNO .184 SI02 :MGO .109	FE203:MGO .644 MNO :MGO .615 AL203:MGO .441 P205 :MGO .406 NA20 :MGO .386 TI02 :MGO .364 SI02 :MGO .135 FE203:TI02 .106	MNO :MGO .405 FE203:MGO .165 P205 :TI02 .040 MNO :TI02 .014	P205 :MGO .940 MNO :MGO .784 TI02 :MGO .780 P205 :AL203 .761 FE203:MGO .725 P205 :MNO .705 P205 :SI02 .695 AL203:MGO .685 P205 :FE203 .684 NA20 :MGO .674 P205 :CAO .645 P205 :TI02 .577 SI02 :MGO .574 TI02 :AL203 .206 TI02 :FE203 .110 CAO :MGO .082	P205 :MGO .932 P205 :CAO .854 MNO :MGO .769 P205 :AL203 .735 TI02 :MGO .729 P205 :FE203 .703 FE203:MGO .702 P205 :TI02 .687 P205 :MNO .650 P205 :SI02 .629 AL203:MGO .597 P205 :NA20 .511 SI02 :MGO .471 NA20 :MGO .381 TI02 :CAO .361 FE203:CAO .271 SI02 :CAO .261 MNO :CAO .244 AL203:CAO .194 TI02 :AL203 .047	P205 :MGO .917 P205 :CAO .808 MNO :MGO .764 FE203:MGO .732 TI02 :MGO .719 P205 :AL203 .681 AL203:MGO .656 P205 :TI02 .647 NA20 :MGO .585 P205 :FE203 .578 P205 :SI02 .562 SI02 :MGO .561 P205 :MNO .503 MNO :CAO .409 FE203:CAO .342 TI02 :CAO .313 SI02 :CAO .282 AL203:CAO .222 NA20 :CAO .191 CAO :MGO .043 TI02 :AL203 .016	JEPPESTOWN
DR PORPHYRIES		AL203:P205 .332 AL203:TI02 .135 TI02 :P205 .095	AL203:P205 .276 TI02 :P205 .102 AL203:TI02 .101	P205 :MNO .328 P205 :AL203 .065 P205 :MGO .034 P205 :FE203 .033	P205 :CAO .234 P205 :MNO .203 P205 :FE203 .091	AL203:P205 .379 TI02 :P205 .259 K20 :P205 .214 AL203:TI02 .075 AL203:FE203 .003	DR PORPHYRIES
DR BASICS		TI02 :P205 .040	TI02 :P205 .048			MGO :MNO .329 MGO :AL203 .296 MGO :FE203 .256 MGO :NA20 .251 TI02 :P205 .215 MGO :P205 .084	

Table 4.7b Separation of formations using trace element ratios

Formation	Allanridge			Rietgat			Makwassie			Lorraine			Orkney			Alberton			Crown			Dominion Gp porphyries			
	COIND	COICE	V IND	CO17R	CO17R	V IND	CO17R	CO17R	V IND	CO17R	CO17R	V IND	CO17R	CO17R	V IND	CO17R	CO17R	V IND	CO17R	CO17R	V IND	CO17R	CO17R	V IND	
Rietgat	.309	.196	.15A	.255	.147	.016	.246	.241																	
Makwassie	.777	.644	.638	.724	.631	.600	.704	.569	.564	.672	.455	.525	.671	.455	.337	.077									
Lorraine	.298	.214	.169	.067	.055	.053	.047	.041	.039	.029															
Orkney	.250	.204	.105	.091	.081	.042	.009																		
Alberton	.784	.404	.366	.729	.356	.317	.667	.274	.254	.192	.179	.155	.539	.143	.120	.120	.049	.041	.037	.036	.007				
Crown	.936	.571	.506	.920	.505	.499	.917	.477	.471	.451	.461	.389	.357	.228	.204	.204	.204	.204	.204	.204	.204	.204	.204	.204	.204
Dominion Gp porphyries	.608	.415	.399	.581	.379	.374	.542	.372	.318	.257	.165	.145	.091	.041	.029										
	.253	.213	.198	.151	.046																				
	.769	.517	.494	.770	.458	.427	.669	.191	.081	.008	.576														
	.969	.467	.313	.966	.262	.253	.961	.242	.215	.192	.190	.180	.146	.118	.091	.044	.016	.013	.006	.006	.006	.006	.006	.006	.006
	.959	.694	.688	.957	.631	.608	.953	.499	.364	.220	.120	.118	.095	.082	.073	.060	.053	.039	.039	.039	.039	.039	.039	.039	.039
	.933	.528	.492	.933	.407	.390	.906	.358	.305	.262	.250	.160	.115	.114	.095	.059									
	.937	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319	.936	.255	.249	.919	.203	.203	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120	.120
	.936	.307	.319																						

Table 4.7c continued

	ALLANRIDGE FORMATION	RIETGAT FORMATION	MAKWASSIE FORMATION					
ALBERTON FORMATION	AL17R .124 TI17R .108 FE17R .012 MG1CE .010	P1CO .702 P1CU .702 AL1ND .647 SI1CE .623 FE1ND .615 SI1ND .614 PIV .607 FE1CE .600 AL1CE .599 SI1LA .561 AL1LA .556 FE1LA .526 P1NI .523 TI1ND .490 AL17R .448 TI1CE .443 AL1NB .428 FE1Y .421 FE17R .393 PI7N .393 SI1NB .391 SI1ZR .390 AL1Y .381	MG1ND .379 TI1LA .369 SI1Y .353 FE1NB .350 MG1CE .312 MN1ND .285 MN1CE .276 CA1CE .261 MG1LA .235 CA1LA .233 CA1ND .229 MN1LA .227 TI1Y .219 TI1ZR .182 TI1CO .162 TI1CU .120 CA1ZR .118 CA1NB .109 TI1NB .101 CA1Y .088 MN1Y .032 TI1V .004	P1NI .912 P1CU .820 P1CR .816 P1CO .800 AL17R .741 PIV .729 TI1NI .725 CA1CE .724 CA1LA .717 AL1NB .715 TI1ZR .715 FE17R .706 AL1ND .705 FE1NB .704 MG1ZR .679 FE1CE .675 MG1CE .675 TI1NB .671 MG1NB .671 MG1ND .669 CA1ND .663 TI1ND .662 SI1ND .661 AL1CE .658	AL1NI .657 MN17R .651 FE1ND .648 MN1NB .642 MN1CE .640 MG1LA .630 SI1CE .624 SI1CU .621 FE1LA .614 SI17R .613 SI1NI .612 TI1CE .605 AL1Y .604 SI1NB .600 CA1ZR .600 AL1LA .590 MN1LA .588 MN1ND .575 TI1Y .549 FE1Y .544 CA1NB .539 FE1NI .539 MG1Y .536 SI1Y .531	TI1CR .527 TI1LA .517 SI1LA .509 MN1Y .492 CA1Y .483 TI1CU .478 TI1CO .447 AL1CR .438 AL1CU .390 SI1CR .389 SI1CO .383 AL1CO .370 TI1V .314 FE1CO .278 FE1ZN .234 AL1V .173 PI7N .163 MG1ZN .144 CA1CU .104 FE1CR .104 PI7R .101 FE1CU .096 SI1V .056 PINB .040		
	CROWN FORMATION	PIV .945 FE1V .869 TI1V .851 AL1V .848 MN1V .843 NA1V .830 P1NI .814 SI1V .779 P1CO .722 MG1Y .668 MG1NB .653 MG1LA .622 MG1CE .614 MG17R .604 MG17N .598 MG1ND .567 P1CR .542 MN1NI .444 MG1V .429 CA1LA .426 NA1NI .408 FE1CO .401 AL1Y .385 CA1CE .366 TI1Y .355 AL1CE .336 CA1V .309 CA1ND .305	TI1NB .305 CA1Y .304 TI1LA .285 AL1ND .282 AL1LA .281 TI17R .280 AL1NI .278 FE1NI .271 TI1CE .260 TI1NI .247 AL1NB .246 SI1NI .239 PI7R .238 TI1CO .236 AL1CO .235 TI1ND .223 AL17N .206 AL17R .180 SI1Y .146 NA1CO .119 SI1CE .117 TI17N .093 SI1ND .089 SI1NB .080 SI1CO .054 FE1CR .039 MN1CO .020	FE1V .888 MN1V .861 AL1V .838 PIV .835 NA1V .823 TI1V .818 SI1V .774 MG17N .705 FE1CR .486 MN1NI .480 FE1NI .475 SI1CR .474 MN1CR .471 CA1V .447 MG1CO .407 NA1CR .386	MG1Y .355 AL1CR .301 P1CR .299 P1NI .230 MG1NB .219 SI1NI .195 AL1NI .145 TI17N .143 MG1V .120 TI1CR .116 NA1NI .067 FE1CO .062 TI1NI .040 FE1ND .036 CA1NI .018	MN1V .816 PIV .783 FE1V .775 TI1V .773 AL1V .759 SI1V .597 NA1V .557 PI7R .486 MN17R .471 MN1NB .469 FE17R .418 TI17R .414 FE1NB .409 PINB .387 FE1CR .385	AL17R .357 TI1NB .303 AL1NB .278 MN1NI .250 MN1CE .240 P1CR .238 CA1V .236 NA1CR .230 FE1NI .215 P1ND .187 FE1CE .183 AL1CR .159 MN1LA .138 MG17N .135 MN1CR .119	FE1ND .119 MN1ND .111 FE1LA .109 TI1CR .106 PIY .100 PICE .089 TI1ND .084 AL1ND .067 NA1LA .065 FE1Y .063 P1NI .061 MN1Y .055 SI1CR .031 AL1CE .016
		DOMINION GROUP PORPHYRIES	SI1CO .496 TI17R .453 P1CO .452 P1NI .444 SI1NI .418 PIV .396 TI1NB .336 SI1V .318 AL1CO .248	AL17R .228 CA1LA .215 K1NI .205 AL1NI .181 AL1V .162 AL1NB .097 TI1Y .082 CA1CE .075 TI1CO .022	SI1CR .420 SI1NI .385 SI1V .303 K1CR .225 AL1V .104 SI1CO .044 AL1NI .030	AL1Y .330 AL1NB .276 AL17R .241 TI1Y .056		
			DOMINION GROUP BASICS		PIV .134 SI1CE .123 TI1ND .091 SI1LA .046 TI1CE .037 FE1LA .035 FE1CE .024 TI1V .004	TI1NB .584 TI17R .550 FE1NB .501 PIV .403 MN1NB .402 TI1ND .398 TI1Y .367 FE17R .359 TI1CE .318 TI1V .314 AL1NB .311 MN17R .238	FE1LA .215 FE1CE .207 AL17R .197 SI1NB .193 FE1Y .191 TI1LA .185 AL1CO .181 AL1V .126 SI1CE .125 SI1CO .118 MN1CE .102	SI1ND .099 MN1LA .093 SI17R .085 P1CO .079 FE1ND .052 CA1CE .049 AL1CE .033 AL1Y .032 FE1CO .030 CA1LA .015 PINB .001

Table 4.7c continued

	LORAINÉ FORMATION			ORKNEY FORMATION			
ALBERTON FORMATION	TIIV	.156		TIIV	.035		
	TI:CO	.051		PIV	.028		
CROWN FORMATION	PIV	.974	PICU .630	PIV	.975	NAIV .766	FEIV .533
	PINI	.968	TI:CR .623	PINI	.964	CA:LA .753	TI:ZR .531
	MG:NB	.940	MN:LA .611	TIIV	.911	CA:NB .751	TI:NB .526
	MG:CE	.925	FE:CE .608	MG:CE	.909	CA:ZR .723	NA:CR .501
	TIIV	.922	MN:CE .605	MN:V	.906	CA:Y .713	MN:CE .500
	MG:LA	.922	NA:CR .601	MG:LA	.903	NA:NI .677	SI:ND .497
	MG:ND	.919	AL:NB .597	MG:ND	.902	AL:CE .662	TI:CR .492
	MG:ZR	.901	FE:LA .586	MG:NB	.898	FE:CE .623	TI:LA .489
	PI:CR	.898	SI:CR .579	FEIV	.896	FE:ND .613	MN:NB .481
	TI:NI	.883	AL:ND .576	MG:ZR	.885	AL:ND .610	TI:ND .473
	MG:Y	.881	TI:CO .572	TI:NI	.883	CA:ZN .601	NA:ND .466
	CA:LA	.880	FE:ND .564	MG:Y	.881	AL:LA .595	SI:LA .449
	CA:CE	.875	FE:NB .559	AL:V	.874	CA:V .594	MN:CO .448
	FE:V	.874	AL:ZR .546	FE:NI	.868	AL:ZR .585	SI:CR .446
	NA:V	.867	FE:CR .545	PI:CO	.860	AL:NB .576	TI:CO .444
	FE:NI	.867	MN:ND .544	MN:NI	.856	PI:CU .576	MN:ND .439
	PI:CO	.865	SI:LA .540	PI:CR	.846	FE:ZR .567	MN:LA .433
	MN:V	.864	MN:CR .538	AL:NI	.835	FE:LA .565	PI:ZN .431
	AL:V	.857	AL:Y .532	SI:V	.832	FE:NB .547	FE:CR .424
	MN:NI	.855	SI:CE .525	MG:ZN	.822	CA:NI .544	SI:NB .418
MG:ZN	.837	AL:CR .524	SI:NI	.820	SI:CE .537	FE:CO .416	
AL:NI	.834	MN:NB .518	CA:CE	.818	TI:CE .537	MG:V .415	
NA:NI	.826	FE:ZR .511	CA:ND	.799	AL:Y .536	TI:Y .415	
SI:NI	.816	FE:Y .508					
SI:V	.809	MN:ZR .506					
CA:ND	.732	CA:V .502					
AL:LA	.638	SI:ND .501					
AL:CE	.630	SI:NB .490					
DOMINION GROUP PORPHYRIES	PINI	.904	FE:LA .602	PINI	.892	MG:CE .614	MN:ZR .386
	SI:NI	.859	SI:CO .602	SI:NI	.863	AL:ZR .609	MN:ND .367
	CA:LA	.836	FE:ND .583	AL:NI	.812	MG:ZR .593	TI:V .355
	CA:CE	.817	AL:ZR .573	TI:NI	.808	FE:LA .582	AL:ND .317
	AL:NI	.811	FE:NB .546	CA:CE	.735	MG:ND .560	CA:Y .313
	TI:NI	.808	SI:CR .535	PIV	.728	TI:NB .546	AL:V .302
	KINI	.793	K:CR .524	PI:CO	.725	MN:CE .533	SI:ZR .291
	PI:CO	.734	CA:ND .519	KINI	.688	FE:NB .533	NA:ND .289
	PIV	.718	AL:NB .517	CA:ZR	.671	MG:NB .512	TI:CO .288
	MG:NB	.714	MN:ZR .508	CA:LA	.662	SI:CO .509	FE:Y .282
	MG:LA	.693	MN:NB .486	FE:CE	.657	MN:LA .497	TI:ND .257
	MG:CE	.684	MN:ND .486	FE:ZR	.650	AL:NB .492	NA:CE .243
	MN:LA	.655	TI:CO .452	TI:ZR	.643	SI:V .480	MN:NI .231
	MG:ZR	.649	TI:V .436	CA:NB	.642	PI:CR .449	NA:LA .221
	FE:CE	.644	SI:V .409	CA:ND	.641	MN:NB .446	AL:CE .221
	PI:CR	.636	FE:NI .407	FE:ND	.630	FE:NI .412	MG:Y .206
	MG:ND	.635	TI:ZR .380	MG:LA	.620	SI:CR .389	
	MN:CE	.632	TI:NB .311				
	FE:ZR	.605	K:V .032				
	DOMINION GROUP BASICS						

Table 4.7c continued

ALBERTON FORMATION							
CROWN FORMATION	PIV .961	NAICR .824	MNICU .520	MGICO .338			
	PINI .960	NAIV .823	FEIND .519	MNICE .336			
	PICR .950	SIIV .823	TIILA .508	MNICO .333			
	MGICE .893	PICO .815	TIIND .491	FEI7N .331			
	MGIND .891	SIINI .794	FEICO .485	MNILA .319			
	TIIV .890	CAICE .770	TIINR .481	ALICO .314			
	MGILA .888	CAILA .769	TI:7R .480	ALICU .311			
	MNIV .886	NAINI .756	CAINI .480	TI:7N .305			
	FEIV .884	CAIND .751	FEIY .471	PIZN .297			
	FEINI .880	CAIY .720	TI:Y .471	MGIV .277			
	MGIY .875	CAINR .679	FEICU .468	SI:7R .269			
	MG:7R .865	CA:7R .675	SIICE .465	MNIY .232			
	MGINB .864	CA:7N .631	FEILA .460	MNIND .212			
	ALIV .863	NAICU .586	SIIND .452	MGINI .206			
	SIICR .863	ALIND .582	TIICO .437	SIICO .195			
	TIINI .857	ALICE .581	TIICU .429	SI:7N .195			
	MNINI .850	ALILA .579	FEI7R .426	KICR .161			
	TIICR .848	CAIV .545	AL:7N .426	MN:7R .160			
	FEICR .847	FEICE .540	SIILA .421	MGICR .136			
	MNICR .843	SIICU .537	FEINR .420	MNINB .131			
MG:7N .839	ALIY .534	CAICR .417	NAICO .119				
ALICR .830	ALINB .527	SI:Y .376	PIY .089				
PICU .826	AL:7R .525	SIINR .349	MN:7N .022				
ALINI .824	TIICE .523						
CROWN FORMATION							
DOMINION GROUP PORPHYRIES	PINI .881	FEICE .583	SIIV .452	TIIV .205	PIV .557		
	SIICR .849	PIV .576	ALINB .434	MNINI .198	PI7R .473		
	SIINI .842	SIICO .571	FEINR .402	FEIY .186	PI:NB .265		
	PICR .820	MGILA .561	MNILA .395	MGIY .169	FEIV .252		
	ALINI .801	CAIND .553	MNICE .380	MN:7R .163	TIIV .239		
	TIINI .766	AL:7R .552	MGINR .351	MGICR .128	ALIY .213		
	KINI .760	MGICE .547	CAIY .331	SI:7R .121	TI:7R .198		
	KICR .753	FEIND .541	ALICO .326	ALIY .119	MNIV .169		
	ALICR .746	CAINR .538	MNICR .318	MNIND .112	FEI7R .083		
	TIICR .702	FEI7R .536	FEICR .290	TIICE .107	PI:RR .049		
	CAILA .685	MG:7R .521	TIIND .282	PI7R .078			
	CAICE .664	MGIND .509	TIICO .280	MNINB .073			
	PICO .636	TIINB .504	ALIND .269	ALICE .034			
	CA:7R .613	FEILA .481	TI:Y .246	SIINB .011			
	TI:7R .605	FEINI .465	ALIV .244	KICO .001			
	CROWN FORMATION				DOMINION GROUP PORPHYRIES		
	DOMINION GROUP BASICS	PIV .915	CAICE .207	SIIV .400	FEIV .897	CAILA .197	SIICO .387
		TIIV .890	TI:7R .180	TI:7R .377	MNIV .887	TIICE .177	TIINR .373
		ALIV .855	TIILA .169	TIIV .205	NAIV .851	PICO .148	ALIV .201
		SIIV .806	CAIY .145	ALICO .125	CAIV .530	MGIND .141	ALICO .125
MG:7N .529		CAIND .119	PIV .067	MGINB .423	ALICO .109		
MGILA .373		TIIND .094		TIINB .344	CAINR .074		
MGIY .329		FEIY .061		FEICO .308	NAICO .061		
FEICO .307		FEINR .024		MGICE .307	FEINR .024		
TI:Y .258		PINI .018		TI:Y .258			
MG:7R .249							

TABLE 4.8a Ranges of major element ratios for each formation

S102 IS102		T102 IS102		AL203:IS102		FE203:IS102		MNO IS102		MGO IS102		CAO IS102		NA2O IS102		K2O IS102		P2O5 IS102								
AR	1.000	1.000	AR	.015	.025	AR	.223	.311	AR	.091	.296	AR	.047	.117	AR	.093	.071	AR	.005	.122	AR	.003	.005			
PR	1.000	1.000	PR	.019	.045	PR	.213	.394	PR	.106	.300	PR	.037	.148	PR	.047	.261	PR	.019	.082	PR	.002	.053	PR	.002	.025
PM	1.000	1.000	PM	.016	.031	PM	.195	.338	PM	.096	.219	PM	.011	.087	PM	.029	.120	PM	.009	.053	PM	.001	.025	PM	.001	.013
KL	1.000	1.000	KL	.008	.016	KL	.163	.301	KL	.178	.237	KL	.002	.004	KL	.014	.269	KL	.003	.049	KL	.002	.024	KL	.002	.002
KO	1.000	1.000	KO	.016	.021	KO	.244	.305	KO	.188	.276	KO	.002	.003	KO	.008	.200	KO	.039	.130	KO	.004	.024	KO	.002	.003
KA	1.000	1.000	KA	.017	.022	KA	.243	.316	KA	.174	.278	KA	.002	.003	KA	.009	.177	KA	.012	.083	KA	.007	.051	KA	.002	.003
JC	1.000	1.000	JC	.015	.022	JC	.179	.390	JC	.172	.374	JC	.002	.005	JC	.030	.065	JC	.001	.084	JC	.002	.022	JC	.002	.003
DP	1.000	1.000	DP	.007	.025	DP	.138	.370	DP	.024	.211	DP	.000	.003	DP	.005	.095	DP	.003	.078	DP	.014	.022	DP	.002	.009
DB	1.000	1.000	DB	.011	.036	DB	.187	.417	DB	.138	.343	DB	.002	.005	DB	.003	.219	DB	.005	.089	DB	.002	.043	DB	.002	.014

Table 4.Bb continued

HA:ICU		NB:ICU		ZR:ICU		Y:ICU		SR:ICU		RH:ICU		ZN:ICU		CU:ICU		NI:ICU		CO:ICU		CR:ICU		V:ICU		LA:ICU		CE:ICU		ND:ICU																							
AR	4.708	148.200	AR	0.046	2.440	AR	1.055	38.600	AR	.172	8.160	AR	1.258	66.400	AR	.040	15.020	AR	.547	27.600	AR	1.000	1.000	AR	.585	45.200	AR	.390	16.400	AR	0.000	17.200	AR	1.134	50.400	AR	.115	4.000	AR	.291	8.800	AR	.152	4.000							
PR	4.375	123.917	PR	.296	6.533	PR	5.271	112.000	PR	.824	19.767	PR	7.030	84.333	PR	.103	6.467	PR	1.323	33.333	PR	1.000	1.000	PR	1.282	64.000	PR	.694	17.000	PR	1.359	180.667	PR	2.565	66.333	PR	.403	16.333	PR	2.000	35.000	PR	1.000	14.500							
PM	2.500	185.600	PM	.843	3.800	PM	16.571	56.500	PM	2.336	7.988	PM	9.417	105.107	PM	.200	18.031	PM	2.786	20.400	PM	1.000	1.000	PM	.746	2.533	PM	.679	3.300	PM	1.043	2.500	PM	3.214	13.250	PM	2.964	10.700	PM	6.296	20.500	PM	2.944	14.000							
KL	.372	24.053	KL	0.000	.189	KL	.271	5.105	KL	.045	1.174	KL	.400	21.053	KL	.091	3.505	KL	.683	4.421	KL	1.000	1.000	KL	.923	12.000	KL	.347	3.053	KL	1.058	60.444	KL	1.529	11.032	KL	0.000	.032	KL	0.000	1.243	KL	0.000	.032							
KO	1.728	15.688	KO	.029	.256	KO	.730	6.000	KO	.162	1.194	KO	1.845	45.500	KO	.168	1.381	KO	.628	4.188	KO	1.000	1.000	1.000	KO	1.066	7.875	KO	.447	2.938	KO	.176	1.813	KO	1.670	11.740	KO	.047	.020	KO	.151	1.563	KO	0.000	.750						
KA	1.476	12.516	KA	.029	.102	KA	.857	2.322	KA	.194	.380	KA	1.400	11.910	KA	.093	1.253	KA	.670	1.610	KA	1.000	1.000	1.000	KA	1.312	3.014	KA	.516	1.000	KA	.525	5.900	KA	1.402	3.186	KA	.058	.220	KA	.243	3.077	KA	2.933	6.923	KA	1.533	3.000			
JC	9.385	46.538	JC	.537	1.208	JC	8.033	19.615	JC	1.737	4.177	JC	7.300	25.826	JC	.331	2.523	JC	5.167	12.415	JC	1.000	1.000	1.000	JC	.467	1.154	JC	.967	2.231	JC	.200	.400	JC	.600	1.345	JC	1.433	3.077	JC	2.933	6.923	JC	1.533	3.000						
JP	4.354	337.667	JP	.139	3.800	JP	2.195	77.867	JP	.325	7.800	JP	3.062	55.000	JP	.797	50.033	JP	1.027	18.143	JP	1.000	1.000	1.000	JP	.097	4.333	JP	.186	4.000	JP	0.000	3.000	JP	.177	16.867	JP	.372	14.750	JP	3.735	24.250	JP	.372	14.750	JP	3.735	24.250	JP	.372	14.750
DB	.431	46.644	DB	.032	1.475	DB	1.026	29.250	DB	.222	5.117	DB	2.200	59.265	DB	.022	7.558	DB	.872	27.667	DB	1.000	1.000	1.000	DB	.328	16.694	DB	.625	13.000	DB	0.000	38.583	DB	2.023	67.000	DB	.045	4.833	DB	.213	10.000	DB	.042	4.000	DB	.042	4.000			

Table 4.8c Ranges of major:trace element ratios for each formation

SI:RA		TI:RA		AL:RA		FE:RA		MN:RA		MG:RA		CA:RA		NA:RA		K:RA		P:RA											
AR	267.97	2494.08	AR	5.78	74.78	AR	82.32	859.43	AR	75.67	988.57	AR	1.17	17.78	AR	12.21	359.31	AR	28.70	571.76	AR	1.19	255.46	AR	4.75	97.79	AR	.77	10.30
PR	167.87	3183.39	PR	5.21	107.34	PR	46.04	939.35	PR	74.85	974.63	PR	.37	19.19	PR	11.83	573.94	PR	18.93	481.91	PR	12.65	261.01	PR	5.59	35.36	PR	1.63	41.02
PM	100.31	12384.24	PM	2.96	293.99	PM	29.43	3216.03	PM	20.98	1784.87	PM	.09	27.88	PM	2.70	400.39	PM	6.26	587.20	PM	1.36	1946.75	PM	9.37	38.84	PM	.92	45.14
KL	162.89	15427.12	KL	1.91	149.88	KL	52.71	2841.68	KL	43.34	5319.81	KL	.84	82.29	KL	41.55	6580.24	KL	26.38	3600.70	KL	4.82	248.31	KL	9.54	64.40	KL	.14	10.64
KO	207.00	1944.82	KO	4.56	43.60	KO	66.16	622.34	KO	77.41	643.67	KO	1.08	7.87	KO	27.00	226.13	KO	36.74	595.21	KO	19.71	240.95	KO	11.06	44.49	KO	.78	4.57
KA	225.92	1709.78	KA	5.42	43.09	KA	64.19	532.37	KA	68.35	521.62	KA	.78	5.50	KA	23.43	164.39	KA	38.29	282.32	KA	4.85	192.42	KA	7.38	44.07	KA	.59	4.56
JC	193.99	2559.97	JC	7.26	51.20	JC	81.16	518.02	JC	93.92	725.20	JC	1.48	10.79	JC	6.39	53.87	JC	16.68	169.89	JC	25.75	165.41	JC	10.21	22.93	JC	2.58	18.31
DP	195.43	1076.22	DP	2.73	15.19	DP	38.44	219.55	DP	10.49	224.88	DP	.13	2.18	DP	1.60	68.08	DP	2.96	80.38	DP	1.46	101.62	DP	13.34	64.33	DP	.67	7.57
DB	99.87	9761.18	DB	1.73	211.74	DB	28.39	2813.76	DB	23.48	3600.51	DB	.39	55.76	DB	10.65	1859.65	DB	12.86	3121.81	DB	3.92	495.59	DB	9.85	44.38	DB	.75	33.34

Table 4.8c continued

SI:CU		TI:CU		AL:CU		FE:CU		MN:CU		MG:CU		CA:CU		NA:CU		K:CU		P:CU											
AR	1352.89	56447.22	AR	42.03	1556.30	AR	476.08	17636.28	AR	273.84	17596.90	AR	93.70	6235.02	AR	233.58	8004.64	AR	9.13	5579.09	AR	20.60	5744.98	AR	5.27	242.64			
PR	3827.84	75594.16	PR	175.56	3305.24	PR	1145.68	28211.69	PR	1128.01	21308.39	PR	13.90	619.60	PR	400.45	14451.90	PR	643.23	30207.99	PR	380.22	8358.74	PR	70.06	2166.07	PR	42.65	999.76
PM	11017.29	23378.45	PM	228.67	810.87	PM	2449.90	9004.72	PM	1603.62	9169.13	PM	11.92	116.18	PM	228.28	3147.66	PM	939.32	3510.96	PM	252.25	4866.86	PM	58.11	4136.40	PM	70.60	250.38
KL	1740.51	13657.92	KL	20.85	277.03	KL	573.62	3986.46	KL	573.96	4409.90	KL	7.56	57.07	KL	195.68	2448.46	KL	36.43	2922.75	KL	16.13	1327.61	KL	21.24	742.81	KL	1.21	27.10
KO	1894.51	17200.32	KO	43.37	346.59	KO	539.35	5216.91	KO	585.56	4838.97	KO	7.91	53.25	KO	227.52	1345.44	KO	361.18	2309.37	KO	143.51	1687.82	KO	43.43	380.16	KO	4.40	31.09
KA	2040.47	4180.46	KA	52.97	119.29	KA	656.42	1494.60	KA	690.33	1449.77	KA	8.11	21.00	KA	237.75	627.98	KA	402.27	841.30	KA	58.48	429.72	KA	28.94	277.20	KA	5.57	12.72
JC	9028.01	24024.36	JC	209.03	480.52	JC	2170.13	4861.42	JC	2725.33	6805.70	JC	43.78	101.28	JC	172.86	505.59	JC	636.73	1594.33	JC	1038.66	1552.28	JC	95.79	716.99	JC	73.32	171.87
DP	2799.44	122972.94	DP	40.64	1318.90	DP	600.97	19284.16	DP	505.05	10963.10	DP	7.54	154.90	DP	26.68	3577.80	DP	181.52	4945.72	DP	288.22	5391.14	DP	182.20	14777.56	DP	8.03	302.57
DB	2861.84	38911.05	DB	42.03	1317.90	DB	864.74	15217.38	DB	693.71	19991.18	DB	10.81	271.08	DB	182.78	5557.65	DB	372.67	5669.95	DB	63.28	3363.28	DB	11.45	1044.67	DB	7.10	282.93
SI:NI	1116.38	9929.65	AR	30.03	313.74	AR	322.15	7040.53	AR	353.23	4255.98	AR	4.06	43.03	AR	125.72	625.33	AR	107.21	1378.62	AR	15.62	796.86	AR	13.25	501.40	AR	4.69	27.64
PR	1181.16	10495.93	PR	37.52	399.75	PR	289.19	3601.36	PR	372.94	3066.17	PR	4.56	40.27	PR	162.30	713.95	PR	169.80	1146.38	PR	65.08	1255.29	PR	4.88	743.86	PR	12.40	114.69
PM	6939.00	18363.32	PM	216.61	432.35	PM	2156.20	4729.45	PM	1517.96	4584.57	PM	5.74	58.09	PM	185.72	1573.83	PM	458.91	2576.17	PM	99.57	2862.86	PM	34.18	1632.00	PM	67.06	130.90
KL	478.36	2402.78	KL	4.65	48.74	KL	88.11	701.37	KL	144.96	775.82	KL	2.55	11.25	KL	183.90	396.33	KL	39.48	610.24	KL	2.73	233.56	KL	1.77	130.68	KL	.33	4.77
KO	1499.98	2343.32	KO	38.93	48.77	KO	476.72	696.67	KO	549.46	769.34	KO	6.76	11.16	KO	170.85	272.76	KO	245.04	532.78	KO	110.78	434.10	KO	19.74	194.71	KO	3.27	5.35
KA	1114.94	2690.10	KA	27.75	59.65	KA	322.26	740.50	KA	380.65	699.40	KA	4.95	11.65	KA	160.03	294.16	KA	223.83	607.18	KA	43.22	327.61	KA	13.96	151.29	KA	2.68	5.92
JC	13040.46	27657.33	JC	416.45	559.17	JC	4213.23	6226.49	JC	5839.99	7578.95	JC	77.45	119.70	JC	370.41	539.35	JC	1167.34	2189.58	JC	1345.31	2434.78	JC	83.02	1314.48	JC	148.96	210.68
DP	17071.79	59780.46	DP	254.45	763.08	DP	3713.92	10586.00	DP	1307.34	8770.48	DP	14.52	154.90	DP	174.20	3220.02	DP	230.90	4237.15	DP	114.14	4271.22	DP	630.95	5728.38	DP	49.70	188.38
DB	379.44	20702.48	DB	10.34	612.35	DB	155.23	4820.41	DB	169.90	6144.73	DB	2.58	118.21	DB	80.07	1113.96	DB	36.75	4013.60	DB	29.68	1510.30	DB	2.05	761.62	DB	1.48	146.31
SI:CO	3076.86	6004.60	AR	90.43	162.12	AR	995.49	1666.73	AR	698.03	1689.87	AR	12.00	37.97	AR	214.94	552.54	AR	241.21	2054.41	AR	23.28	613.20	AR	40.17	747.18	AR	10.67	21.08
PR	3449.99	11400.96	PR	142.33	344.61	PR	991.51	3104.62	PR	1253.43	2643.25	PR	15.06	41.70	PR	442.20	964.80	PR	491.51	1776.94	PR	169.35	1082.15	PR	12.86	916.41	PR	47.03	99.87
PM	8282.03	19931.69	PM	215.64	486.23	PM	2370.62	4699.63	PM	2010.78	2834.41	PM	8.61	54.22	PM	241.20	953.84	PM	435.32	2982.93	PM	151.35	2433.43	PM	29.06	2480.64	PM	70.22	153.66
KL	2804.93	4742.44	KL	27.25	90.75	KL	516.67	1533.21	KL	967.24	1482.73	KL	14.52	28.60	KL	434.58	1247.04	KL	94.10	1633.41	KL	16.02	488.42	KL	10.38	336.53	KL	1.93	10.21
KO	2880.63	5855.43	KO	72.94	117.99	KO	972.27	1775.97	KO	1178.16	1647.31	KO	13.90	21.37	KO	352.47	558.90	KO	512.05	1209.32	KO	255.25	905.62	KO	44.80	370.07	KO	6.70	10.59
KA	3560.37	1088.41	KA	95.63	119.31	KA	1138.00	1494.60	KA	1241.72	1452.28	KA	13.11	25.82	KA	396.26	637.08	KA	592.65	1265.44	KA	81.49	612.75	KA	49.29	370.91	KA	9.48	14.01
JC	6344.01	12771.71	JC	212.10	270.59	JC	2179.26	2840.58	JC	2819.31	3287.18	JC	38.73	54.42	JC	178.82	262.39	JC	583.67	898.87	JC	695.85	1187.04	JC	42.94	657.24	JC	75.84	102.50
DP	11920.37	39853.64	DP	165.72	595.10	DP	2217.39	7632.51	DP	1406.57	4846.43	DP	15.49	63.37	DP	71.26	1894.13	DP	171.02	2236.59	DP	148.38	2843.95	DP	371.15	4433.22	DP	38.44	150.41
DB	2073.13	7303.13	DB	56.52	239.56	DB	795.78	1941.21	DB	928.29	1950.40	DB	14.08	40.84	DB	177.55	901.33	DB	200.77	1773.45	DB	50.04	653.64	DB	9.77	612.69	DB	7.67	64.65
SI:CR	1135.80	31357.35	AR	34.06	846.63	AR	391.38	8704.04	AR	391.18	8420.78	AR	5.82	123.92	AR	163.63	1971.81	AR	256.01	5386.24	AR	56.53	3056.63	AR	13.25	3191.66	AR	4.69	110.07
PR	278.64	15435.20	PR	8.82	587.86	PR	67.97	5296.11	PR	89.59	4509.07	PR	1.49	59.23	PR	52.83	1049.93	PR	50.85	1685.85	PR	15.30	1846.02	PR	1.39	1093.91	PR	2.91	168.66
PM	6583.15	28430.67	PM	213.42	594.05	PM	2286.58	6370.85	PM	1496.72	5393.61	PM	5.74	98.57	PM	213.06	1851.56	PM	547.94	2153.63	PM	140.14	2317.55	PM	27.67	2686.83	PM	65.90	183.29
KL	94.28	12357.17	KL	.92	250.65	KL	17.37	3606.80	KL	32.51	3989.91	KL	.50	51.63	KL	32.59	1200.26	KL	22.00	2644.39	KL	.54	1201.17	KL	.35	672.07	KL	.07	24.52
KO	2642.98	16256.76	KO	62.61	337.97	KO	793.95	4826.55	KO	948.88	5088.79	KO	11.18	77.45	KO	338.18	1686.63	KO	443.56	3385.89	KO	271.52	1500.76	KO	57.61	1364.64	KO	5.53	37.09
KA	594.83	4025.48	KA	15.11	100.94	KA	175.50	1285.21	KA	207.29	1340.88	KA	2.63	17.60	KA	95.46	486.17	KA	142.59	907.89	KA	36.65	529.76	KA	7.60	348.43	KA	1.46	12.10
JC	29341.04	104105.58	JC	664.78	2082.26	JC	7574.87	21066.14	JC	8765.81	29491.37	JC	111.87	438.88	JC	562.80	2190.90	JC	1556.46	6908.77	JC	3008.82	6726.56	JC	415.10	2006.32	JC	240.50	744.79
DP	26610.64	313999.32	DP	338.72	5341.55	DP	5050.40	70132.25	DP	1888.38	46300.28	DP	25.82	619.60	DP	557.78	16100.10	DP	333.53	29660.05	DP	164.87	29898.57	DP	1411.34	19703.41	DP	67.28	1274.29
DB	161.55	264361.44	DB	3.53	8470.93	DB	45.85	78283.47	DB	62.24	99734.44	DB	1.02	1394.10	DB	34.59	22130.10	DB	15.90	63608.30	DB	8.93	22331.19	DB	.49	6724.62	DB	.56	1501.22
SI:IV	929.16	1919.84	AR	27.91	51.83	AR	314.91	532.90	AR	193.48	568.54	AR	3.21	10.52	AR	65.19	164.51	AR	67.33	569.43	AR	6.45	190.39	AR	12.71	207.10	AR	2.88	6.74
PR	707.05	1963.08	PR	38.32	63.25	PR	248.75	569.84	PR	239.09	485.15	PR	2.95	9.34	PR	74.28	253.57	PR	77.97	455.40	PR	36.57	198.62	PR	3.23	157.79	PR	12.02	20.14
PM	1635.30	3507.60	PM	55.67	78.95	PM	581.47	845.10	PM	376.01	975.44	PM	1.42	12.36	PM	43.23	334.86	PM	121.76	629.73	PM	31.53	496.62	PM	5.93	516.80	PM	17.46	26.58
KL	948.10	1663.70	KL	11.27	27.07	KL	255.43	501.64	KL	269.98	548.24	KL	4.73	9.11	KL	114.05	591.48	KL	23.82	410.62	KL	4.95	148.78	KL	5.13	103.55	KL	.68	3.14
KO	947.66	1463.86	KO	23.44	30.95	KO	307.04	443.99	KO	361.36	452.93	KO	4.42	6.31	KO	114.51	168.68	KO	156.79	334.15	KO	76.98	262.63	KO	12.38	101.55	KO	1.97	3.04
KA	1144.29	1542.98	KA	32.08	38.17	KA	373.02	481.01	KA	394.63	501.85	KA	4.24	7.65	KA	139.99	208.34	KA	191.65	375.22	KA	26.02	198.27	KA	16.26	123.64	KA	3.13	4.73
JC	8693.64	20154.29	JC	347.04	398.87	JC	3511.02																						

It has already been shown, however, that this problem can be overcome using discriminant analysis, or, if this is unavailable, there are several interelement discrimination plots which can be used. These are discussed in the following section.

4.5 Interelement Discrimination Plots

The relative importance of the various major and trace elements and selected ratios in discriminating between formations was shown quantitatively in Table 4.2 and discussed in section 4.2.1.2. Several interelement plots were constructed using the most important discriminating variables indicated by discriminant analysis and those suggested by orthogonal discrimination. The plots which were most successful in distinguishing between formations are:

Ti:Zr vs Ti:P (Figure 4.6)

Zr:P vs P:Ti (Figure 4.7)

TiO₂ vs Zr (Figure 4.8)

The dotted lines drawn on these plots are arbitrary and serve merely as convenient separators between the units. Figures 4.6 and 4.7 contain the entire sample population and provide very successful discrimination diagrams. The Rietgat and Goedgenoeg Formations plot as a single unit, as suggested by discriminant analysis, and, while the Klipriviersberg Group cannot be separated into its three separate chemical entities, it does, as a whole, lie apart from the other formations.

In order to distinguish between the individual formations of the Klipriviersberg Group a plot of TiO₂ vs Zr (Figure 4.8), is the most effective. Figure 4.8 shows the existence of a small, but consistent gap between which no samples fall. All samples with concentrations below this gap belong to the Loraine-Edenville Formations. Although two Loraine-Edenville samples plot above the Zr gap, one (KL-145) lies completely off the trend demarcated by the other samples of the set, while the other (KL-465) lies amongst the Orkney samples. This is the lowermost sample of the Loraine Formation in borehole JBF-1, and although

the upper purple zone, which marks the top of the Orkney Formation, is present some metres below the position of sample KL-465, this probably implies uncertainty in the positioning of the Orkney-Lorraine boundary, rather than inconsistent chemistry.

All Orkney samples lie within the range of 90 to 110 ppm Zr, while only three Alberton samples have Zr concentrations of less than 110 ppm. These are KA-346, which lies at some distance from the TiO_2 -Zr trend, KA-347 and KA-452, which lie within the Orkney field, near the other Alberton samples.

Figure 4.6 Ti:Zr vs Ti:P discrimination plot for all groups
(See pages 62 & 72 for symbols)

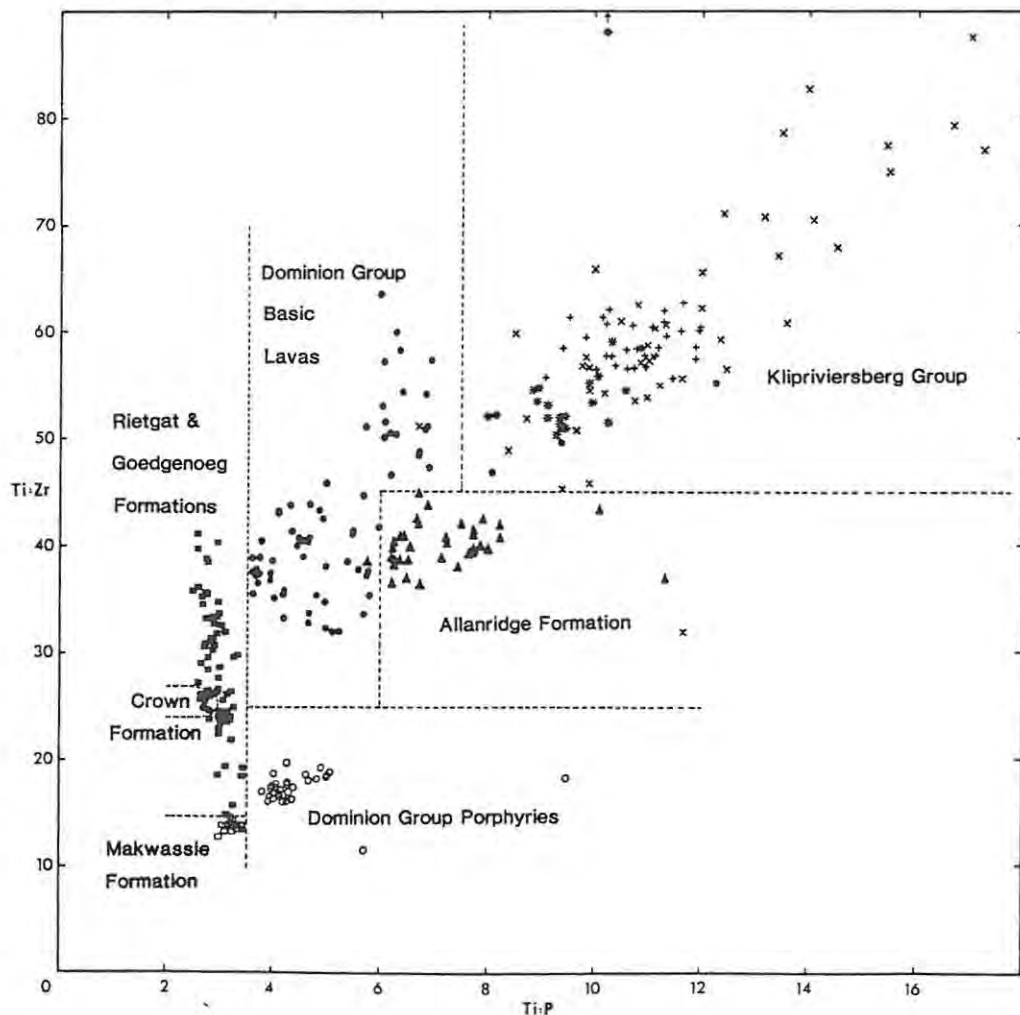


Figure 4.7 Zr:P vs P:Ti discrimination plot for all groups

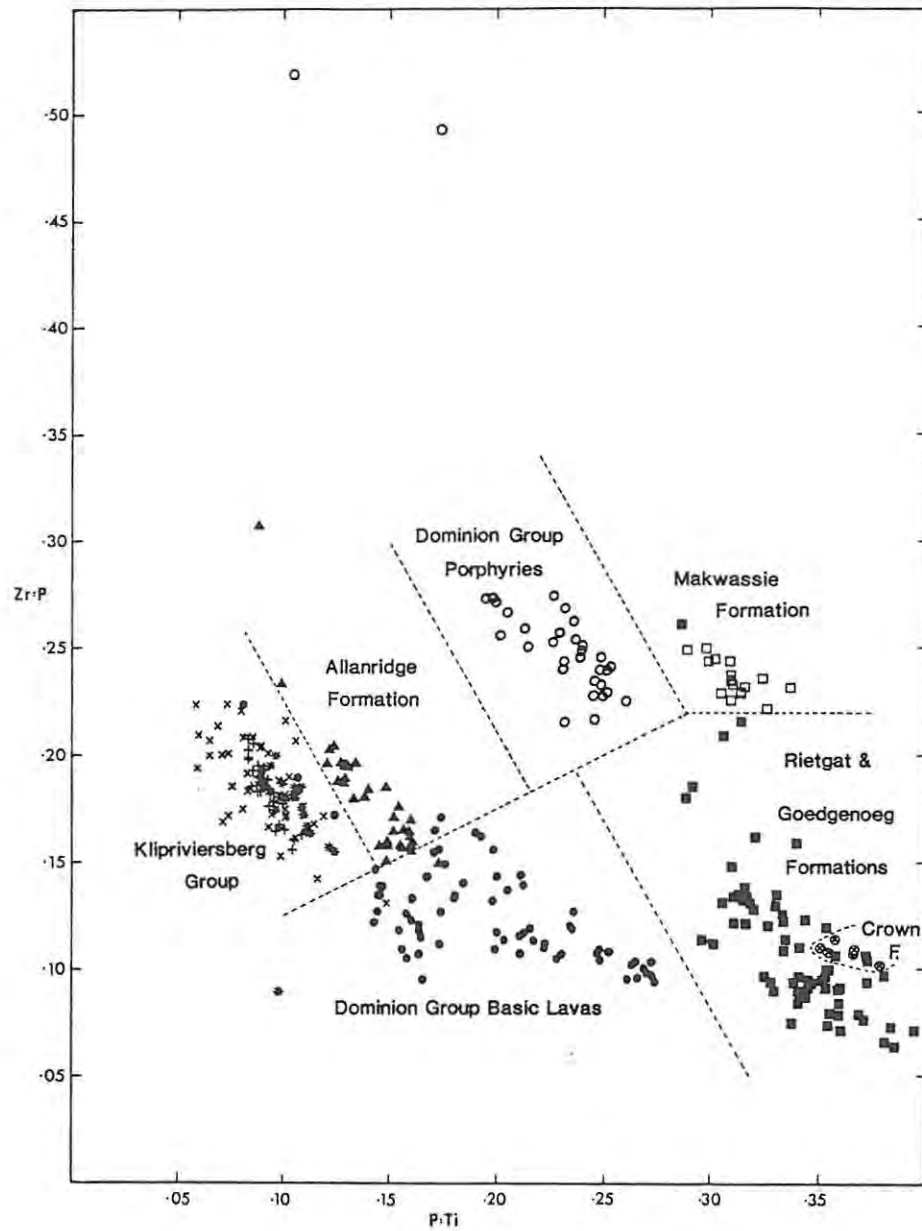
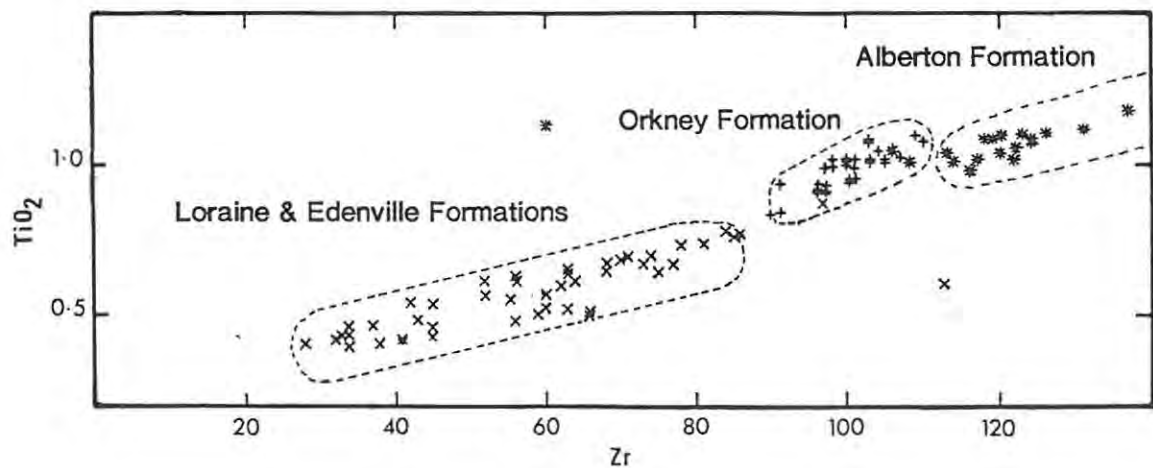


Figure 4.8 TiO₂ vs Zr discrimination plot for the Klipriviersberg Group



The major advantage of these three discrimination plots is that they require the analysis of only three elements, Ti, P and Zr. A large number of samples can thus be analysed within a short space of time in order to determine their stratigraphic affinities with reasonable confidence.

4.6 A Case Study of the Use of Various Discrimination Techniques

During 1981 a prospecting company drilled a borehole to the northwest of Klerksdorp. The volcanic rocks intersected were not able to be identified with any certainty and fourteen samples were sent to Rhodes University for identification. The analytical work and subsequent interpretation were carried out jointly by the author, T.B. Bowen, and her brother, M.P. Bowen.

Table 4.9 shows the list of samples together with the respective depths from surface at which they were taken.

Table 4.9 Unknown samples NL-781 to NL-794 showing depths from surface

NL-781	103 m
NL-782	226 m
NL-783	318.5 m
NL-784	377 m
NL-785	500 m
NL-786	627 m
NL-787	724 m
NL-788	782 m
NL-789	890 m
NL-790	891.5 m
NL-791	1106 m
NL-792	1155 m
NL-793	1230 m
NL-794	1250 m

The above samples were initially analysed for eleven major and five trace elements, viz. Nb, Zr, Y, Sr and Rb, upon the completion of which identification was possible. Analyses of a further ten trace elements were later done so that these samples could be added to the previously compiled data base.

The unique sample NL-789 consists of numerous light ocelli up to 2 cm in diameter set in a darker matrix. These fractions were separated and analysed individually (NL-789-C and NL-789-M). Thus it should be borne in mind that the whole rock analysis in the following diagrams lies somewhere along a mixing line joining the two fractions, which are labelled "C" and "M". The analyses are presented in Table 4.10.

Table 4.10 Major and trace element analyses for unknown samples NL-781 to NL-794. (Oxides and trace elements recalculated volatile-free, all Fe as Fe₂O₃, TOTAL is original total before recalculation).

SAMPLE:	NL-781	NL-782	NL-783	NL-784	NL-785	NL-786	NL-787	NL-788
SiO ₂	53.43	53.08	53.43	52.77	53.02	53.85	52.79	54.72
TiO ₂	0.444	0.353	0.473	0.419	0.388	0.579	0.613	0.583
Al ₂ O ₃	11.41	11.01	12.24	11.29	10.28	13.52	13.47	13.14
Fe ₂ O ₃	10.53	12.38	10.29	10.95	11.59	10.95	12.10	11.04
MnO	0.17	0.20	0.18	0.16	0.21	0.17	0.18	0.17
MgO	13.56	13.98	10.53	12.95	14.50	7.91	8.46	7.98
CaO	7.92	7.17	9.07	9.19	8.23	9.22	8.30	7.84
Na ₂ O	2.11	1.56	2.71	1.62	1.34	3.42	2.24	2.64
K ₂ O	0.38	0.23	1.02	0.60	0.39	0.32	1.77	1.80
P ₂ O ₅	0.054	0.040	0.064	0.051	0.048	0.072	0.076	0.073
TOTAL:	99.59	99.29	99.68	100.04	100.04	99.43	99.81	99.67
LOI:	4.47	4.92	3.21	4.09	4.72	2.82	2.82	2.47
Ba	119	73	342	287	146	106	909	730
Nb	1.2*	1.4*	2.6	1.7*	0.7*	2.5	2.7	1.9*
Zr	49	36	55	47	39	62	65	59
Y	12.2	10.0	13.2	12.8	9.8	15.2	17.6	15.3
Sr	92	63	168	64	46	122	206	138
Rb	12.5	9.0	40.4	20.4	16.8	15.4	82.5	79.3
Zn	65	119	61	71	101	66	82	69
Cu	43	14	55	21	36	50	44	52
Ni	336	497	283	381	505	245	226	221
Co	72	87	66	78	85	67	69	67
Cr	1358	2754	883	1619	2308	470	453	425
V	160	155	178	170	164	206	206	199
La								
Ce								
Nd								

SAMPLE:	NL-789-C	NL-789-M	NL-790	NL-791	NL-792	NL-793	NL-794
SiO ₂	57.92	42.09	53.75	54.19	53.16	55.45	54.60
TiO ₂	0.577	0.571	0.537	0.977	1.057	1.274	1.277
Al ₂ O ₃	12.62	14.70	11.64	15.07	16.10	14.08	16.45
Fe ₂ O ₃	7.64	15.15	10.73	12.62	13.93	13.17	11.88
MnO	0.13	0.22	0.19	0.14	0.15	0.14	0.20
MgO	7.78	12.10	9.40	5.01	5.50	4.22	5.25
CaO	7.97	14.87	9.65	7.34	5.78	7.03	4.20
Na ₂ O	4.50	0.09	3.98	3.87	3.24	3.76	3.47
K ₂ O	0.75	0.16	0.06	0.67	0.93	0.69	2.46
P ₂ O ₅	0.099	0.049	0.068	0.126	0.133	0.185	0.213
TOTAL:	99.67	100.39	99.12	98.38	99.27	99.87	98.46
LOI:	1.52	5.18	2.34	2.17	2.81	1.96	2.49
Ba	201	63	14*	276	370	281	1194
Nb	2.5	2.8	2.7	4.8	6.2	8.3	10.5
Zr	52	53	50	100	110	133	162
Y	9.0	22.8	14.1	21.4	20.7	26.6	28.0
Sr	174	32	60	258	211	172	237
Rb	31.2	7.4	4.3	29.4	41.1	29.8	94.5
Zn	43	93	56	97	101	100	110
Cu	22	59	37	65	71	70	110
Ni	206	388	324	169	194	191	147
Co	49	91	72	69	69	68	62
Cr	870	868	819	52	72	71	35
V	181	266	203	222	209	206	212
La							
Ce							
Nd							

* below LLD

Application of the criteria established during the study for the identification of lava types in the Klerksdorp area shows these fourteen samples to be unambiguously representative of an almost complete section through the Klipriviersberg Group. These conclusions are illustrated by means of the following diagrams:

Figure 4.9 shows the boundary lines only of Figure 4.6, a plot of Ti:Zr vs Ti:P. The NL samples have been plotted and lie very clearly within the Klipriviersberg field.

Figure 4.9 Unknown samples plotted on Ti:Zr vs Ti:P discrimination diagram

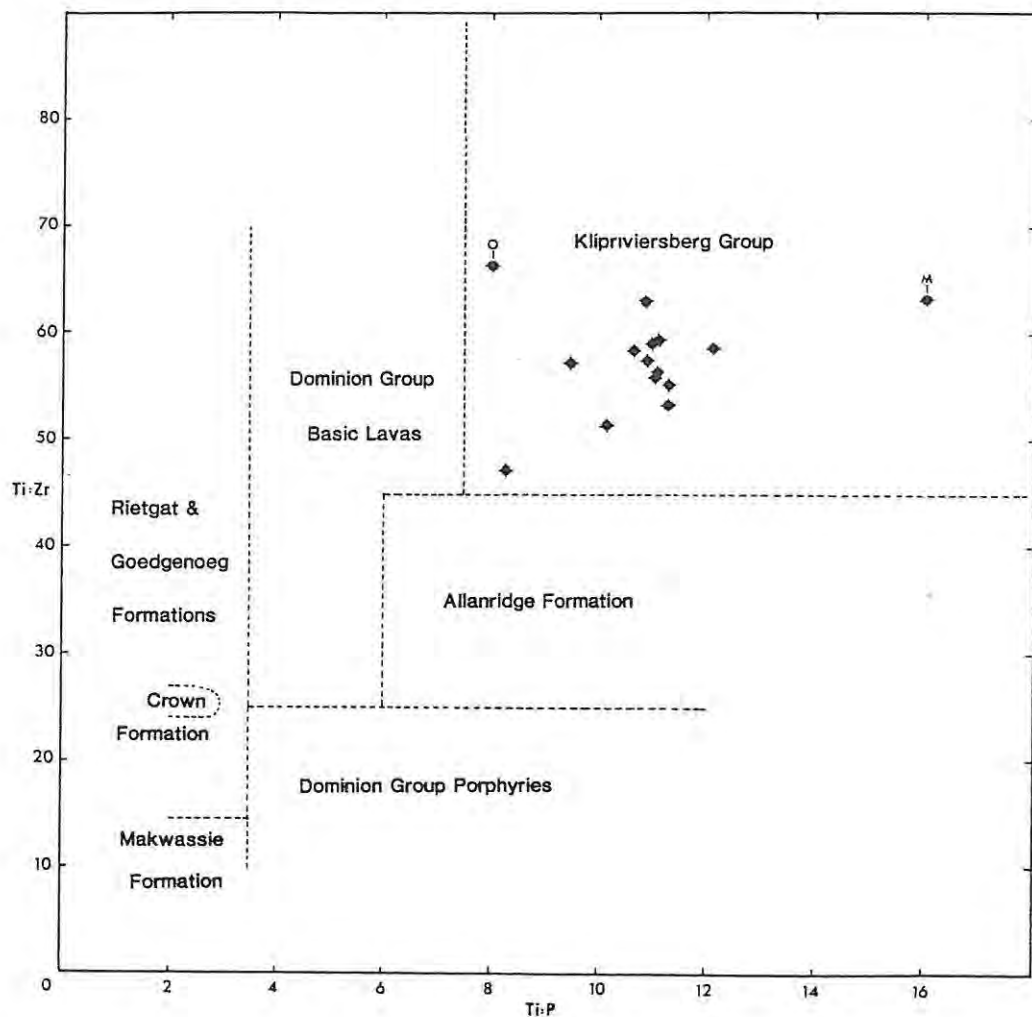
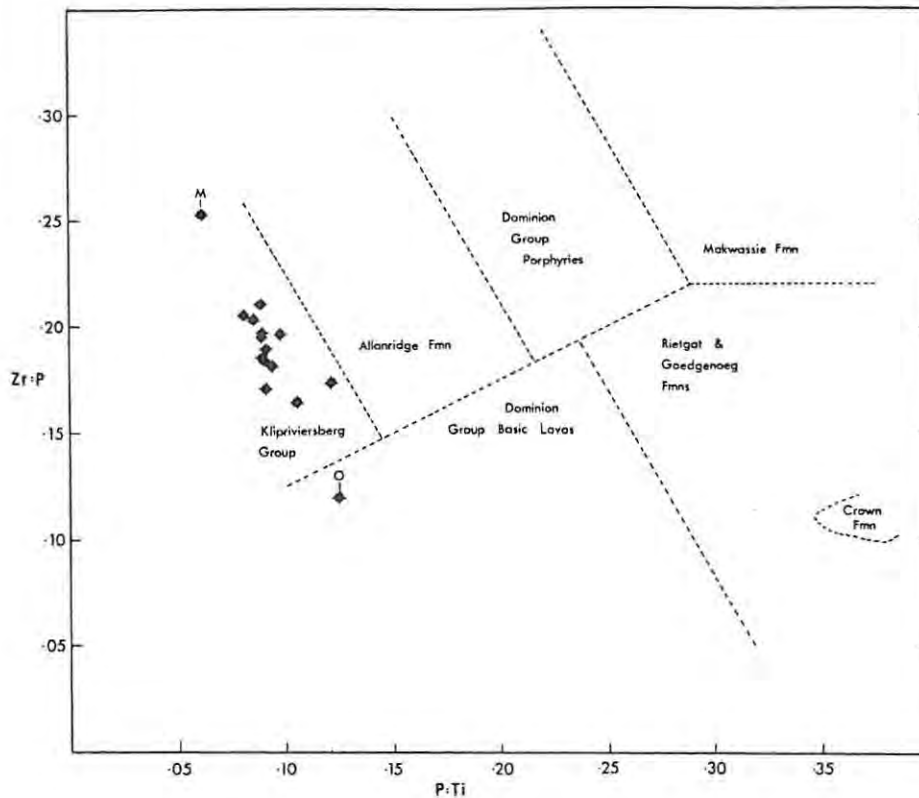


Figure 4.10 is a reproduction of the boundary lines only of Figure 4.7, which uses a different combination of the same variables. Again the NL samples alone have been shown and once more they plot within the Klipriviersberg field.

Figure 4.10 Unknown samples plotted on Zr:P vs P:Ti discrimination diagram



Having established that the NL samples have undoubted Klipriviersberg affinities, a third graph was employed, in which TiO_2 is plotted against Zr. This enables samples from the three chemically distinct units constituting the Klipriviersberg Group to be differentiated from one another. Figure 4.11, showing the boundaries of the three sets of volcanics from Figure 4.8, enables the NL samples to be subdivided as follows: samples NL-781 to NL-790 are representative of the Loraine-Edenville formations, while samples NL-791 and NL-792 represent the Orkney, and NL-793 and NL-794 the Alberton Formation respectively.

Figure 4.12b demonstrates that major element variation of the NL samples is also consistent with that of the Klipriviersberg Group. Within the boundary lines from Figure 4.12a, samples NL-781 to NL-790 plot within the Loraine-Edenville field, which displays a large variation in MgO relative to Fe_2O_3 , while samples NL-791 to NL-794 plot in the Alberton/Orkney field, in which the variation is in the opposite direction.

Figure 4.11 Unknown samples plotted on TiO_2 vs Zr discrimination diagram

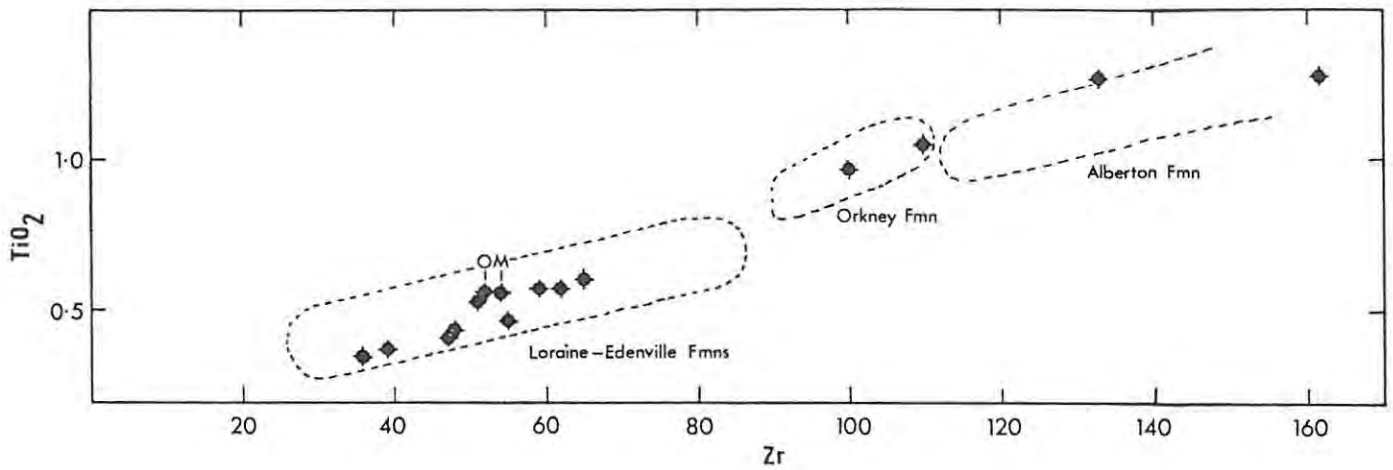
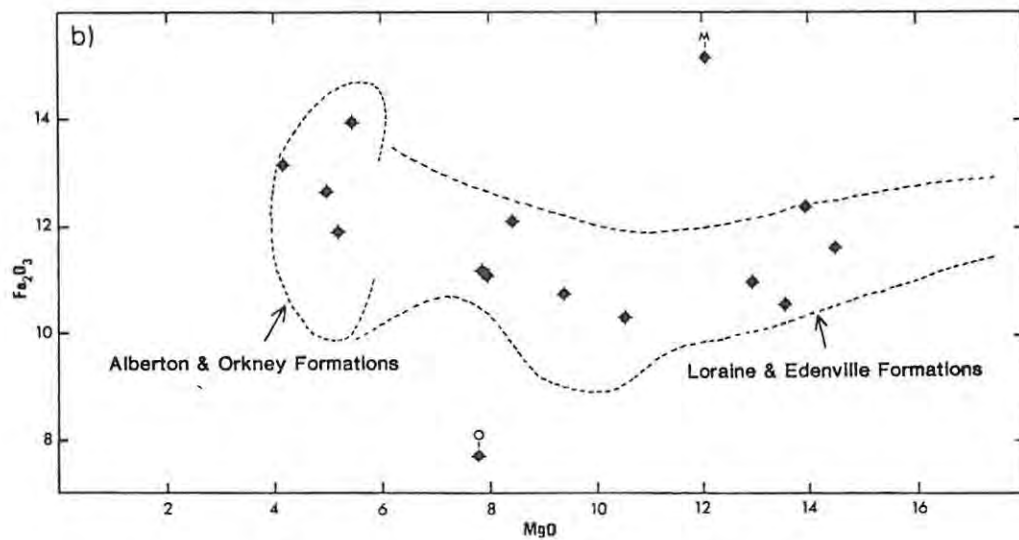
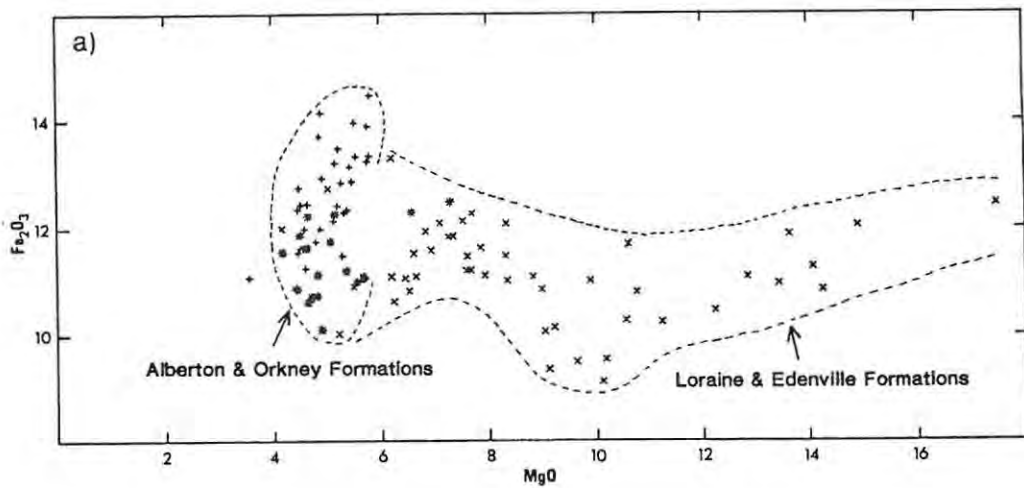


Figure 4.12 a) Fe_2O_3 vs MgO discrimination plot for Klipriviersberg Group
b) Unknown samples plotted on Fe_2O_3 vs MgO discrimination diagram.



The use of orthogonal discrimination enables classification of the NL samples to be achieved in a more quantified manner. It has been strongly indicated that they have clear Klipriviersberg affinities. Table 4.6 suggests an examination TiO_2 and P_2O_5 values as separators for formations within the Klipriviersberg Group. Table 3.2a shows that samples NL-781 to NL-790 do indeed lie within the TiO_2 range for the Loraine-Edenville Formations, i.e. between .400 and .878% TiO_2 . NL-791 and NL-792 fall within the Orkney range - .940 to 1.020% TiO_2 , while NL-793 and NL-794 have slightly higher values - 1.274 and 1.277% TiO_2 - than the Alberton Formation, whose upper limit for this study is 1.174% TiO_2 . Similar conclusions are obtained from P_2O_5 values.

Amongst the trace elements, Zr is the most useful discriminator within the Klipriviersberg Group, as shown in Figure 4.8, although it is not listed in Table 4.6 due to the presence of a few stray samples which cause overlapping of the fields.

With major element ratios, separation is possible only between the Alberton and Loraine Formations, using $TiO_2:Fe_2O_3$, $TiO_2:SiO_2$, $P_2O_5:SiO_2$ or $TiO_2:Al_2O_3$ ratios. As these are not adjacent formations, these ratios will not be helpful in the current identification process.

The Alberton and Orkney Formations can, however, be separated using Ti:V and P:V ratios. As sample NL-792, which has thus far been identified as an Orkney sample, lies very close to the Alberton field in Figure 4.11, it is of interest to check whether this is indeed an Orkney member. The range for the ratio Ti:V within the Alberton Formation is 32.08 to 38.17 (from Table 4.8c), while that for the Orkney Formation is 23.44 to 30.95. Ti:V for sample NL-792 is 30.32, confirming its Orkney affinity. Similarly for P:V, NL-792, with a value of 2.78, lies within the Orkney range of 1.97 to 3.04, while the Alberton samples lie between 3.13 and 4.73.

It can therefore be concluded that, on the basis of data accumulated from the Klerksdorp area, the unknown samples may indisputably be identified as Klipriviersberg Group lavas, with samples NL-781 to 790 representing the Loraine-Edenville Formations, NL-791 and 792 the Orkney Formation, and NL-793 and 794 the Alberton Formation.

5. STATISTICAL EVALUATION OF DATA

5.1 Graphical Representation of Data

To analyse geochemical data, Koch and Link (1971) suggest one must first decide what an anomalous observation is. They say that current practice is to make an orientation survey, to calculate the mean and standard deviation of the observations, and then to define any observation more than two standard deviations above or below the mean as possibly anomalous and any one more than three standard deviations above or below the mean as certainly anomalous.

Normal probability graph paper is scaled so that if observations are normally distributed, their relative cumulative frequency in percent plots as a straight line. The mean of the observations is estimated by reading the fifty percent point on the straight line. The standard deviation of the observations is estimated by subtracting the sixteen percent point from the eighty four percent point and dividing the result by two (Koch and Link, 1970).

For fewer than 100 observations, they suggest it is desirable but not essential to replace the relative cumulative frequency (r.c.f.) by a plotting percentage obtained by the formula

$$\text{plotting percentage} = 100 \times \frac{3 \text{ (r.c.f.)} - 1}{3n + 1}$$

where n is the total number of observations. An example of the calculation is shown in Table 5.1. The graph of the plotting percentage plotted against the upper bound of the class interval is shown in Figure 5.1.

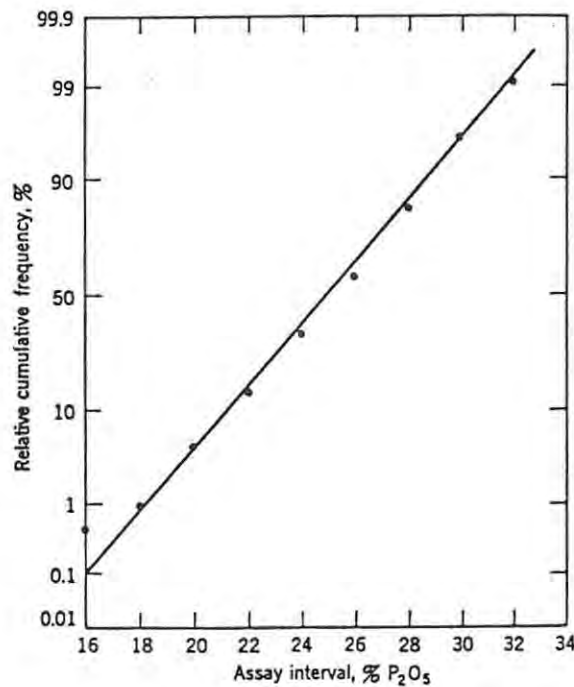
5.2 Probability Plots of Witwatersrand Triad Data

In order to examine the distribution of the chemical data of the Witwatersrand triad samples used in the current study, normal probability plots of the type discussed above were plotted using a program written by C. Beuthin for use on the Rhodes University ICL 1904S computer. The samples were first separated into their respective formations and then all major element analyses within each formation were tested for normality. The normal probability plots for major elements are shown in Figure 5.2a.

Table 5.1 Example of frequency distribution of 224 phosphate assays and plotting percentages (from Koch and Link, 1970)

Assay interval (% P ₂ O ₅)	Interval midpoint, <i>w</i>	Frequency, <i>f</i>	Relative frequency (%)	Cumulative frequency, c.f.	Relative cumulative frequency, r.c.f. (%)	Plotting percentages
14-16	15	1	0.45	1	0.45	0.30
16-18	17	1	0.45	2	0.90	0.74
18-20	19	8	3.57	10	4.47	4.31
20-22	21	21	9.37	31	13.84	13.67
22-24	23	44	19.64	75	33.48	33.28
24-26	25	54	24.12	129	57.60	57.36
26-28	27	56	25.00	185	82.60	82.32
28-30	29	30	13.39	215	95.99	95.69
30-32	31	7	3.12	222	99.11	98.81
32-34	33	2	0.89	224	100.00	99.70

Figure 5.1 Plot on normal probability graph paper of phosphate data from Table 5.1 (from Koch and Link, 1970)



A similar procedure was adopted for the trace elements, the results of which are illustrated in Figure 5.2b. Each horizontal line of graphs represents a different formation, while each column corresponds with a different element.

The Y-axis has been modified to show the mean, given at 0.00, and one, two, three or four standard deviations shown by ± 1.00 , ± 2.00 , ± 3.00 and ± 4.00 respectively. Concentrations of major element oxides (in weight percent) and trace elements (in parts per million) are shown on the X-axis.

Figure 5.3 shows several hypothetical populations in order to illustrate some of the basic patterns which occur in Figures 5.2a and b. In Figure 5.3a, three lines are plotted, each with a different mean (read opposite the "0"), but the same standard deviations. Figure 5.3b show three lines, all with the same mean, but different standard deviations. The steeper the line, the smaller the standard deviation.

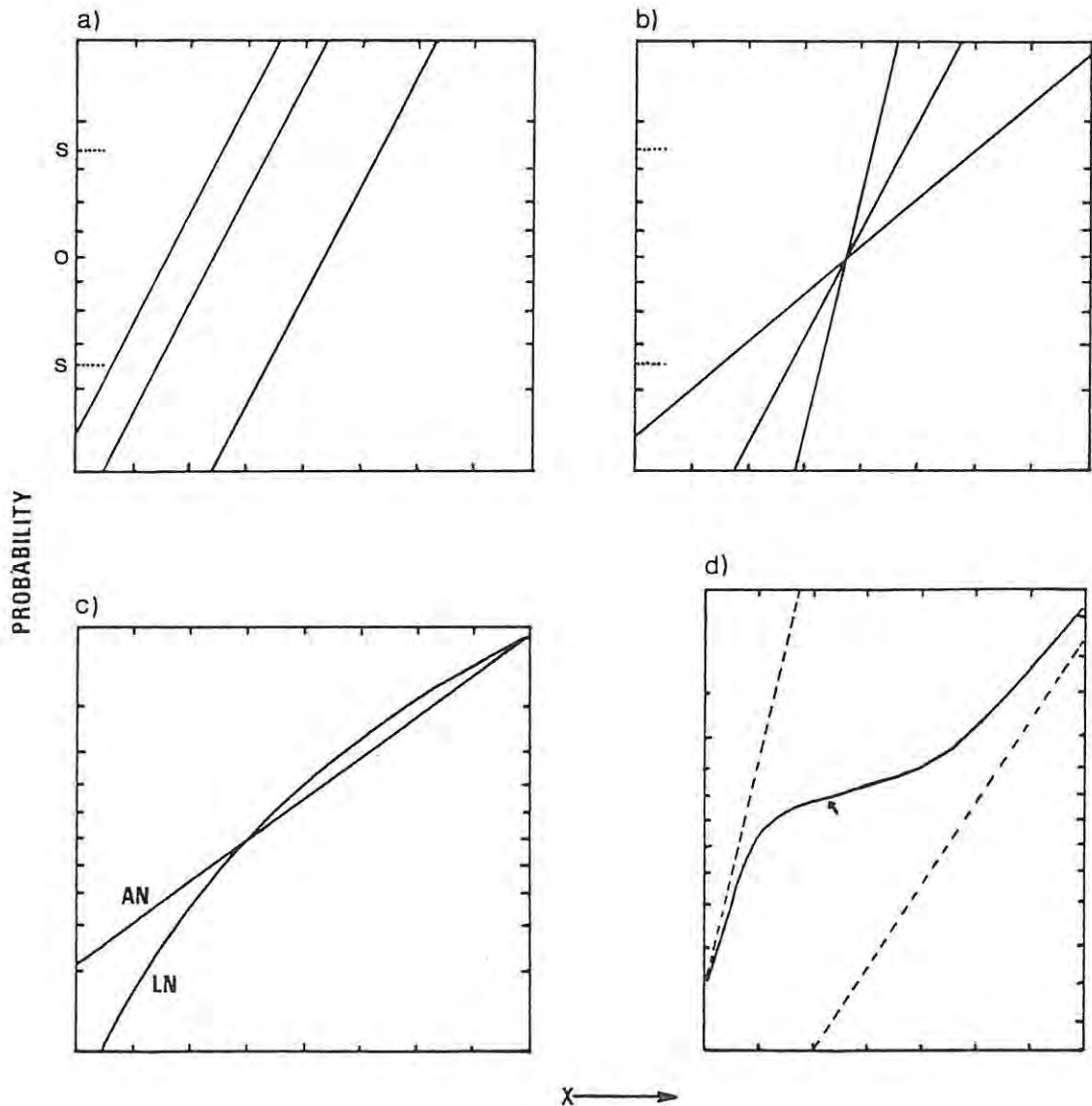
Figure 5.3c shows the difference between an arithmetic normal population, which plots as a straight line (AN), and a log normal population, which plots as a curved line (LN) on arithmetic probability paper. In general, concave-down graphs, such as LN, are skewed towards high values, i.e. the majority of points have low values, while the remainder tail off towards the higher end of the scale. The converse is also true.

Finally, Figure 5.3d shows how a combination of two normal populations produces a bimodal probability curve, which is recognised by the presence of an inflection point (marked with an arrow). Lognormal populations will produce similar patterns, with varying amounts of curvature. It is possible to separate the bimodal curve into its two constituent populations by means of a process called "partitioning". Details of the procedure are provided by Sinclair (1976), but this was not attempted in the current study.

The inflection point marks the cumulative percentile that represents the proportions of the two populations. Similarly, curves which have two inflection points indicate the presence of three populations, etc. The manifestation of more than one population may mean that that

particular element is present in more than one mineral in differing degrees of concentration, for instance, or it may represent concentrations in the matrix and phenocrysts of a porphyritic rock, etc.

Figure 5.3 Hypothetical probability plots illustrating some basic patterns (after Sinclair, 1976)



Sinclair (ibid.) suggested that, although some authors recommend a minimum of 100 values in a probability plot, as few as 70 data points seem valid and useful. The only unit in the current study which fulfils this requirement, are the Dominion Group basic rocks (69 samples). Reasonably well-defined curves are, however, produced by the Loraine-Edenville, Allanridge, Rietgat, Goedgenoeg, Dominion porphyry

Figure 5.2a Normal probability plots of major elements for each formation

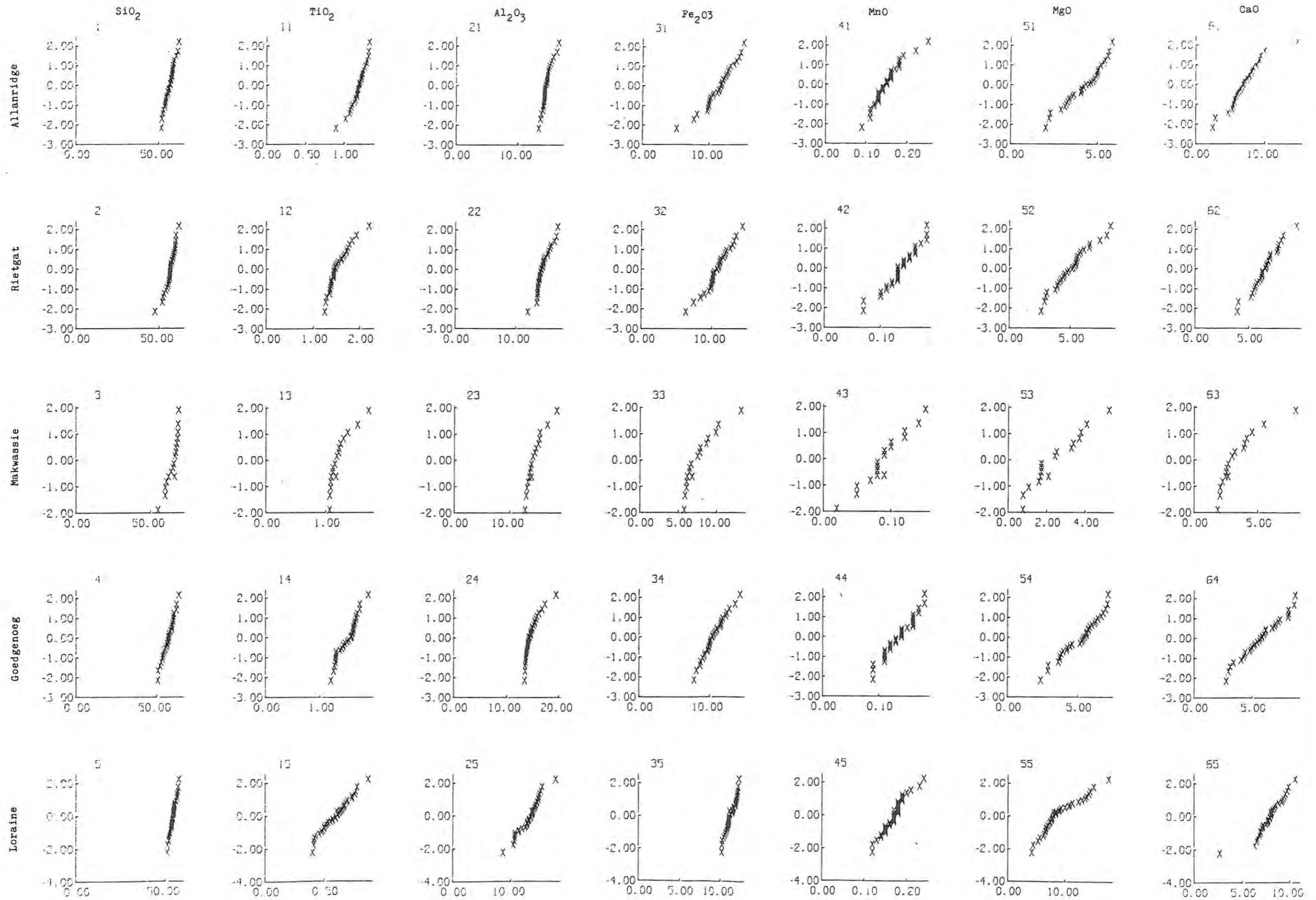


Figure 5.2a continued

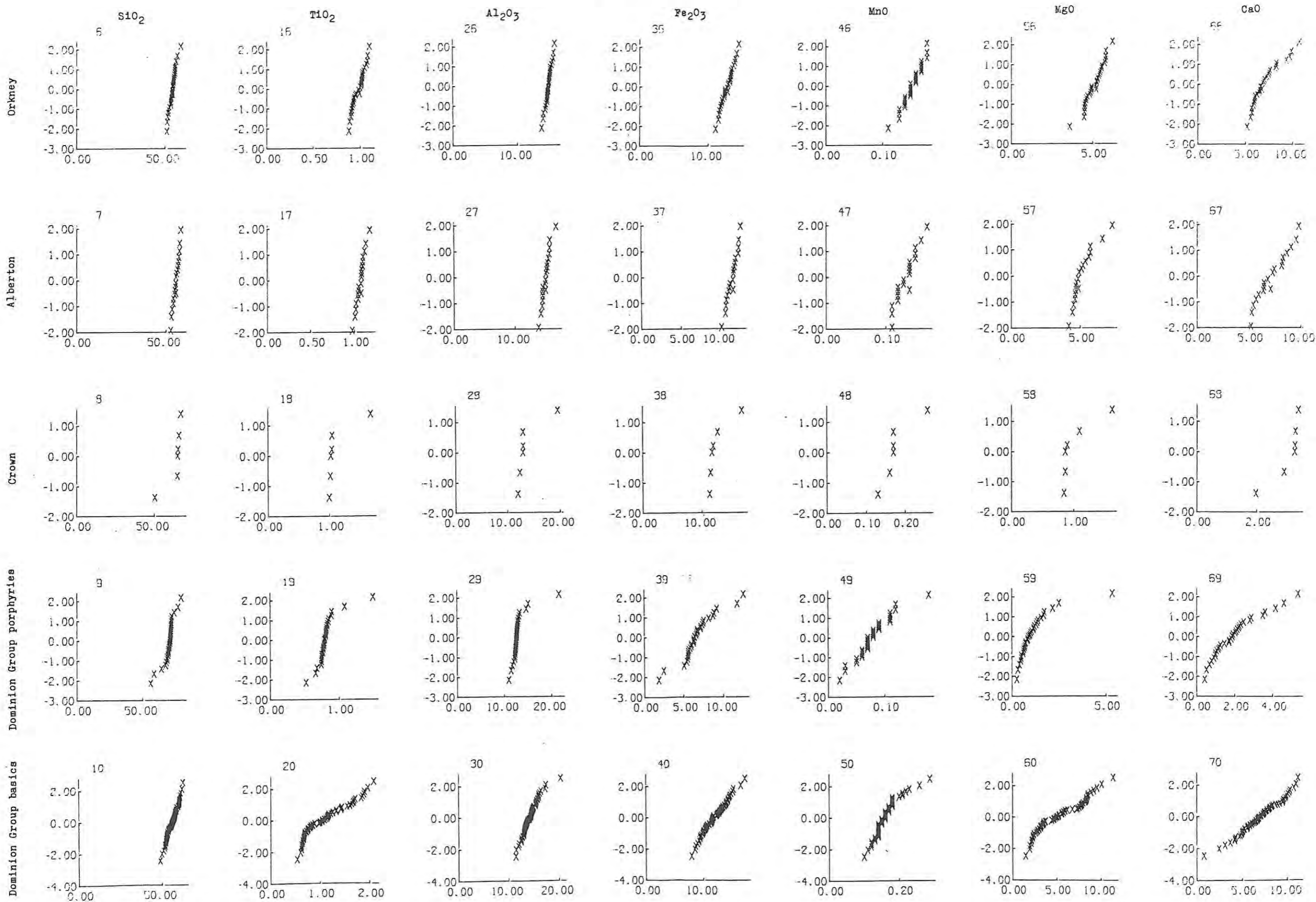


Figure 5.2a continued

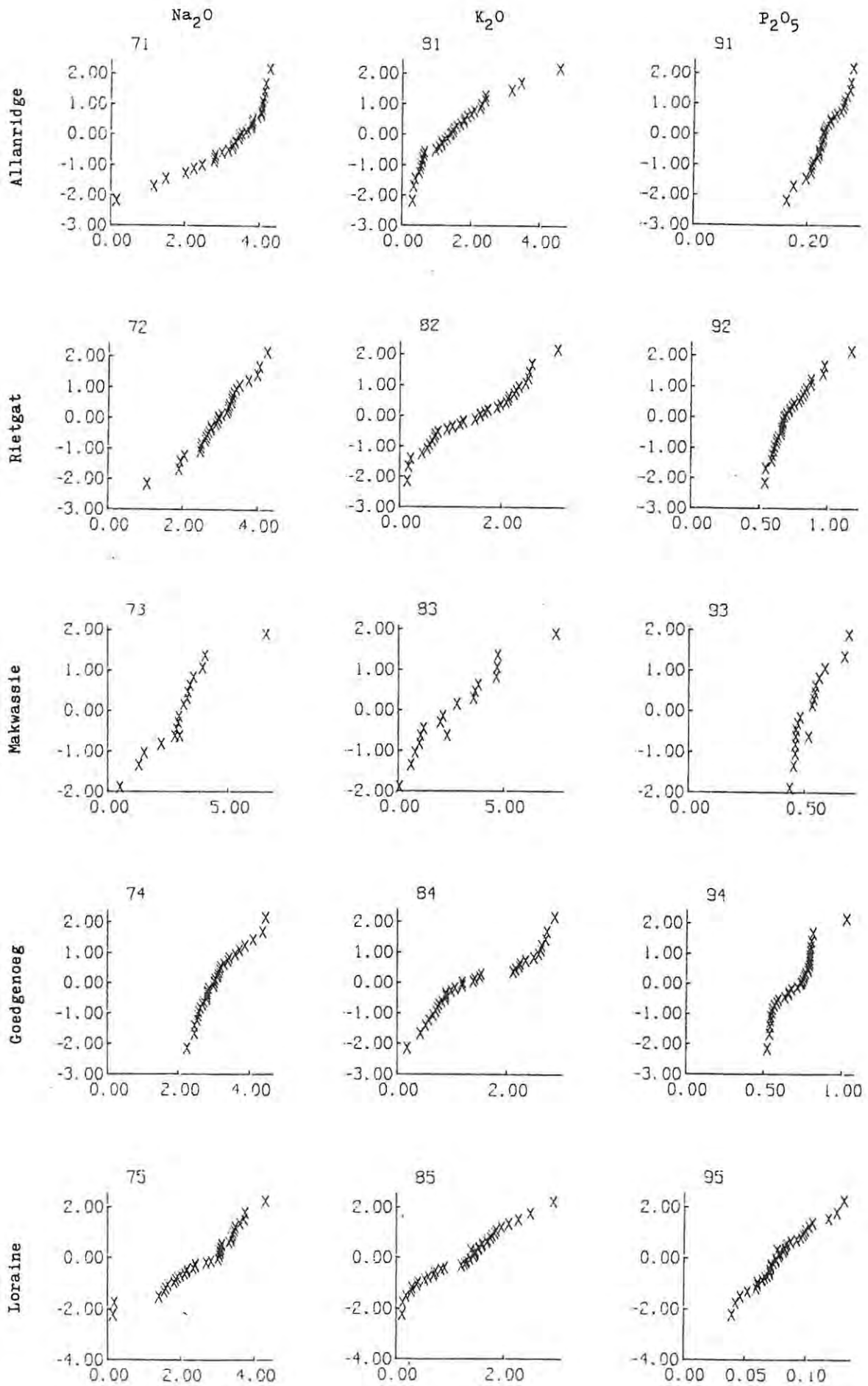


Figure 5.2a continued

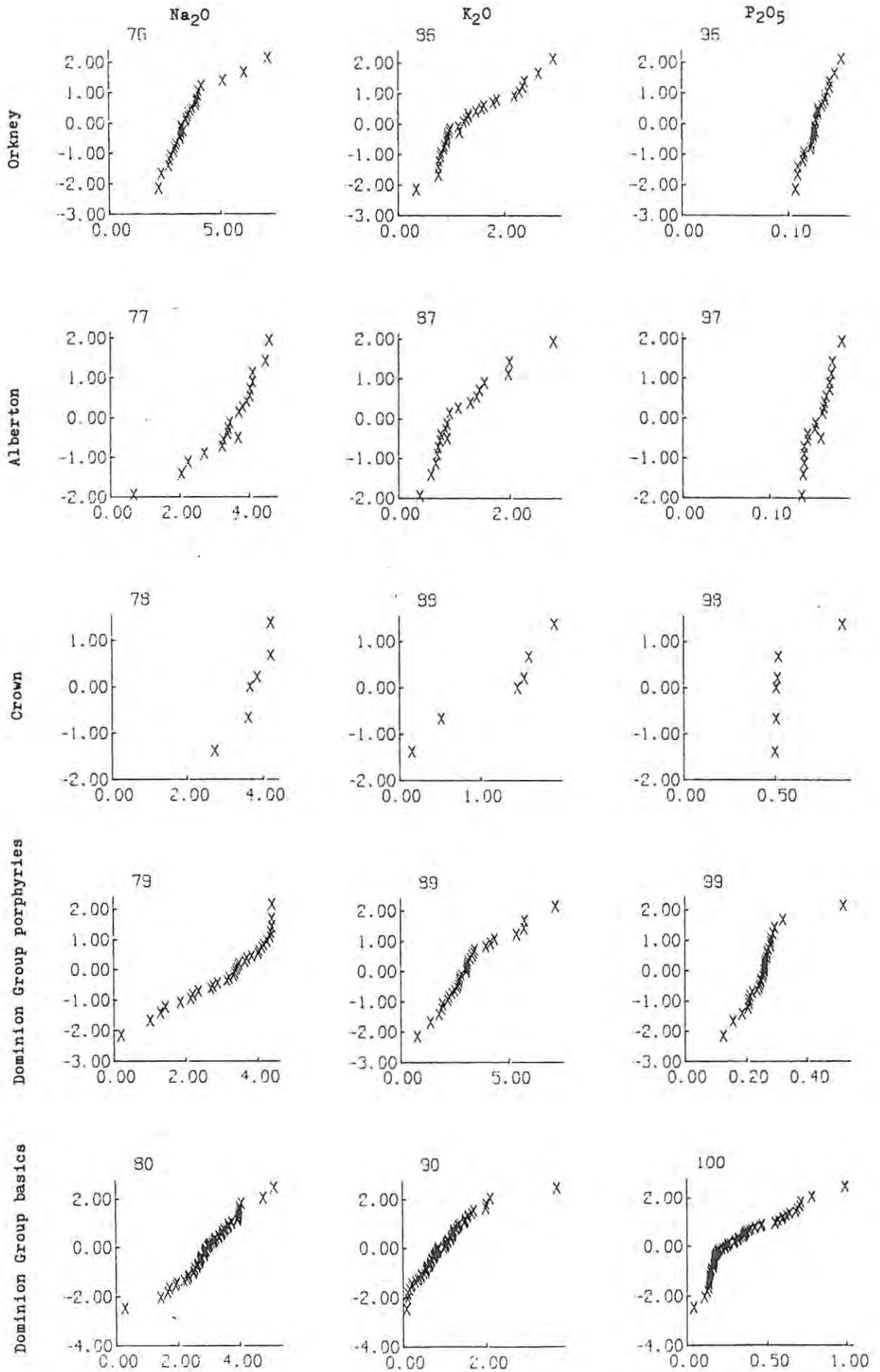


Figure 5.2b Normal probability plots of trace elements for each formation

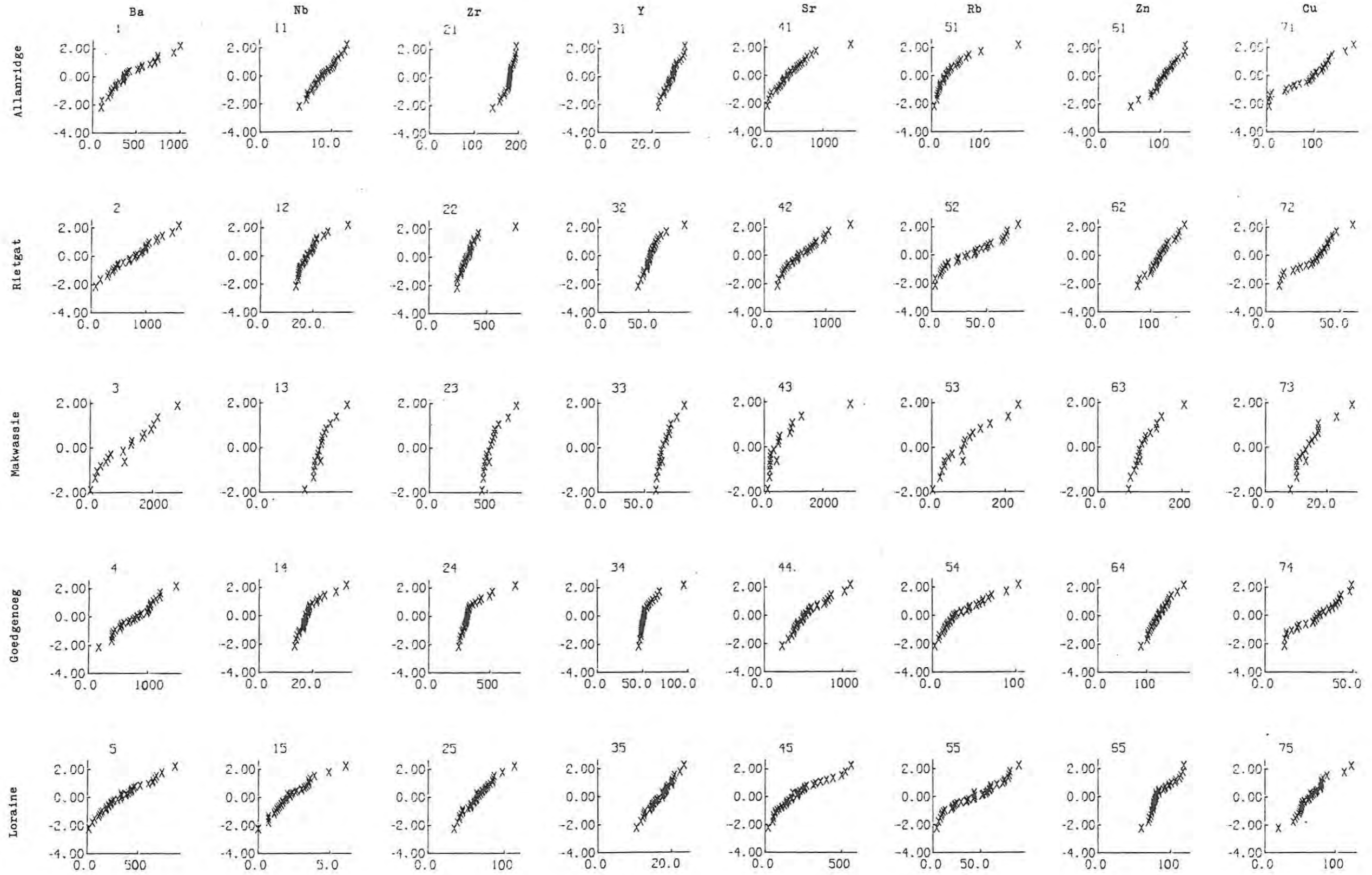


Figure 5.2b continued

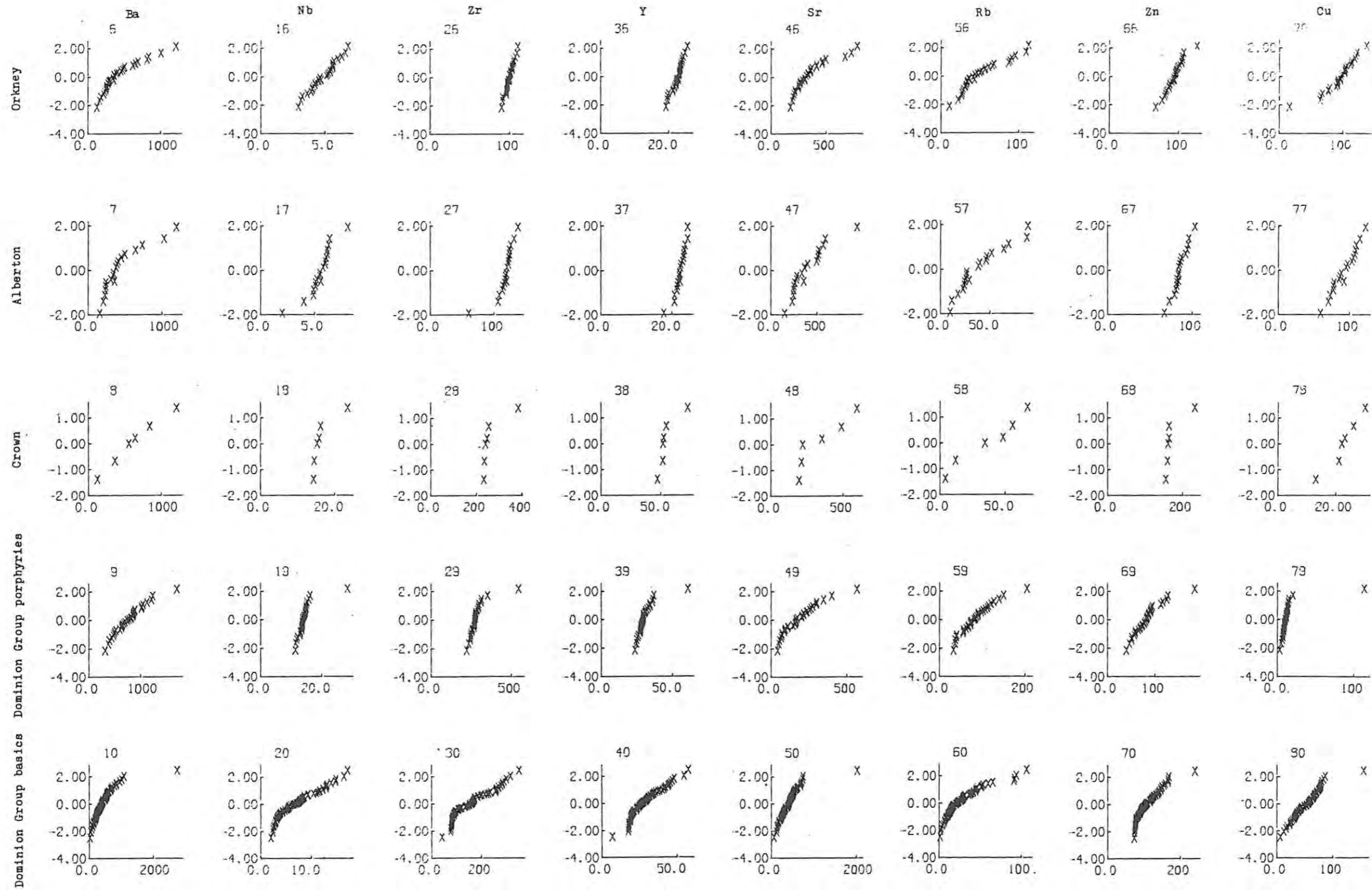


Figure 5.2b continued

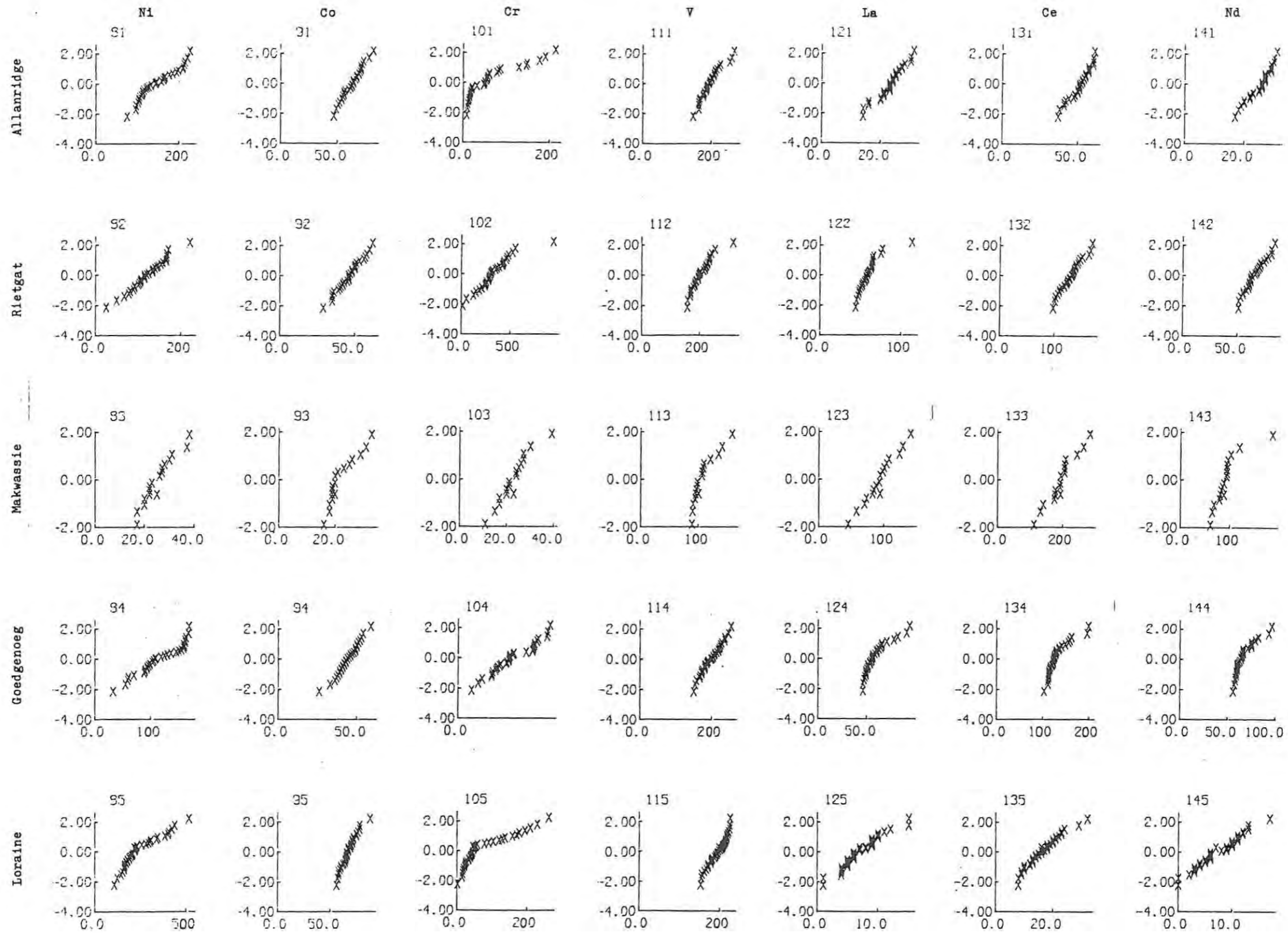
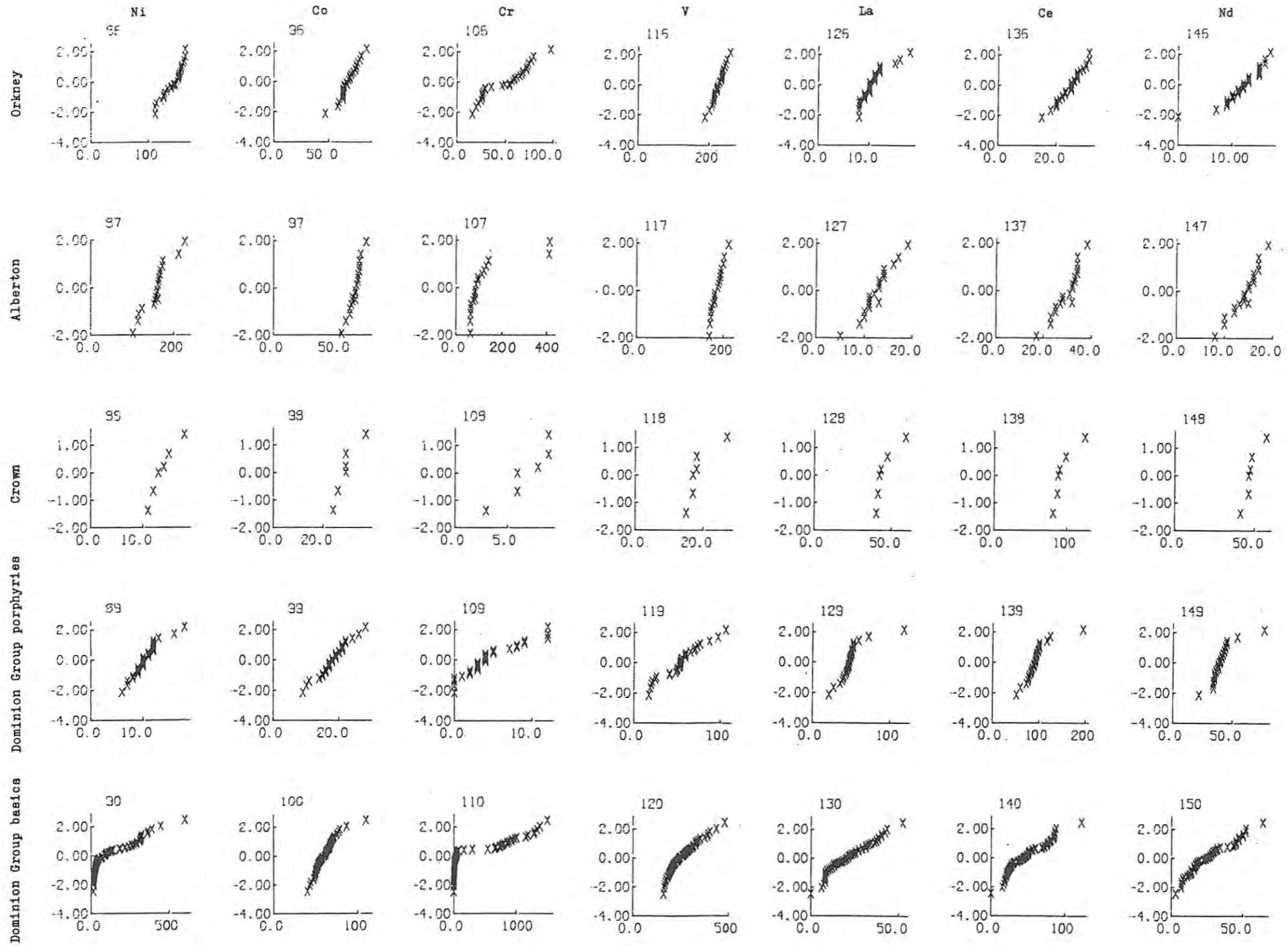


Figure 5.2b continued



and Orkney units, which have 46, 35, 34, 33, 32 and 31 data points respectively. It must be borne in mind in such cases that fluctuations in the probability curve may possibly be due to the low density of sample points, and this is certainly well-illustrated by the curves of the Alberton (19 points), Makwassie (17) and Crown (6) Formations.

One method of recognising "wild" samples is to plot confidence limits of a probability curve. This can be done graphically, but in the current study, the computer program which produced the probability plots also printed out a list of outliers for both the 90% and 98% confidence limits or intervals. An example of the output is shown in Figure 5.4. This isolates samples which plot away from the main curve, both at the tails and in between, so that they can be examined further.

Figure 5.4 Sample of output showing outlier samples for 90% and 98% confidence intervals

```

72 ELEMENT SI          FORMATION JEPP
73 FOR THE 98% CI THE OUTLIERS ARE
74 FOR THE 90% CI THE OUTLIERS ARE
75                          AI-371
76
77
78 ELEMENT TI          FORMATION LDOM
79 FOR THE 98% CI THE OUTLIERS ARE
80                          AI- 13
81 FOR THE 90% CI THE OUTLIERS ARE
82                          AI- 13
83                          AI- 69
84                          AI- 72
85                          AI- 73
86                          AI- 75
87                          AI- 91
88
89
90 ELEMENT TI          FORMATION UDOM
91 FOR THE 98% CI THE OUTLIERS ARE
92                          AI- 33
93 FOR THE 90% CI THE OUTLIERS ARE
94                          AI- 33
95                          AI-263
96                          AI-276
    
```

In Table 5.2 different confidence limits are shown together with the standard deviation which corresponds to each limit (from Spiegel, 1961).

Table 5.2 Confidence limits with corresponding standard deviations.

Confidence limit (CL)	99.73%	99%	98%	96%	95.45%	95%	90%	80%	68.27%	50%
Std deviation (s)	3.00	2.58	2.33	2.05	2.00	1.96	1.645	1.28	1.00	.6745

Thus, a 90% confidence limit implies that 90% of the samples from a particular population should be found within the interval $-1.645s$ to $+1.645s$. At the beginning of this chapter it was mentioned that Koch and Link (1971) suggested that samples which lay outside the $2s$ limit (or 95.45% CL) were possibly anomalous. By lowering this to the 90% CL, obviously a number of acceptable points would also be isolated, but it would ensure that all possibly anomalous values be scrutinised.

These normal probability plots were used in conjunction with the various methods described in Chapter 4 in attempting to classify samples into their correct formations. Naturally, some samples isolated using the probability plots were valid members of their particular group and simply represent extreme values. These were left unchanged, but certain samples were removed to other units, or left unclassified, as documented in Chapter 4.

To conclude, it is of interest to note some of the advantages and disadvantages of this simple graphical technique, as summarised by Sinclair (1976):

Advantages

1. A simple form of graphical representation of data.
2. Rapid, qualitative analysis of density distributions.
3. Rapid estimation of parameters of normal and lognormal distributions.
4. Compact graphical representation of several sets of data on a single diagram.
5. Recognition of polymodal distributions.
6. Partitioning of polymodal distributions.
7. Rapid recognition of certain abnormalities in data.

Disadvantages

1. Data might not approximate normal or lognormal density distributions.
2. Data might be too sparse for meaningful analysis on probability paper.
3. Tails of cumulative distributions are commonly not well-defined. This can cloud interpretation in the upper value range which is of particular interest in most mineral exploration data.

4. A small, but significant proportion of plots appears to be uninterpretable using the procedures outline in Sinclair (1976). This is particularly true if several populations are present with extensive overlap and/or different types of density distributions. Except in very special cases, four populations are about the maximum that can be treated successfully, and difficulties are not uncommon if three or four populations are present.
5. The lower the quality of data, the more ambiguous is the interpretation of the corresponding probability plot.

6. SUMMARY AND CONCLUSIONS

The volcanic rocks of the Witwatersrand triad were found to be primarily subalkaline tholeiitic rocks. The use of the word "andesite" for describing the majority of the basic volcanic rocks in the Witwatersrand triad in the past, is a misnomer, for two principal reasons. Firstly, the name "andesite" is generally confined to calc-alkaline rocks, associated with which are very specific tectonic regimes. The Witwatersrand triad volcanic rocks do not fit into this calc-alkaline association and M P Bowen (in prep.) has discussed this in some detail. Secondly, the Al_2O_3 content of the Witwatersrand triad volcanic rocks is too low for them to be called andesites.

The Dominion Group basic rocks, all the formations of the Klipriviersberg Group, and the Goedgenoeg-Rietgat Formations were classified as basalts, the latter formation having tholeiitic andesitic to andesitic tendencies in the non-genetic sense of Irvine and Baragar, (1971). About a third of the Loraine-Edenville samples were designated Mg-rich tholeiites. The Allanridge Formation rocks are basalts to tholeiitic andesites, while the Crown Formation samples are all tholeiitic andesites. The Makwassie Formation comprises predominantly tholeiitic andesites to andesites, but with a fair proportion of both basalts and dacites. The Dominion Group porphyries are mostly dacites, with an andesitic to tholeiitic andesitic tendency.

With regard to potassium content, the Crown Formation rocks are K-poor, the Alberton, K-poor to average. The Allanridge, Orkney and Dominion basalts are average, the Dominion porphyries are average to K-rich, and the Loraine-Edenville, Goedgenoeg-Rietgat and Makwassie Formations are K-rich to average. The Klipriviersberg Group becomes progressively more potassic upwards.

Major and trace element analyses have revealed that a clearly-defined geochemical stratigraphy exists within the Witwatersrand triad volcanic rocks. The Dominion Group consists of bimodal basaltic and dacitic lavas, both of which transgress the formal formational boundaries defined by SACS (1980). Within the overlying Witwatersrand Supergroup, the Jeppestown Amygdaloid or Crown Formation occurs in the study area, with distinctive chemistry, but the Bird amygdaloid marker is absent.

In the overlying Ventersdorp Supergroup, the lower Klipriviersberg Group, the middle Platberg Group and the upper Pniel sequence are each chemically distinctive units, with further subdivisions possible within each of these units. The lowermost Westonaria Formation of the Klipriviersberg Group is absent in the Klerksdorp area, being confined to only the deepest parts of the Witwatersrand basin, and the Jeannette Agglomerate Formation has not been recognised. Three chemically distinct units are present - the Alberton Formation, the Orkney Formation, and the Loraine and Edenville Formations together, which form a single geochemical entity.

The Platberg Group comprises three stratigraphic units, two of which, the Goedgenoeg and Rietgat Formations at the base and top respectively, have the same chemical signature, while the intervening Makwassie Formation is chemically distinct. This group is overlain by the Pniel sequence, the only volcanic unit of which, the Allanridge Formation, is also chemically distinctive.

Several techniques were used to distinguish between the 9 different geochemical units present. The simplest of these, orthogonal discrimination, assesses the range of each major and trace element, as well as that for each interelement ratio, for every geochemical unit. Discrimination is achieved when ranges are discrete, i.e. do not overlap. This method is successful in distinguishing all geochemical units from one another, except the Dominion Group basalts from the Allanridge, Loraine-Edenville, Orkney and Loraine basalts. The other techniques, however, do enable these distinctions to be made successfully.

The second discrimination technique involves the use of three discrimination plots, namely Ti:Zr vs Ti:P, Zr:P vs P:Ti, and TiO_2 vs Zr. The first two are able, with a minimum of overlap, to distinguish seven separate units, the three Klipriviersberg entities plotting together in one field. The third plot, TiO_2 vs Zr, enables these three Klipriviersberg units to be separated from one another.

The third and most sophisticated discrimination technique, discriminant analysis, requires the use of a computer, and has two main functions. Firstly, previously grouped data is assessed and an evaluation of the

success of the parameters defining the groups is provided. In the current study, 99.07% of all the samples were correctly classified. The second function of discriminant analysis is to classify unknown samples. In this respect a number of intrusive samples from the current study were able to be assigned to various formations.

A useful test of the aforementioned discrimination techniques was afforded when fourteen samples from an unknown succession in the Klerksdorp area were acquired. They were positively identified as representative of a complete sequence through the Klipriviersberg Group, two of the samples being from the Alberton Formation, two from the Orkney Formation and ten from the Loraine-Edenville Formations.

Diagrams of major and trace element variation with depth show that, despite their having been subjected to low grade greenschist metamorphism, variation patterns are still reasonably clear. The greatest amount of scatter is shown by Na, K, Mn, Ba and Rb, and these elements should not be used for any petrogenetic deductions. In contrast, the immobile elements Ti, P, Nb, Zr and Y, together with the light rare earth elements La, Ce and Nd, appear to have been little affected by any post-extrusive processes.

Normal probability distributions of each major and trace element for each formation were plotted, enabling a qualitative estimation of the density distributions to be made. Confidence limits placed on these curves served to isolate "wild" samples, which were then examined further using discrimination techniques to clarify their true affinities.

With regard to the SACS (1980) nomenclature of the individual formations included in this study, several changes are recommended. It is proposed that the name "Rhenosterhoek Andesite Formation" be shortened to "Rhenosterhoek Formation", as the lava type present is primarily, but not exclusively, basaltic. For the same reason, the name "Allanridge Formation" is proposed in place of "Allanridge Andesite Formation".

The original purpose of this study, namely, to ascertain whether any geochemical criteria exist which may serve to discriminate between the various volcanic formations within the Witwatersrand triad, has been fulfilled with highly encouraging results. The successful identification of a sequence of unknown samples has demonstrated that the various techniques used have undoubted application in exploration and development in the Klerksdorp area. Chemical data from south of Johannesburg (Wyatt, 1976) appear to uphold these criteria in that region, and it remains for further analytical work to show whether these criteria can be extended to include areas further afield. In addition, closer sampling on either side of formational boundaries is required in order to resolve situations in which geochemical and visual delineation of boundaries does not concur.

7. ACKNOWLEDGEMENTS

The writer would like to thank the following people :

Professor H.V. Eales, who initiated this study and who, together with Dr Goonie Marsh, offered suggestions and criticisms during its progress;

Michael Bowen, for his brotherly cooperation;

Mr Meiring Strydom and the staff of Anglo American Prospecting Services, Klerksdorp, for providing access to borehole core and related information;

Anglo American, for the provision of a grant to finance this study, and the C.S.I.R. for a post-graduate bursary;

David Bowen, Mark Hannam and Julian Misiewicz, for help in the preparation of samples; Archie Cintso for making thin sections and Dennis Gouws and Johan Kruger, for their help with the spectrometer;

Craig Beuthin, Hugh Murrell and Pat Terry, for assistance with graph plotting;

Shirleen Reynolds and June McDougall, for typing this manuscript.

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APPENDIX I

HISTORY OF INVESTIGATION OF THE WITWATERSRAND TRIAD LAVAS

The literature review which follows is the result of an attempt to collect together as much published information as possible about the stratigraphic relationships, nomenclature, correlation and general descriptions of the volcanic rocks of the Witwatersrand triad. Much unpublished material has also been referenced, but this review has not exhausted all possible sources. The literature has been reviewed more or less in chronological order and the spelling of names, descriptions, terminology and views expressed are those of the original authors. The locality map in Figure A.1 shows the positions of places mentioned in the review.

The important contributions of the early writers, the most prolific of whom were Molengraaff (1903a, 1903b, 1904, 1905), Hatch (1903a, 1903b, 1903c, 1904a, 1904b), Corstorphine (1903), Holmes (1904, 1906), Rogers (1905, 1906, 1907, 1908) and du Toit (1906, 1907, 1908), were collated and consolidated in a book by Hatch and Corstorphine (1909). The next set of major contributions included those of Kynaston (1907, 1912, 1913), Humphrey (1909, 1910a, 1910b, 1911), Horwood (1910, 1912a, 1912b), Mellor (1911, 1912, 1913), McDonald (1911) and Rogers (1921, 1922, 1929). Rogers (1929) provided a comprehensive overview of the current state of knowledge of the Ventersdorp System to that date. Subsequently, more work was done by Nel (1927, 1933, 1934a, 1934b, 1935) and by Nel and others (1935, 1939), while Jacobsen (1943) produced a detailed thesis on the Ventersdorp System. Truter (1949) provided a detailed review in a comprehensive overview of all volcanism in South Africa. Further investigations by Von Backström (1952, 1962), Von Backström and others (1953), Matthysen (1953), Pienaar (1956), Coetzee (1960) and Winter (1963b, 1965b) provided additional detailed information, with the literature to that date being reviewed exhaustively by Winter (1965b). A very useful bibliography was provided by Simpson (1964). More recent studies included that of Whiteside (1970) and Wyatt (1976), while the reader is referred to SACS (1980) for the latest recommendations regarding nomenclature and correlation of stratigraphy.

Before the turn of the century the amount of literature published on any of the Witwatersrand triad lavas was negligible. The earliest account of the Ventersdorp rocks was that of Wyley (1859), who wrote: "About five or six miles below Hope Town, the Orange River flows between rocky hills of greenstones, amygdaloid, claystone and other traps.". This was followed by Stow's (1874) map and description of the "amygdaloidal and associated rocks of Pniel", on the Vaal River near Barkly West.

Despite the occasional heretical statement from respected geologists such as Hans Merensky (1907), who emphatically announced that "amygdaloidal diabase, not kimberlite, is the source of river diamonds", the observational powers of the geologists of that time were acute and much of their original mapping and descriptions are still valid today.

The year 1903 heralded the beginning of an interest in these and associated rocks which lasted for about ten years, waned and then revived again during the 1930's. Molengraaff (1903a), in describing the strata circumscribing the central granite boss of the Vredefort Mountain-Land, reported the presence of "amygdaloidal diabase and diabase porphyrite merging gradually into one another" and that "these rocks bear a very close resemblance indeed to the amygdaloidal rocks of the Klipriversberg and the Suikerboschrand."

Sawyer (1903), in a discussion about an unnamed amygdaloidal diabase formation near Heidelberg, recognised that in places its flows were conformable to the underlying Witwatersrand Beds, while in others the relationship was clearly unconformable. He also found that in some places the diabase overlies the Elsburg Reefs, but in others it overlies much older beds.

Hatch (1903a), describing two sections in the Potchefstroom District, observed "one great unconformity ... between the Upper Witwatersrand Beds and the Klipriversberg Amygdaloid", and probably another between the Black Reef and the Klipriversberg Amygdaloid. He deduced that, due to its vesicular structure, the amygdaloid was a flow which had spread over the Witwatersrand Beds after they had been tilted. He also noted the great disparity in thickness between the south-eastern and the

north-western portions of the amygdaloid and added that this variability in thickness was characteristic of the Klipriversberg Amygdaloid in the Southern Transvaal, and concluded that the amygdaloid must be a series of flows.

Also in 1903, a hitherto unrecognised formation, younger than the Witwatersrand Beds, but older than the Black Reef Series, and comprising "shales, schists and shaly conglomerates, banded cherts, amygdaloidal diabase tufas, grits and very coarse conglomerates", was described by Molengraaff (1903b) from the south-western and southern Transvaal. He noted its unconformable relationship to both the underlying Witwatersrand Beds and the Black Reef Series above, but was unsure of its position regarding the "Klipriver Amygdaloid" and the "Vaal River Amygdaloid".

Hatch (1903b) wrote about a very thick, coarse conglomerate developed between the Black Reef and the Klipriversberg Amygdaloid near Reitzburg in the Orange River Colony, and correlated it with that referred to by Molengraaff (1903b).

His interest being revived by Molengraaff's earlier 1903 paper, Hatch (1903c) described a boulder bed containing boulders of various rock types (conglomerates and quartzites of the Upper Witwatersrand Beds, banded slates from the Lower Witwatersrand Beds, quartz, gabbro, diabase and quartz-porphyry) ranging up to several feet in diameter. This boulder bed, which occurs north-west of Ventersdorp, lies unconformably on the Lower Witwatersrand Beds. Hatch was convinced that this formation was the same as that described by both Dr Molengraaff and himself previously, and proposed for it the name "Ventersdorp Beds", and suggested that it include the underlying Klipriversberg Amygdaloid, which he showed was intimately related to it.

Dörffel (1903) wrote that "the Black Reef Formation on Kromdraai (NW of Pretoria) is deposited unconformably on a schistose conglomerate, belonging to the new formation described lately by Dr Molengraaff, Dr Hatch, and others, and for which Dr Hatch had suggested the name Ventersdorp Beds."

Corstorphine (1903) described the same conglomerate in the Krugersdorp area and in the valley of the Suikerbosch Rand River, also associated with masses of diabase and having the same relationship to enclosing rocks. He was convinced that this diabase was a geological continuation of that of the Klipriversberg, and that this together with the amygdaloid, quartz-porphyrines and quartzites of the Vaal River, Kimberley and Hopetown areas, constituted a major volcanic series.

Hatch (1904a) reported the presence of "Ventersdorp beds" north of the Dwarsberg in the Marico District, but noted that they rest immediately on the Basement Granite, not on Lower Witwatersrand Beds, as at Ventersdorp. In a supplement to this paper, Hatch (1904b) described a thin section of amygdaloidal diabase, "composed of lath-shaped crystals of plagioclase feldspar, imbedded in a chloritic mass derived probably from the decomposition of augite; the amygdules consist of calcite and quartz."

Hatch and Corstorphine (1904a) described the amygdaloidal diabase and diabase breccia occupying the Bezuidenhout Valley and postulated that it belonged to the Ventersdorp beds now called the Ventersdorp Series. They noted that quartz- and chalcedony-filled amygdules were abundant and that occasionally a porphyritic habit was observed in which the phenocrysts were stellate aggregates of feldspars. Microscopically, they described the diabase as a mass of minute felty laths of feldspars and augite with interstitial quartz, while the larger feldspars were sometimes orthoclase, sometimes plagioclase. The diabase breccia included mainly basic igneous fragments similar to the amygdaloidal diabase, but also pieces of quartzite. The breccia matrix was crystalline, containing mostly small feldspar crystals, with a kaolin cement.

To the north of the valley, the amygdaloid becomes slaty and is associated with slates, quartzites and conglomerates similar to those classified as part of the Ventersdorp Series in the Western Transvaal. From their observations, Hatch and Corstorphine inferred an episode of "long-continued volcanic activity, producing alternate outpourings of lava and the ejection of fragmental material, (accompanying) the deposition of the sediments.". In addition, they suggested "removing (the Elsburg Series) from their hitherto accepted position at the top

of the Witwatersrand Series, and including them with the Ventersdorp Series.". Molengraaff (1904) seems to have also given support to this idea, and categorically stated that he recognised "the Klipriver Amygdaloid as representing a separate system - the Vaal River System - overlying the Witwatersrand Beds unconformably".

In a discussion on the paper by Hatch and Corstorphine (1904a), Frames (1904) suggested that the Elsburg Series and the diabase in the Bezuidenhout Valley had been inverted by localised thrust faulting from the south, which induced the slaty structure displayed at the bottom of the diabase. In two drill holes he found a "transition layer between the (diabase) sheet and the sedimentary beds composed of an intimate intermixing of grains of sand with igneous matter." In other instances they were abruptly separated by a "3 to 5 inch layer of quartzite". He accepted this as strong evidence that the ancient lava flow interrupted the deposition of the sedimentary strata, resulting in conformity between the Elsburg Series and the diabase. He concluded that the diabase was poured out as a lava, when the underlying beds were still horizontal, and was later tilted along with the Elsburg Series. Frames did not think it advisable to include the diabase and the Elsburg Series in the Ventersdorp Beds, until more evidence was collected, his reasoning being that the latter had characteristics not present in the Elsburg Series, viz. the Ventersdorp Beds were largely made up of fragments of pre-existing conglomerate, while the Elsburg Series of conglomerates were true and original pebble beds. In reply to Frames, Hatch and Corstorphine (1904b) reiterated their conviction that the Amygdaloidal Diabase was part of the Ventersdorp Series, but were not entirely sure whether the Elsburg Conglomerates were.

Holmes (1904) recognised amygdaloidal diabase and volcanic breccias belonging to the Ventersdorp Series near Vryburg. He found that to the west the diabase thins out to expose the underlying boulders of granite imbedded in the diabase. Further to the west he noted that the granite was penetrated by both acid and basic dykes, which appear to be older than the overlying Black Reef Formation.

Luttman-Johnson (1904) detailed the geology of the Fortuna Valley, Heidelberg, in which the Hospital Hill Quartzites of the Witwatersrand Series are unconformably covered by a great thickness of amygdaloidal diabase of the Ventersdorp Series, together with associated rocks, viz.

breccia, tuffs, conglomerates and quartzites. The coarse breccia is found interbedded with, and grading into, the fine tuffs. The Black Reef overlies the Ventersdorp Series unconformably.

Jorissen (1906) agreed with Hatch and Corstorphine (1904a) and Molengraaff (1904) in separating the Elsburg Series from the Witwatersrand System and placing it as the basal series of the Vaal River or Ventersdorp System.

The name "Dominion Reef Series" was given by Molengraaff (1905) to a series of reefs found in the Klerksdorp district. The reefs comprise sericite schists interbedded with gold-bearing conglomerates and quartzites, and are overlain by amygdaloidal diabase and porphyrites of what Molengraaff mistakenly called the Vaal River System, even though he saw clearly that they underlay conformably the Witwatersrand Beds. He later contradicted himself by saying that "the Dominion Reef Series and the Lower Witwatersrand Beds disappear under the overlying amygdaloidal diabase of the Vaal River System.". He evaded explaining this inconsistency by concluding that "a great diversity appears to exist in the sequence of the strata of the Vaal River System ...".

Two groups of amygdaloidal rocks, called the Beer Vley series and the Zeekoe Baard amygdaloid, found in the Prieska and Hope Town districts, were described by Rogers (1905). The Beer Vley group consists of andesitic amygdaloidal lavas, rhyolites and some agglomerates. He noted that this group is less altered than the Zeekoe Baard amygdaloids which are more basic and which he said contain very few recognisable original constituents. Rogers noted the similarity between the Zeekoe Baard amygdaloid and those in the Transvaal underlying the Black Reef Series. In the Prieska area, however, the correlation proved difficult and he suggested that these amygdaloids belonged to the same group as the "Boschveld volcanic rocks" of the Transvaal.

Samples of a number of igneous rocks from the Transvaal were collected by Rand (1905), who published descriptions of their petrographic features. The samples included two from the Klipriviersberg Amygdaloid, six from the "eruptive occupying the floor of the Bezuidenhout Valley" and three from the vicinity of Klerksdorp.

Rogers (1906) mentioned the presence of the "Pniel Volcanic Series (Ventersdorp Beds)" between Barkly West and the Langeberge in the northern Cape. He stated that they undoubtedly underlie the Campbell Rand Group and the Black Reef Series. They consist of compact and amygdaloidal lavas, some of which are porphyritic, breccias and tuffs. He noted their resemblance to the lavas and breccias at T'Kuip, which have a lower basic portion, overlain by a thick group of porphyritic acid lavas.

In his survey of the Vryburg-Mafeking area, Du Toit (1906) detailed the occurrence of the Ventersdorp Series, consisting of lavas, breccias, conglomerates and other rocks of both sedimentary and igneous origin. In the Vryburg area he reported the presence of what he called the Zoetlief Beds, which unconformably underlie the Ventersdorp diabases and amygdaloids, but due to the similarity of rock types, he preferred to group the two formations together.

The Zoetlief Beds described by Du Toit consist of a lower portion of arkoses and quartzites, resting upon granite. This is followed by a middle portion consisting of acid and intermediate lavas: quartz porphyries, trachytes and allied rocks. The uppermost division comprises quartzites and flagstones, which grade into cherts. He noted that the quartz-porphyries and trachytes had been highly silicified.

The Ventersdorp Diabase formation contains compact and amygdaloidal diabases, volcanic tuffs and breccias, conglomerates, grits, sandstones and shales. Du Toit recognised finely developed flow-structures in the volcanic rocks, which are mostly andesitic and usually altered. In one area he detected a "pillowy-structure" in the lava, but he maintained that the greater portion of the volcanic material was laid down sub-aerially.

Hall and Humphrey (1906) reported the presence of Ventersdorp Beds (mainly conglomerates) north of Krugersdorp. These were covered by schists and various sedimentary rocks, which Hall and Humphrey decided were best included in the Swaziland Series. These in turn are overlain by the Black Reef Series.

Holmes (1906), in describing the geology of the south-western Transvaal, stated that the predominating rock of the Ventersdorp System in this area is amygdaloidal diabase, the nature of which had already been commented upon by previous authors. The second most important rocks are the acid quartz porphyries, well-exposed in the "Makuasiberg". They contain porphyritic crystals of feldspar and quartz and are highly resistant to weathering. South-west of Wolmaransstad he noted the occurrence of a porphyrite dyke, 23 metres wide, cutting through the quartzites underlying the diabase. He was unable to determine the exact relationship between the diabase and the porphyrite, but he postulated that the porphyrites occurred as intrusive dykes rather than interbedded lava sheets, and classed them together with the Ventersdorp System Rocks.

Holmes also recorded the presence of volcanic conglomerate containing fragments and boulders imbedded in a lava matrix, which is both overlain and underlain by amygdaloidal diabase. He concluded that "the sources of origin of the enormous masses of diabase which form the bulk of the Ventersdorp Beds are not apparent; they possibly belong to the class known as fissure eruptions, though anything of the nature of volcanic cones would have long ago become unrecognisable in a formation so ancient."

Du Toit (1907) in a survey of the eastern Griqualand West district found the same two volcanic formations as he had in the Vryburg-Mafeking area, i.e. the Zoetlief Beds and the Diabase Formation, which he now grouped with the Pniel Volcanic series of Stow, who in 1873 had given this name to the amygdaloidal and associated rocks of the Vaal River Valley. Again he found that the Zoetlief Beds were unconformably overlain by the diabase, which, at the contact, often contains angular fragments of the underlying rhyolite. He determined the presence of three zones in the Pniel series: a lower zone of porphyritic lavas (acid and basic), a zone of normal diabases, and an upper zone of quartzites and conglomerates. Tuffs, breccias and brecciated lavas occur in all three zones. Du Toit remarked on "the great proportion of brecciated igneous material in (the Pniel volcanics), which sometimes forms beds 25 to 50 feet in thickness". He was unable to find any vents or fissures from which the Pniel volcanics could have erupted, but he did note the presence of many veins or dykes of felsitic material, which he suggested marked the sites of fissures from which the higher-lying acid volcanics may have issued.

Kynaston (1907) described the presence of Ventersdorp sediments and diabase breccia in the area occupied by the Pretoria-Johannesburg Granite. The breccia extends south-westwards into the Bezuidenhout Valley, as described by Hatch and Corstorphine (1904).

Rogers (1907), working adjacent to that area surveyed by Du Toit (1906) reported the same lavas as had Du Toit, i.e. Pniel and Zoetlief Beds. He expressed doubts as to whether the Pniel lavas and overlying Black Reef Series ought to have been separated, as he found them to be "very intimately connected" in this area, shown by a certain degree of interbedding of the two series. It appeared to him that the great unconformity between the two groups in the Transvaal, did not seem to exist in the Vryburg area.

To the west of this region, Rogers (1908) encountered very little outcrop of the Ventersdorp or Vaal River System. Although the Koras rocks in this area differed in important respects from those of the Vaal River System, Rogers suggested that they may have been formed at about the same time, and therefore placed them as a distinct sub-group in the Vaal River System.

Du Toit (1908) reported the occurrence, in the Hopetown-Britstown-Prieska-Hay area, of the Zoetlief Series, consisting of an upper group of acid volcanics and a lower group of sediments with occasional lava flows. South of the T'Kuip Hills he found a possible volcanic neck. Succeeding the Zoetlief rhyolites with an inferred unconformity, is a group which du Toit called the Kuip Series. It comprises diabase and amygdaloid with some interbedded sediments, and is indistinguishable in hand specimen from those of the Pniel Series. It reached 457 metres in thickness in one area. Unconformably overlying this is the Pniel Series - a lower sedimentary group, mainly quartzite, and an upper volcanic one of "diabase amygdaloid" and breccia.

Humphrey (1909) found almost the whole of the country along the western Transvaal border to be occupied by rocks of the Ventersdorp System, i.e. mainly diabasic rocks with amygdaloids, porphyries and acid lavas with tuffs and breccias. He suggested that they all fell "within the limits of the Pniel Series as constituted in the Cape Colony", mapped by du Toit and Rogers in previous years.

Hatch and Corstorphine (1909), in their book on the geology of South Africa, mentioned sheets of amygdaloidal diabase which occur at different horizons in both the Lower and Upper Witwatersrand Beds, interbedded with slates and quartzites. They maintained that these "bursts of volcanic activity" were "a prelude ... to the violent volcanic outpourings which characterise a later period" i.e. the Ventersdorp lavas. They documented the Ventersdorp System in a fair amount of detail, consolidating all the information gathered until then. They described it as "a series of acid and basic lavas, tuffs, breccias, and conglomerates, which, being typically developed at Ventersdorp in the Western Transvaal, may be conveniently grouped together as the Ventersdorp System." They dismissed Molengraaff's name, the "Vaal River System", as ambiguous. They included within the Ventersdorp System the Elsburg Series, consisting of alternating conglomerates and quartzite. They wrote that "the basic lavas are well represented by the rocks of the Klip River Hills south of the Rand, where they have an aggregate thickness of some 5000 feet. The rock there exposed is a fine-grained amygdaloidal and porphyritic diabase, long known as the Klipriversberg amygdaloid. The amygdales consist chiefly of silica ..., calcite and chlorite."

Furthermore they stated that "acid lavas are of less frequent occurrence in the Transvaal, but they have been observed in the Makwassie Hills in the Wolmaranstad district", and at Platberg, and between Kimberley and Mafeking. "These form the lower division of the system, and, like the upper basic flows, are associated with tuffs and breccias, conglomerates and quartzites."

Hatch and Corstorphine wrote that, in the geological succession, the position of the fragmental rocks - breccias and conglomerates - had but recently been understood, as their connection with the "Klipriversberg amygdaloid" had not been recognised for a long time. At Ventersdorp clastic rocks of two distinct types are found: (1) volcanic breccias, grading into fine tuffs, and (2) boulder and pebble beds. The fact that the boulder beds are, to a large extent, made up of the debris of the Witwatersrand System points to a considerable denudation of the latter after its elevation above sea level and before the deposition of the Ventersdorp System.

They also noted "that the period to which the Ventersdorp System belongs was marked by great volcanic activity is shown in the enormous accumulation of lava, volcanic ash, and breccia. But this activity must have been to some extent localised, so that while lava and fragmental rocks were being ejected in one place, sand and gravel were simultaneously accumulating elsewhere: only in some such way can the great diversity of the beds grouped together as the Ventersdorp System be explained."

They mentioned the distribution of the Ventersdorp System in the Heidelberg district, in the Bezuidenhout Valley near Johannesburg, at Kromdraai north of Krugersdorp, in the Marico district, north of the East Rand, in the Venterskroon-Parys district, at Hoopstad, near Reitzburg, and at Ventersdorp. In the south-western area of its exposure, from Mafeking to the T'Kuip Hills, the Ventersdorp System had been grouped into two series, the Pniel and the underlying Zoetlief Series. They concluded by supporting the idea of Merensky (1907) that river diamonds were associated with the amygdaloids of the Ventersdorp System.

Rogers and Du Toit (1909) wrote a detailed account of the rocks of the Ventersdorp System in the Cape Colony. They included in their succession, as did Hatch and Corstorphine (1909), the Pniel Series, the Kuip Series and the Zoetlief Series. They also placed in this system the Koras Series, though admitting that its true position was not known.

In the Marico district of the western Transvaal Humphrey (1910a) found the following succession of Ventersdorp rocks: a basal sequence of shales and conglomerates, followed by acid lavas and porphyries. Overlying this is a second sequence of shales and conglomerates, succeeded by basic amygdaloid. He contended that the two sedimentary facies represent one period in the middle of which the acid rocks were erupted, as the whole succession appeared conformable. He also observed that the "close-grained diabase" which had previously been mapped as an intrusive, was simply "the central and lower portion of massive flows while the amygdaloidal rocks are those which consolidated more at the surface."

The Ventersdorp System in the Klip River Valley was found by Humphrey (1910b) to consist of the Elsburg Series overlain by the Klipriversberg amygdaloid. The latter forms the Klipriversberg to the north of the valley and the Suikerboschrand to the south-east. It consists of basic igneous rocks, frequently amygdaloidal, but also "close-grained." Occasionally it becomes markedly porphyritic with large well-developed crystals of felspar, which are usually saussuritised.

In their description of the geology around Zeerust and Mafeking, Hall and Humphrey (1910) presented a report virtually identical to that given by Humphrey (1909).

Horwood (1910) gave a detailed description of the petrography of the Klipriversberg Amygdaloidal Diabase. In addition he presented four major element analyses.

Mellor (1911) recorded the similarity of the amygdaloidal diabase in the Bezuidenhout Valley, Johannesburg, to that of the Klipriversberg, south of Johannesburg. He noted that it had been subjected to earth movements as shown by the intense shearing evident throughout the diabase mass.

Humphrey (1911) described the geology north of Zeerust, which appeared to be merely an extension of the rocks described by him in 1909 in the Marico district.

McDonald (1911), discussing the intrusive rocks of the Witwatersrand, noted the close resemblance of many of the dykes to the Klipriversberg volcanics: "The great similarity in mineralogical composition appears rather to indicate that the dykes are the hypabyssal equivalents of the lavas. The absence of volcanic necks, marking the centres from which the lavas could have flowed, indicates that the eruptions were of the fissure type, and the dykes may now fill the channels up which the mass of molten material flowed."

The Ventersdorp System in the Rustenburg district was described by Kynaston (1912) as consisting of two divisions, a lower sedimentary sequence of quartzites, quartzitic sandstones, grits and conglomerates, and an upper volcanic series. The lower portion of the latter

comprises fine-grained and frequently amygdaloidal basic rocks. These are overlain by acid lavas, showing characteristics such as banding (indicating flow structure), brecciation, nodular and spherulitic structures. However, after extending his work into the Marico district, Kynaston (1913) found that the volcanic rocks were intercalated with conglomerates, quartzites and shales, and there appeared to be no regular succession for the whole area.

A number of writers have recorded the fact that the Ventersdorp System underlies the Black Reef formation in some places with a well-marked unconformity, yet in others the relationship is apparently conformable (Sawyer (1903), Molengraaff (1903b), Hatch (1903b)). Horwood (1912a) believed that, although the Ventersdorp diabases had always been considered older than the Black Reef formation, at least some of the diabases are intrusive in the Black Reef formation. He supported this view by describing a number of occurrences in which diabase had forced its way up from below through vertical joints in the Black Reef quartzite and had spread out above it. He concluded that "it would appear, therefore, that the eruptive activity which characterised the Ventersdorp period did not entirely die out during the time that the Ventersdorp System was being laid down; but that it extended into that period to which the Black Reef Series belongs.". In a criticism of this paper, Mellor (1912) stated that Horwood's conclusions had been based on the misidentification of shale as shaly diabase, and that these conclusions were therefore erroneous. In reply to Mellor, Horwood (1912b) refuted Mellor's evidence and maintained his former stand.

Mellor (1913) observed that the Ventersdorp System in the western Witwatersrand was represented by the Klipriviersberg Diabase. Instead of forming hills, as it does south of Johannesburg, it tends here to form the lower ground. Mellor explained that this was a function of a greater ease of weathering, which was a result of the somewhat less acid character of the local rock, sometimes associated with a coarser texture. The diabase was found as usual to be markedly amygdaloidal and occasionally porphyritic, with its base apparently lying conformably on the Elsburg Beds, but overlain unconformably by the Black Reef Series. Mellor preferred to retain the Elsburg Beds in the Witwatersrand System, until further elucidation of its relationships showed otherwise.

Kynaston and Humphrey (1920) incorporated the reports of Kynaston (1912) and Kynaston (1913).

In 1921 Rogers (1921) discovered a succession of coarse agglomerate and lava belonging to the Jeppestown Series in the Heidelberg neighbourhood. The agglomerate was found to be considerably altered and the lavas, which appeared to overlie them, were highly vesicular. No feeder or vent from which the lavas might have come was found. Referring to the transition from the Witwatersrand to the Ventersdorp System, Rogers found indications "that volcanic activity began whilst the small pebble conglomerate and grit taken as the top of the Kimberley-Elsburg Series in Heidelberg was being formed". Rogers suggested that eventually "the arenaceous deposits were outweighed by the volcanic", and that "it would be difficult to find better evidence of the conformable succession of volcanic rocks on sedimentary than that offered by the Elsburg-Ventersdorp passage beds in that part of Heidelberg". This transitional zone is followed by at least 1524 metres of amygdaloidal lavas of the Ventersdorp System, containing a "strikingly porphyritic amygdaloidal lava about 300 feet above the base", similar to that found in Bezuidenhout Valley.

In a further description of the Heidelberg area Rogers (1922) stated that the Ventersdorp System there was divisible into two groups. The lower group, mainly lavas, displays stratification, indicating a large number of flows. Tuffs were seen only in the lowest rocks, and have undergone more intense alteration than the lavas. The phenocrysts of the porphyritic lava, also mentioned by Rogers (1921), were found to be albite-oligoclase feldspars. The upper group contains conglomerates and finer grained sedimentary rocks with some interbedded lavas.

In the Vredefort area the transition from the Elsburg Series to the Ventersdorp amygdaloid is very abrupt, unlike that near Heidelberg (Nel, 1927). In both these areas, the relationship between the two formations is conformable, unlike that in the Western Transvaal. The Ventersdorp Amygdaloid is the only representative of the Ventersdorp System in the Vredefort area. A porphyritic lava, the phenocrysts of which were found to be albite, is located near the base (Nel, 1927). This is similar to that described from the Heidelberg district by Rogers (1921). Nel did not notice tuffs or volcanic breccias in the

Vredefort area. The thickness of the amygdaloid was calculated to be between 3353 and 4267 metres, indicating "a close proximity to the place of origin of the lavas".

Rogers (1929) summarised the state of knowledge of the Ventersdorp System to that date. He listed the areas in which it occurs, i.e. in the central, west and south-western Transvaal, northern Orange Free State, and northern Cape Province. No Ventersdorp Beds had been recognised in either the Eastern Transvaal or Natal. Rogers stated that "the time represented by the Witwatersrand and Ventersdorp beds was one of considerable earth movements and widespread volcanic activity in the southern and western Transvaal and northern districts of the Cape; elsewhere in the Union there is no evidence yet known bearing on the period.". In neither the south-western Transvaal nor the Cape had rocks been recognised as Witwatersrand beds, but Rogers said that "the latter may be represented in the Cape Province by the lavas and sediments of the Zoetlief and Kuip series which are unconformably overlain by the Pniel amygdaloidal diabase.". To Rogers it seemed certain that the Zoetlief Series was represented by some part of the Witwatersrand-Ventersdorp succession in the southern Transvaal, but exact correlation could not yet be drawn. He also gave a good resumé of the different types of Ventersdorp rocks which were prominent in the various areas mentioned.

The Ventersdorp System in the Kimberley area was described by Williams (1932). Here only rocks belonging to the Pniel and Zoetlief Series were found, which lay "between the bottom of the Karroo system and the top of the Swaziland system of schists, gneiss and intrusive granite.". The Zoetlief beds are represented at the base by the "Vaal River conglomerate" which rests directly on top of the Swaziland schists with an unconformity, followed by quartzite and then quartz porphyry. This is succeeded unconformably by rocks of the Pniel Series, which consists of a quartzite with an underlying conglomerate, followed by a melaphyre, "a basic porphyritic rock of highly altered character", or diabase, as it was more commonly called. Williams presented chemical analyses of five diabase samples and three quartz porphyries.

Nel (1933) discussed rocks of the Witwatersrand System outside the Rand. He mentioned the volcanic rocks occurring at the base of the

Witwatersrand System in the Parys-Vredefort and the Klerksdorp-Ventersdorp areas. These appear to be overlain conformably by the Orange Grove quartzites, and consist in the former area of basic amygdaloidal lavas, and in the latter of "lavas and light-grey bedded ashy-looking rocks", some of which are cherty and amygdaloidal. In addition, "dense quartz-porphyry" is intercalated with the amygdaloidal varieties.

Nel mentioned the volcanic group discovered higher up in the Witwatersrand succession by Rogers (1921) in the Jeppestown beds. It is described as consisting of "agglomerate and lava of rather acid composition" in the Heidelberg area. It was also reported from south-west of Greylingstad, while on the West Rand he said it assumes an amygdaloidal character.

Nel (1934a) described the Witwatersrand System and associated rocks in the Klerksdorp and Ventersdorp districts. He noted the presence of a marked unconformity at the base of the Ventersdorp System, as the Ventersdorp Beds were deposited on a surface of older formations that had a fair amount of relief, and can be found resting on any of the various horizons of the Witwatersrand System, even extending right across on to the Old Granite. He reported that they lie nearly flat, or are gently folded. Nel maintained that "the relation of the Ventersdorp to the Witwatersrand System and the deposition of the former on an old land surface of marked relief, are factors of great importance in the search for those horizons known on the Rand in the concealed bodies of Witwatersrand beds around Klerksdorp."

Nel, in discussing the structure of the Witwatersrand, stated that "dislocations of the Witwatersrand System started in pre-Ventersdorp times, and were renewed at intervals until after the deposition of the Black Reef and Dolomite Series". These movements produced "a number of prominent N.E.-S.W. trending strike or oblique faults as well as dip-faults with a general M.W.-S.E. trend; and it seems very likely that the great Ventersdorp volcanic eruptions followed or were correlated with these fissures". He noted that "the general distribution of the volcanic agglomerates and boulder beds, rock types that could not have been formed far from the source of their constituents, along or close to some of these lines of fracture is

significant". He also concluded that these faults "are responsible for several reduplications of the rocks across a wide stretch of country", and "have all the characteristics of step-faulting with the downthrow sides north-west of the lines of fracture."

Twenty nine years after Molengraaff (1905) discovered the Dominion Reef beds and erroneously classified the associated lavas as Ventersdorp volcanics, Nel (1934a) carefully reviewed the information on hand. He realised that the two volcanic formations separating the Dominion Reef beds from the overlying Lower Witwatersrand beds formed a conformable succession with the Dominion Reef sedimentary groups and belonged to a sequence older than the Hospital Hill Series. He proposed the extension of "Molengraaff's term 'Dominion Reef Series' to embrace not only the quartzose sediments, but also the two volcanic groups following conformably above them, and to look upon the series as the lowest sub-division of the Witwatersrand System in the Klerksdorp-Ventersdorp area."

Nel (1934a,b) described the Dominion Reef volcanics as "lavas and bedded light-grey, ashy-looking rocks." Some of the lavas are "curious cherty amygdaloidal rocks" and "appear to be acid igneous rocks partly or wholly silicified"; others are "darker amygdaloidal lavas that bear a close resemblance to the Ventersdorp amygdaloid." Intercalated with the lavas is a "dense dark porphyry", with phenocrysts of mainly plagioclase and a little quartz. The thickness of the whole series varies from about a metre near Ventersdorp to over nine hundred and fifteen metres south-west of Hartebeestfontein. These accounts were followed by extremely detailed descriptions of the Dominion Reef Series, the Witwatersrand and Ventersdorp Systems in the Potchefstroom, Klerksdorp and Ventersdorp areas by Nel (1935), Nel et al. (1935), and Nel et al. (1939).

As a result of the discovery of a second volcanic horizon in the Witwatersrand System, the first being the Bird Amygdaloid in the Bird-Main Reef Series, Kovaloff (1934) suggested a reinterpretation of the boundary between the Upper and Lower Witwatersrand. The presence of the new horizon, the Jeppestown Amygdaloid, allowed for a better correlation of Witwatersrand Beds than had previously been achieved.

Beetz (1936) mentioned the presence of magnetic lavas in the lower part of the Ventersdorp System in the Klerksdorp District. The magnetic rock is "a dark fine-grained diabase with a high amount of magnetite", and the horizon is found "between the Ventersdorp sediments and the lavas and tuffs of the lowest part of the Ventersdorp System.". Beetz concluded that "where they occur at the surface in the Klerksdorp district, the underlying contact with the Witwatersrand beds can confidently be expected at a reasonable depth, say, 2,500 to 3,000 feet.". He stated that "the stratigraphical succession of rocks forming the Ventersdorp System ... in the Klerksdorp district ... is regular save that any one of the rock groups may increase or decrease in thickness or may be missing altogether in places.". He attributed this variation to a number of factors: the Ventersdorp had been deposited on an uneven surface; overlapping had been caused by the spread of later beds across the edges of older ones; lava flows and sheets of fragmental material thrown out from vents and fissures had built great fans with the thickness greater next to the eruptive centres; warping and fracturing had formed horsts and grabens.

Frommurze (1937) documented the results of a survey on the water-bearing properties of the Ventersdorp System and other important formations.

Truter and Strauss (1941), after studying the pre-Transvaal rocks in the Taungs area, proved that they belonged to two distinct formations, separated by an unconformity. The upper formation was correlated with the Pniel series of the Ventersdorp System, while the lower formation was correlated with the Zoetlief Series, which "can very possibly be correlated with the Dominion Reef series."

In a major consolidation of data Jacobsen (1943) described the stratigraphy and petrography of the Ventersdorp System in the southwestern Transvaal and northern Orange Free State in great detail. From boreholes drilled in the Buffelsdoorn area near Klerksdorp, he derived the following succession for the Ventersdorp System:

"Horizon Y : Dark green dense porphyrite with many small yellowish specks and occasional larger felspar phenocrysts.

- Horizon A : Quartz porphyrite.
- Horizon B : Green, dense to fine-grained porphyrite with small black specks.
- Horizon C : Large to fine volcanic breccia, dark quartzite, fine-grained tuffs and tuffaceous shales.
- Horizon C₁ : Green porphyrite with numerous feathery greenish-black specks.
- Horizon D : Mainly dense green-grey almost siliceous porphyrite. Characteristic are the numerous parallel white bands in the lower and middle portion of the lava flows.
- Horizon E : Dark green to grey-green porphyrite with zones of numerous yellowish specks and often with small black groups of crystals.
- Horizon F : Dark green porphyrite, very well distinguished by very large, white, stellar clusters or aggregates of plagioclase.
- Horizon G : Detritus-agglomerate, sporadically developed."

These horizons were grouped together in the following manner :

- Upper Ventersdorp System : Y
- Middle Ventersdorp System : A, B, C
- Lower Ventersdorp System : C₁, D, E, F, G

From his study Jacobsen showed that a "lava stratigraphy" can be evolved. He used this to discuss the Ventersdorp succession in the Klerksdorp area, the northern C.F.S., the Ventersdorp-Hartebeestfontein areas and the Wolmaransstad area. He rejected "the term Ventersdorp diabase, so firmly embedded into the South African literature" as incorrect, "as typical minerals of the diabase group are nowhere to be found. The absence of olivine especially is noted by all authors who describe the petrography of the Ventersdorp porphyrites.". He regarded the coarse Ventersdorp sediments to be of fanglomeratic origin, while the quartz porphyries he considered to be intrusive. He recognised the strong secondary alteration to which the Ventersdorp rocks have been

subjected, ascribing it to "autohydrothermal and solfataric processes". Jacobsen included the Elsburg Series in his study as many previous authors had considered it to be part of the Ventersdorp succession. He concluded, however, from his work, that it belonged to the Witwatersrand System.

Van Eeden (1946) described a series of rocks in the Schweitzer Reneke district which had up till then been grouped with the Ventersdorp System. He realised that there were actually two groups of rocks separated by an unconformity and maintained that the lower group, the Zoetlief series and the Dominion Reef series all represent rocks of the same age in different areas. The upper group he correlated with the Pniel or Ventersdorp lavas.

By examining borehole cores, Baines (1949) compared the geology of the Odendaalsrus area with that of the Klerksdorp area, and deduced that they had much in common.

In a major review of all volcanism in South Africa, Truter (1949) suggested giving the Dominion Reef rocks "the independent status of a system" because of the "phenomenal thicknesses" sometimes attained and their unconformable relationships to overlying rocks. Apart from the type area west and north-east of the Dominion Reef mine, Truter listed the following correlatives: north of Ottosdal; in the Vredefort area; the Insuzi series in the south-eastern Transvaal; the Godwan beds along the escarpment; between Denilton and Groblersdal; east and north-east of Crocodile Pools to beyond Thabazimbi; south of Lobatsi to between Zeerust and Mafeking.

He stated that the Witwatersrand System "was interrupted twice by volcanic eruptions, the first time during the Jeppestown epoch and subsequently during the Main-Bird epoch" and noted that the Jeppestown amygdaloid was first found in the Heidelberg area, and subsequently near Klerksdorp, near Randfontein and in the West Witwatersrand area, but is absent on the Central Rand and in the Vredefort area. Referring to the Bird amygdaloid, he noted its absence from the Klerksdorp area and on the Central Rand, but its presence throughout the East Rand, Far East Rand, Heidelberg and Vredefort areas. Also, the suggested correlatives of the Witwatersrand - the Mozaan series in Swaziland and

Natal, the Uitkyk formation north-east of Potgietersrust and the Stinkfontein series in the Richtersveld - all appear to contain similar volcanic zones.

He described the Zoetlief period as being characterised by frequent and extensive outpouring of lava, often on a stupendous scale" and reported that outcrops extend from Zoetlief, north of Vryburg, past Taungs and Kimberley to the T'Kuip hills near Britstown. North west of Prieska and near Upington they are referred to as the Koras series and in South West Africa as the Kunjas and Sinclair series, (correlations which are not supported today). Truter also included the quartz porphyries north-east of Kimberley and north and north-east of "Maquassi" in this system, as well as those east of Wolmaransstad and Cttosdal.

Truter stated that surface exposures and deep drill-holes indicated that the Ventersdorp System extended from the T'Kuip Hills in the south to Lobatsi in the north and Kuruman in the West to Leslie in the East. For the area south and east of the Rand and in the Cape Province, Truter inferred from the minimum of pyroclastic material contained in the lavas, that the eruptions were mainly of the effusive, non-explosive type and "were piled up more or less uninterruptedly in successive highly mobile flows building typical lava plains or plateaux.". Between Potchefstroom and Delareyville, however, breccia- and agglomerate-filled pipes or diatremes occur and the lavas are frequently interspersed with pyroclasts or volcanic conglomerate. According to Truter, this proved that "the eruptions were partly of the effusive and partly of the explosive type and that there were long intervals of quiescence during which active erosion in some localities proceeded simultaneously with rapid accumulation in others.". The distribution of diatremes and pyroclasts along normal faults invited the conclusion that "the eruptions were causally connected with tensional features and were therefore mainly of the fissure type.". He furthermore suggested that the frequently ripple-marked character of the sedimentary tuffs indicates subaqueous accumulation and that the breccias and agglomerates are partly vent-fillings and in part possibly of nuée ardente origin.

Brock (1950) believed that, in the absence of volcanic necks or throats, dykes constituted the feeders for the Ventersdorp "plateau

lava". He submitted that the constancy of markers of porphyritic lava, for instance, is evidence that the lava had been exceedingly liquid and had flowed for great distances horizontally. He declared: "The characteristic of plateau lava country is block faulting. This is a worldwide rule. Not only is there a potential void created by the vacating magma, but there is the additional weight of the lava on top of the crust. Gravity naturally is the major force concerned, and block faulting is the result. Examples of plateau lava country are the Deccan in India, Snake River in Washington-Oregon, etc. The structural aspect in such areas is the domination of block faulting. Lateral compression does not come into the picture."

Truter (1950), referring to the statement by Brock (1950) regarding the absence of volcanic necks, noted that "although such necks are not abundant, they are fairly common in certain parts, especially in the Western Transvaal, and many more are undoubtedly concealed under lava flows."

Von Backström (1952) presented a detailed description of the Dominion Reef Series between Wolmaransstad and Ottosdal. In addition he listed a number of chemical analyses, two andesites and one rhyolite from the Dominion Reef Series, and one quartz porphyry from the Ventersdorp System.

By 1953 the Zoetlief series had been tentatively correlated with the Dominion Reef series. Matthysen (1953) embarked on a project to accumulate evidence for or against this correlation, with reference to the Taungs-Christiana area. He divided the Ventersdorp System of Taungs into the Zoetlief and Pniel Series, while the Ventersdorp System of Christiana was correlated with the Pniel Series only. He found insufficient stratigraphic evidence to correlate the Zoetlief with the Dominion Reef Series.

He then embarked on a study of borehole cores in the Orange Free State from which Matthysen was able to divide the Ventersdorp System into three easily recognisable series :

Upper Ventersdorp or Pniel Series

Middle Ventersdorp or Zoetlief Series

Lower Ventersdorp or Klipriviersberg amygdaloid Series

He gave two reasons for placing the three series together as a system :

- (1) All three are characterised by large masses of lava, while the sediments almost always contain large quantities of ash-like material.
- (2) The system lies stratigraphically between two large lithologically different sedimentary systems.

From field and literature studies, it seemed to Matthysen that every known occurrence of Ventersdorp rocks could be accommodated in one of these divisions.

Matthysen then selected the Klerksdorp-Ventersdorp area as his prototype, as the Ventersdorp rocks attained their maximum development in this area. He drew attention to a number of features: a porphyritic andesitic marker bed in the Klerksdorp area; a time lapse between the deposition of the lowest lavas and the overlying pyroclastics; the existence on the Klerksdorp townlands of a younger group of andesitic lavas on top of the quartz porphyries, but indistinguishable in hand specimen from those below; a zone of large phenocrysts and the purple zone (in the Orange Free State), useful as markers in the Klipriviersberg amygdaloid, which he said occurred only in the Witwatersrand basin; and an unconformity below the Zoetlief Series everywhere except in the Klerksdorp and Le Roux Pan areas. He noted that, in hand specimen, rocks of the Pniel Series are very similar to those of the Klipriviersberg Series, but can, in certain instances, be distinguished by textures, size of amygdales and colour. He said the Zoetlief Series was folded, faulted and eroded in places before the deposition of the Pniel rocks, which lie disconformably on those of the Zoetlief Series.

Von Backström et al. (1953) documented the geology of the Lichtenburg area, the major part of which is underlain by rocks of the Ventersdorp System.

In the third edition of his book, "Geology of South Africa", published in 1954, Du Toit retained the Dominion Reef rocks as a series and relegated it to the base of the Witwatersrand System. He correlated these rocks with those in the Cape grouped as the Zoetlief series. In

the Jeppestown series of the Witwatersrand System, the thin Jeppestown Amygdaloid at about the middle of the series was mentioned as a marker of great value in deep boring.

Du Toit also gave a general description of the Ventersdorp System and listed the usual correlatives, but also included the "Deweras Series" of Southern Rhodesia. He emphasised the presence of volcanic breccias which are very characteristic of the lavas. Although the Ventersdorp magma had long been presumed to have been poured out from fissures, Du Toit noted that "breccia-filled necks penetrate Witwatersrand Beds to the west of Ventersdorp and Klerksdorp and between Greylingstad and Villiers, while sporadic shattering of the lavas, areas of coarse breccias enclosing in places large broken masses of Witwatersrand beds, and long belts of thick breccias north of Klerksdorp, all suggest extrusion from vents aligned along south-west trending fractures."

Although Holmes (1906) regarded the quartz porphyries of the Makwassie Hills as intrusive into the basic lavas, du Toit (1954), on the other hand, wrote that "just north of Makwassi Station the porphyry appears to be overlain unconformably by a quartzite body and that in turn by Ventersdorp lavas.". He postulated that the Ventersdorp Contact Reef forms the real stratigraphical base of the Ventersdorp System and that the relation to the Witwatersrand System may be a disconformable one in the Central and East Rand and Vredefort areas.

Another overview of the distribution, lithology and stratigraphy of the Dominion Reef, Witwatersrand and Ventersdorp Systems was published by Hamilton and Cooke (1954) in their book on South African geology.

The exploration of and opening of the Orange Free State Goldfield provided a wealth of data on the extension of the Ventersdorp System beneath the Karoo cover. Pienaar (1956), in a detailed stratigraphic and a preliminary structural study of the Ventersdorp System in the Orange Free State Goldfield, proposed the following succession:

Ventersdorp Upper Lava	- maximum	488 m
Ventersdorp Upper Sediments	- maximum	1676 m
Ventersdorp Lower Lava	- maximum	1829 m

All the flows of the Lower Lava (except in the northwest) dip towards the centre of the Witwatersrand basin, with distribution and thickness controlled by block faulting and erosion. He reported that the composition of the Lower Lava ranges from a diabase at the base to an augite andesite at the top. Pienaar also documented a number of marker zones occurring in this group - the upper and lower purple markers and the upper and lower altered zones. The Upper Lava is also andesitic but differs in certain respects from the Lower Lava. East of Odendaalsrus a conglomerate was found overlying the Ventersdorp Upper Lava. Pienaar realised that it may represent a still younger group, but described it with the Ventersdorp Upper Lava. Within these lavas he found three porphyritic zones which could be used as markers.

Nel and Jansen (1957) described the geology of the Vereeniging district. In this area the Dominion Reef System is restricted to one intercalation of lava between the Archaean Granite and the Orange Grove Quartzite. The Jeppestown Amygdaloid is well developed and described in detail, as is the Ventersdorp System.

In an interpretive review Visser (1957) described the Dominion Reef System in terms of a geosynclinal phase. He stated: "In the Transvaal, as well as in the northern Orange Free State and the northern Cape, the Dominion Reef System and its correlates, the Insuzi, Godwan, Uitkyk, Zoetlief and Koras Formations, are exposed along the margins of the great basins in which the Witwatersrand and the Transvaal Systems were deposited later. They rest with an unconformity on an uneven floor of Archaean granite and enclosed remnants of Archaean Formations. It would appear, therefore, that during the Proterozoic the old Archaean geosynclinal belt became active once more and a depression was created which was only slightly smaller than the Archaean basin. At first sediments were deposited, followed by the eruption of volcanic rock.". They "suffered only a slight degree of tilting prior to the deposition on them of younger formations.".

Referring to the orogenic and first magmatic phases of the Witwatersrand-Ventersdorp Cycle, Visser said: "Whilst the floor of the geosyncline was gradually being depressed, the marginal areas were tilted and elevated relative to the central position, whereby the size of the basin was decreased and erosion ... took place. In the

Klerksdorp and Odendaalsrus areas ... contemporaneous faulting also took place. ... eventually failure had to take place by way of tension-fractures, along which the Ventersdorp lava poured out to cover the last sediments in the basin conformably and to transgress across the tilted, older sediments along the margin of the basin."

A mineralogical aside was provided by Barbour, Bird and Gold (1957) who recorded the existence of axinite-bearing veins in Ventersdorp lavas from the Kinross area, Transvaal. They noted that, "apart from an occurrence in the greenstones of Southern Rhodesia, there is no published record of axinite from elsewhere in southern Africa."

In Botswana Wright (1958) tentatively correlated the Lobatsi Volcanic Series, a succession of felsites and porphyries with interbedded tuffs and agglomerates, lying between the intrusive Gaberones granite and younger sediments of the Ventersdorp and Transvaal Systems, with the Dominion Reef System in South Africa. Sediments of the Mogobane Series occur in the Gaberones area and although no amygdaloidal lavas occur here, evidence found outside of this area suggested a correlation with the Ventersdorp System.

Boocock (1960), however, reported that the Lobatsi Volcanic Series was composed of rocks of two different age groups, and assigned the volcanic succession, mainly felsites, to the Dominion Reef System. The younger sedimentary and volcanic succession, which includes the Mogobane Series, he correlated with the Ventersdorp System. He also included in the Ventersdorp System a succession of basic lavas, sediments and acid lavas in the Sequani area. This correlation was confirmed by Jones (1963) who noted that the Sequani lavas formed the north-westerly continuation of a similar succession which is developed in the Derdepoort area of the Transvaal. He also mentioned rocks of the Dominion Reef Series in the neighbourhood of Oodi Hill and along the Transvaal border.

In a description of the petrology and mineralogy of the rocks of the Dominion Reef System near Klerksdorp, Malan (1955) described mainly the lower sedimentary groups, but briefly mentioned the overlying lavas. He described two sequences of lavas, the "Lower Volcanics" separated by a quartzite from the overlying "Dominion Reef Lavas". He found that the Lower Volcanics are amygdaloidal, porphyritic, andesites, while the

Dominion Reef succession comprises basic lavas at the base followed by an upper acid lava sequence. Malan studied only the lower hundred feet of the former, which he found to be highly altered amygdaloidal andesite, with a thin basal tuffaceous horizon.

In a comprehensive report of the geology of the Orange Free State gold fields, Coetzee (1960) described the Jeppestown amygdaloid as being about 107 metres thick and built up of several lava flows of andesitic composition. He noted that the amygdaloids are sometimes surrounded by a white mantle and when the cavity filling is chloritic, the appearance resembles birds' eyes. A concentric banding observed in some core samples suggests that it might be a pillow lava. A rounding of detrital grains in the tuff also suggests deposition in water. The Jeppestown amygdaloid is regarded as a valuable marker in correlating the Witwatersrand sediments in the area around Odendaalsrus and Virginia.

Coetzee discussed the Ventersdorp System in terms of three subdivisions: the Lower Volcanic Stage, the Sedimentary Stage and the Upper Volcanic Stage. He described the Lower Volcanic Stage as a succession of superimposed flows ranging from 7.5 to 92 metres thick, which follow each other without significant hiatus. He reported at the top of some flows "a brecciated zone composed of swirled blocks of amygdaloidal lava", followed by "a highly amygdaloidal zone grading into an essentially non-amygdaloidal phase lower down, and finally a chilled basal sector containing scattered amygdaloids.". Alteration in the form of dolomitisation and epidotisation were also observed.

Of note in the Lower Volcanic stage is the presence of four markers, first described in detail by Pienaar (1956) - the upper and lower "altered zones", and the upper and lower purple bands. The altered zones are variolitic lava flows, while the purple bands have a mottled appearance and usually contain dark, rounded chlorite-bearing amygdaloids. Boreholes may contain two, one or none of the bands, which vary from a few centimetres to several metres in thickness and occur 6 to 9 metres below the top of a lava flow. Both Coetzee and Pienaar said the colour of the purple lava cannot be related to any particular feature in the chemical analyses, but in thin section finely-disseminated reddish-brown iron oxide is seen. He described a major

development of volcanic breccia, sometimes exceeding 200 metres in thickness, which has been found at both the top and the bottom of the Lower Volcanic Stage.

Coetzee described the Sedimentary Stage as consisting of "clastic sediments and, in places, detrital tuff". Where present, it is usually less than 305 metres thick, but can exceed 1067 metres and has been known to attain a thickness of 1981 metres. The tuff is black, aphanitic and not bedded and its chemical composition differs from that of the Lower Volcanic Stage.

The Upper Volcanic Stage reaches 488 metres in thickness in the area reported on by Coetzee, and in one part is overlain by later Ventersdorp sediment. One particular borehole displayed an interfingering of lava and sediment for almost 1830 metres. Like the Lower Volcanic Stage, this stage is comprised of augite andesite lavas, which also exhibit brecciated flow tops and amygdaloidal and non-amygdaloidal portions. It has, however, a darker green colour, porphyritic zones, epidotised bands and a coarser texture. Chemically it contains more normative quartz and the normative plagioclase is more albitic.

In 1961 the Economic Geology Research Unit at the University of the Witwatersrand decided to investigate the problem of correlation within the Ventersdorp System. Their motives for embarking on this study were very similar to those of the present author, and are quoted in Chapter 2. Annual progress reports were compiled by Winter (1961, 1962, 1963a, 1964 and 1965a), who consolidated the results of the investigation in his Ph.D thesis (1965b).

In 1962 a study of borehole data from the Klerksdorp-Bothaville area resulted in the compilation of a tentative table of stratigraphic subdivisions for this area. The various formation names were derived from localities in which well-developed outcrops are present or in which maximum thicknesses have been recorded in borehole cores. This table was amended in 1963, to include subdivisions of the Ventersdorp System in other localities. In the Vierfontein area a sequence of porphyritic and non-porphyritic lavas which could not be correlated with either the Makwassie or Rietgat Formation, was termed the Garfield Porphyry Facies. Features of note emerging from this compilation are

the fact that rocks of the Godwan Formation, north of Potgietersrus, were found not to resemble members of the Ventersdorp System, and the recording of algal structures in sediments of the Middle Division, by Winter (1963b). These stromatolites occurred in impure dolomitic limestone post-dating the quartz porphyry, and "were probably deposited in shallow lakes in an area of low relief after eruptions of acid and basic volcanic rocks." Winter concluded that "the discovery provides a link between the accepted prevalence of biogenic processes 2,000 million years ago during the time of deposition of the Transvaal System and the problematic origin of oolitic structures in the Swaziland System formed more than 3,000 million year ago." He suggested that such biogenic processes might have played a part in the deposition of "carbon" in the form of seams and granules in the Witwatersrand System.

In 1964 chemical data showed that distinctions could be made between the Jeppestown Amygdaloid, the Bird Amygdaloid and lavas of the Ventersdorp System. It was not certain whether similar variations would be found between the lavas within the Ventersdorp System.

A revised correlation of the Ventersdorp System was presented by Winter (1964). In this revision the Lower Division was associated with the Witwatersrand System and the Upper with the Transvaal System, while the Middle Division was renamed the Zoetlief System, comprising three series, the New Kameeldoorns, Makwassie and Rietgat Series. The Lower Division was renamed the Klipriviersberg Series and the Upper Division was correlated again with the Godwan Formation.

In 1965 a revised stratigraphic table was compiled. The three primary subdivisions of the "alleged" Ventersdorp System were found to be identifiable throughout the Western Transvaal and Northeastern Cape Province and investigations showed that the conclusions drawn from the type-area remain valid.

The findings of the EGRU investigation were consolidated by Winter (1965b) in his Ph.D. thesis, in which he documented in detail the structure and stratigraphy of the Ventersdorp System in the Bothaville district and adjoining areas. He pointed out that, although previous authors had noted many of the key facts in their attempts at stratigraphic subdivision, this information does not appear in any one

publication. With the aid of this previous literature and over two hundred borehole logs, Winter devised a stratigraphic column based on three primary subdivisions incorporating six formations, as shown in Table A.1. He selected the Bothaville area as his type area as it includes all the type boreholes, maximum thicknesses of most of the formations, and most of the geographical areas from which the formation names have been derived.

TABLE A.1 Stratigraphic column of the Ventersdorp System in the Bothaville type area (from Winter, 1965b)

Primary Subdivision	Formation	Representative Rocks
Upper	Allanridge (A)	Dark green amygdaloidal lavas
	Bothaville (B)	Quartzites, conglomerates, subordinate dark shales
Middle	Rietgat (R)	Green-grey amygdaloidal and porphyritic lavas; minor intercalated tuffaceous, calcareous and cherty shales; impure limestones; quartzites; conglomerates
	Makwassie (M)	
	New Kameeldoorns (N)	Quartz and felspar porphyries, green-grey amygdaloidal lavas, minor intercalated sediments
Lower	Langgeleven or Lower Volcanic (L)	Conglomerates; quartzites; tuffaceous, calcareous and cherty shales, impure calcareous rocks
		Green-grey pyroclastics and amygdaloidal lavas

Winter constructed isopach maps and structure contour maps of each formation. He presented detailed descriptions of all lithological types present. He stated that a number of beds within the Langgeleven Formation may be tuffs, but because all gradations exist between a well-defined tuff and a lava flow in this formation, it is difficult to distinguish between them, especially when the tuffs are amygdaloidal. He provided a list of ten distinctive features by which tuffs might be distinguished from lavas. He suggested that both agglomerates and flow-top breccias, which are very similar in appearance, occur in the Langgeleven Formation.

Winter described in detail the type sections for each formation. He considered the first volcanic effusion to be the beginning of the Ventersdorp System. Although a major change in lithology occurs,

Winter drew attention to Mellor's (1913) statement that there was no great time interval between the deposition of the last sediments of the Witwatersrand System and the first effusions of the Ventersdorp System, and that there are areas in which volcanic material and sediments are intimately admixed in a transitional zone. Winter suggested a possible origin for the purple bands as "manifestations of contemporaneous weathering acting variously on different kinds of lava". He noted their prevalence on updomed areas, suggesting that in basins the flows followed upon one another at shorter intervals, thus preventing most flows from being weathered. Winter subdivided the Langgeleven Formation into five zones characterised by the following features:

5. green chalcedony amygdales and clustered milky quartz amygdales
4. variolitic bands or "altered zones"
3. purple bands
2. porphyritic lava
1. talcose tuffs

The New Kameeldoorns Formation overlies the Langgeleven Formation with an extensive unconformity and its distribution bears no relation to the depositional basin of the Langgeleven Formation. The sequence consists of clastic wedges of immature sediments which accumulated in fault troughs when the bounding faults were active.

The Makwassie Formation overlies the New Kameeldoorns Formation without a break and consists of quartz porphyries with quartz-free porphyritic and non-porphyritic volcanic rocks and minor bodies of sediments. The lower limit of the formation was taken as the base of the first lava flow following upon New Kameeldoorns sediments. The upper limit is marked by the advent of quartz-free porphyritic and non-porphyritic lavas. Although not conclusively proved, Winter favoured an ash-flow mechanism of emplacement of the quartz porphyries. He introduced a facies designation - the Garfield Formation - to include one or both of the Makwassie and Rietgat Formations in areas where these two cannot be separated or where doubt exists as to which of the two is represented.

The Rietgat Formation consists predominantly of green-grey amygdaloidal porphyritic and non-porphyritic lavas, but contains subordinate sediments and pyroclastic rocks. Algal-bearing limestones occur in this formation. Winter claimed that the volcanic rocks of this formation cannot be distinguished from the basic lavas of the Makwassie Formation. The uppermost flows may be intensely altered to calcite, epidote, clayey minerals and iron oxides, an important criterion for distinguishing between Rietgat and Allanridge lavas.

The Bothaville Formation consists of sediments suggestive of a shelf facies - conglomerates, grading into quartzite and shale, and then again to conglomerate. The sediments are of a subgreywacke composition and define an unconformity.

The Allanridge Formation lies conformably upon the uppermost conglomerate of the Bothaville Formation. It consists exclusively of alternating porphyritic and non-porphyritic lava flows.

With reference to the geological history of the Ventersdorp rocks, Winter concluded that the volcanics of the Langgeleeven Formation were extruded in rapid sequence as "fissure eruptions or at least lines of volcanoes along deep-seated fractures.". "The peculiar features of the purple bands and altered zones suggest subaerial effusion, as also does the association of purple bands with abundant agglomerates. No pillow structures were noticed.".

The Makwassie Formation covers a large area and its great thickness suggested to Winter a similarity to some of the rhyolitic provinces such as the Taupo-Rotorua district, New Zealand, and the Basin and Range Province of the United States of America. Winter submitted that the Makwassie volcanics exhibit characteristics much in common with these.

The close lithological and structural relationship between the Rietgat and Makwassie Formation suggested to Winter that the former could represent the waning phase of the Makwassie cycle of volcanism. Sedimentary beds are more numerous and the nature of the flows and the oxidation of their tops suggest a subaerial environment for both formations. Winter assumed that the acid Makwassie magma, probably

formed during New Kameeldoorns times by anatexis or differentiation, had become depleted by Rietgat times, the younger magma coming from a source more common to that of the Langgeleven lavas.

The Allanridge Formation consists of andesitic plateau lavas, subaerially extruded and uninterrupted by sediments. Suspected feeder dykes have been found.

Winter stated that the three primary subdivisions of the Ventersdorp System behave as distinct structural entities. The Lower Ventersdorp System conforms structurally with that of the Upper Division of the Witwatersrand System, while the structure of the Middle Ventersdorp System is less complicated and deformation has occurred along different axes from those of the underlying sequences. There is also a difference in degree of warping between the Upper and Middle Ventersdorp System, leading to angular unconformities in elevated areas. Winter thus did not regard the Ventersdorp System as a unit. He pointed out that, lithologically, the only uniform factor within the System is the abundance of andesitic lavas in all three subdivisions. He therefore suggested that the Lower Ventersdorp System be included with the Upper Witwatersrand System as a major subdivision, to be named the Klipriviersberg Series. Similarly the Upper Ventersdorp System could be considered as the lowermost subdivision of the Transvaal System and could be called the Pniel Series. He proposed that the Middle Ventersdorp System is of sufficient importance to warrant classification as a separate system. He recommended that the term "Ventersdorp System" should be discarded to avoid confusion, and the name "Zoetlief System" be adopted.

A.B. Simpson was responsible for documenting the petrology and geochemistry of these rocks, but his results have unfortunately never been published, although a bibliography of the geology of the Ventersdorp rocks was (Simpson, 1964).

Nel and Verster (1962) described the geology of the area between Bothaville and Vredefort. They found the Dominion Reef System to form a narrow band between the Archaean rocks and those of the Witwatersrand System to the west of Vredefort. The Dominion Reef System consists of altered basic amygdaloidal lava with lenses of metamorphosed

sediments. They noted the presence of the Jeppestown amygdaloid in this area, where it is about 90 metres thick and capped by a volcanic breccia. They found that much of the area is underlain by Ventersdorp rocks, which may reach 3050 metres in thickness. Quartz-feldspar porphyry occurs on different horizons. Near the base is a zone of large albite phenocrysts. Pyroclastics and tuffs are also present, as well as lenses of sediment.

The geology of the Ottosdal area in the Transvaal was described by Von Backström (1962), who found that rocks of the Dominion Reef System cover relatively large parts of this area and consist of a basal auriferous and uraniferous conglomerate with quartzite, followed first by andesitic lava which contains lenses of porphyry, then by rhyolite porphyry containing tuffaceous material. Microscope examination and the presence of cordierite indicates that these rocks have been subjected to intense metamorphism. Von Backström noted that the major part of the Ottosdal area is covered by Ventersdorp rocks, most of which are andesitic, with some porphyritic acid lavas. Chemical analyses are presented for most of the rock types.

An interpretation of a number of age measurements on Dominion Reef, Witwatersrand and Ventersdorp rocks and their supposed correlates was presented by Nicolaysen (1962). Age determinations on uraninite and monazite from the basal conglomerate of the Dominion Reef System resulted in an age of 3050 ± 100 my. Rb-Sr measurements on muscovite from the underlying basement granite indicate an age of 2900 ± 100 my. The monazite was known to be detrital and by inference, so was the uraninite, thus a maximum age of 3050 ± 100 my. for the Dominion Reef System was obtained. A tentative age of 2100 my. was assigned to the deposition of either the Witwatersrand or Ventersdorp rocks. Pegmatites underlying the Stinkfontein Series were dated at 950 my. Correlation of this series with the Witwatersrand System was thus no longer tenable. Similarly, the rocks underlying the Koras Series were dated at 950 my., invalidating their correlation with the Dominion Reef System.

These relations were later supported by Van Niekerk and Burger (1964) who dated zircon crystals from a quartz porphyry in the Ventersdorp System by the U-Pb isotope method. They obtained an age of 2300 ± 100 my for the lava.

Haughton (1963) noted that the Insuzi Series, the Godwan Formation and the Zoetlief Series were still correlated with the Dominion Reef System, and the Mozaan Series with the Witwatersrand System, while in the same year Winter (1963a) stated categorically that "there is no evidence in the area for believing that the former Pongola System (Mozaan plus Insuzi Series) can be correlated either with the Dominion Reef or Witwatersrand System."

Pretorius (1963) gave a brief description of Ventersdorp rocks in the Central Rand. He also provided a number of chemical analyses of igneous rocks.

The geology of the Schweitzer-Reneke area was described by Van Eeden et al. (1963). In this account the series of quartz- and feldspar-porphyrries and tuffs found there were correlated with the Zoetlief Series and the Dominion Reef System, all of which Van Eeden et al. stressed belonged to the same period of volcanicity. They furthermore found that the greater portion of the area is occupied by the rocks of the Ventersdorp System, consisting mainly of andesitic lava, but also more acidic rocks, with intercalated pyroclastic and clastic sediments, volcanic breccia and agglomerate and tuff. Boreholes revealed clearly the measure of variation in this system in respect of stratigraphic succession and thickness of the components, within comparatively short distances.

Describing the Ventersdorp System in the Loraine area Olivier (1965) indicated a preference for Winter's (1962) subdivision into a Lower, Middle and Upper Stage.

In a critical review of stratigraphic terminology in South Africa, Truswell (1967) commended Winter's (1965b) nomenclature of the six stratigraphic units of the Ventersdorp System, but felt that the use of "Lower Volcanics" as a formational name was confusing, preferring the term "Klipriviersberg Series".

Van Niekerk and Burger (1968) obtained an age of 2500-2700 my from Pb-isotope dating of acid lavas from the Zoetlief volcanics. The mean age was 2638 ± 105 my. In the following year an age of 2800 ± 60 my was obtained for Dominion Reef lavas (Van Niekerk and Burger, 1969). They

suggested that the Zoetlief succession might be more appropriately accommodated in a separate stratigraphic system.

Haughton (1969) used the term "Witwatersrand Triad" as a chapter heading in his book on the geological history of southern Africa, in which he produced a very thorough literature review.

Whiteside (1970) described the extrusive rocks of the Witwatersrand triad, i.e. the Dominion Reef System, the Witwatersrand System and the Ventersdorp System, which comprise the three lowest members of the Proterozoic in South Africa. He listed the succession for the Dominion Reef System in the type locality, near Hartbeesfontein, as follows :

(a)	Upper or Acid Volcanic Formation	6700 ft
(b)	Lower or Basic Volcanic Formation	2100 ft
(c)	Basal Sedimentary Formation	120 ft

He described these rock types, as well as those from presumed correlatives. The Bird Amygdaloidal Diabase and the Jeppestown Amygdaloid are the only two lavas known within the Witwatersrand System and these are briefly described. Whiteside together with Strydom (1968, in Whiteside, 1970) suggested a modification of the Ventersdorp succession as proposed by Winter (1965b). This included the addition of the Vaal Bend Formation below the Langgeleven Formation and the relegation of the New Kameeldoorns Formation to a sub-stage of the Langgeleven.

Crockett (1971) described rocks in the Lobatse and Ramotswa areas of Botswana which he correlated with the Ventersdorp System. The succession is divided into three units: the Lower Volcanic Assemblage of Si- and K-rich igneous rocks, the Mogobane Assemblage of argillaceous sedimentary rocks, and the Upper Volcanic Assemblage of andesitic lavas. Crockett discussed the petrogenesis of the Lower Volcanic Assemblage in detail, and presented eight chemical analyses. In the past these rocks have been assigned to both the Dominion Reef and Ventersdorp Systems. A porphyry occurring in the Lower Volcanic Assemblage has been dated at 2215 ± 100 my, which corresponds well with the age of 2300 ± 100 my obtained from a quartz porphyry in the Ventersdorp System by Van Niekerk and Burger (1964).

A major event in the development of a stratigraphic nomenclature of the Witwatersrand triad lavas occurred in 1971 when the South African Committee for Stratigraphy decided to follow international recommendations and to adopt a lithostratigraphic instead of a chronostratigraphic system of nomenclature.

Labuschagne (1974) made a study of the lower 850m of the Ventersdorp Group on the East Driefontein Mine east of Carletonville, where only the Klipriviersberg Formation is present. He distinguished two types of flows, aphanitic and phaneritic. A significant component of the aphanitic type are "petrabullae", large infilled cavities up to several metres in diameter. Pillow structures have been found in places and the occurrence of brecciated lava layers capping aphanitic lava flows was noted. From the mineral assemblages present, Labuschagne deduced that they represented the quartz-albite-epidote-biotite sub-facies of the greenschist facies of metamorphism. The dense, compact nature of the aphanitic rocks resulted in their being less markedly metamorphosed than the phaneritic lava. Labuschagne realised that petrographic classification of these rocks is virtually impossible because of their altered and fine-grained nature. He therefore applied various classifications based on their chemistry and settled on the name "basaltic trachyandesites". He offered an informal stratigraphic sub-division, comprising four units, based on lithology, petrology and chemistry. Labuschagne concluded that the rocks belong to the calc-alkaline or Pacific-suite type.

Visser et al. (1975) compiled the results of a large number of theses which were initiated for the purpose of elucidating the Ventersdorp rocks and their relationships in the northern Cape Province. They found that rocks correlated with the New Kameeldoorns, Makwassie, Rietgat, Bothaville and Allanridge Formations occurred. The equivalents of the Langgeleeven Formation did not outcrop anywhere, giving credence to Winter's (1965b) assertion that this formation is restricted to the area of the Upper Witwatersrand Basin. They deduced from field relationships and age dating that the porphyries of the Zoetlief Formation in the Northern Cape were the equivalent of those of the Makwassie Formation of Winter (1965b). They furthermore concluded that "the stratigraphic nomenclature as proposed by Winter (1965b) for the area around Bothaville can be successfully adopted in the Northern Cape."

Also working in the northern Cape, Grobler and Emslie (1975) found that the volcanic-sedimentary sequence of rocks at the T'Kuip Hills corresponded with the middle division of the Ventersdorp System in the Bothaville area, viz. the New Kameeldoorns, Makwassie and Rietgat Formations, and that these are followed by andesitic flows corresponding to the Allanridge Formation. They concluded that the Zoetlief Series, the Kuip Series and the Pniel Series are all part of the same cycle and should therefore be grouped with the "Ventersdorp Group".

The widespread occurrence of stromatolites in primary limestones of the T'Kuip Hills was described by Grobler and Emslie (1976). The rocks in this area correspond to the New Kameeldoorns, Makwassie and Rietgat Formations and the stromatolitic beds form part of the latter. An isotopic age of 2310 my (error and method not stated) for a quartz porphyry in the acid lava zone confirms this correlation with the Ventersdorp Supergroup.

Wyatt (1976) investigated the nature of the Klipriviersberg volcanics south of Johannesburg from a stratigraphic, morphological and petrochemical point of view. Apart from a minor lithic tuff at the base, he found all the volcanics to be lavas, which overlie the Elsburg Formation with apparent conformity. He subdivided them into the Westonaria, Alberton, Orkney and Kibler Park Formations, based on the presence of distinctive lithologic units, including a porphyritic zone, amygdaloidal and breccia horizons. Wyatt concluded that the latter represent fragmented lava flow tops. This, together with the identification of pahoehoe flows, but no pillow lavas, suggested that the lavas were subaerially extruded. Wyatt analysed forty nine samples, to which he applied cluster analysis in order to group chemically similar samples together. He classified the three upper formations as alkali-rich oceanic tholeiites, the result of a partial melt of a wet Precambrian mantle, while a genetically unrelated, 100-metre thick MgO-rich basal unit, the Meredale Member, was possibly formed as a result of liquid immiscibility.

Winter (1976) updated a portion of his thesis of 1965, proposing a framework of lithostratigraphic subdivision of the Ventersdorp for consideration by SACS. He presented a review of some of the more

recent literature on stratigraphic aspects. A subdivision founded on six formations was evolved. Winter described briefly the type stratigraphy of the Ventersdorp succession as elaborated upon in his thesis earlier (1965b), but made a number of changes. He shortened the name "New Kameeldoorns" to "Kameeldoorns". The Klipriviersberg Formation was elevated to group status and divided into six formations. He classified the group as "almost exclusively andesitic". Table A.2 shows the new subdivision. Of note is the fact

TABLE A.2 Subdivision of the Klipriviersberg Group (from Winter, 1976)

Formations	Marker beds	Distinctive Characteristics	Contacts in type borehole, Langgeleven No. 1	Thickness	Correlative Areas (List Incomplete)	Pienaar's (1965) Classification
Edenville		Green chalcedony and clusters of milky quartz amygdales. 1.	1 827 2 395	568	Allanridge, Klerksdorp, S. East Rand	Zone U
Loraine	AZ 1 AZ 2	Variolitic and spherulitic structures surrounded by whitish groundmass hence "altered zones"	AZ 1 flow 2 395-2 401 AZ 2 flows 2 492-2 535 2 612	217	O.F.S., Bothaville, Klerksdorp, East Rand	Zone M
Jeannette Agglomerate		Prominent thick agglomerate at top. Many agglomerate beds in contrast with above. 2.	2 709	97	O.F.S., Bothaville East Rand	
Orkney	PZ 1 PZ 2	Purple-topped ash-flow tuffs. 3. 4.	PZ 2 flow 2 945-2 965 3 189	480	O.F.S., Klerksdorp, Bothaville, Edenville	Zone L
Alberton Porphyry	Porphyritic Marker	Porphyritic lavas including "porphyritic marker". 6.	Porphyritic marker flow 3 352-3 360 3 360	171	South Vaal, Ventersburg (?), Edenville, West Rand, Klipriviersberg, East Rand, Vrededorf, Evander, Greylingstad	
Westonaria		Dark green "talcose tuffs"	Absent		South Vaal, Westonaria Alberton, S.E. Rand, Greylingstad	Absent

1. Some variolitic zones noted by Winter [1965] from Allanridge area.
2. Some porphyritic lavas south of Orkney
3. Pienaar distinguishes PZ 1 and PZ 2 in Klerksdorp and O.F.S. areas from other purple tops by their round chlorite amygdales.
4. Pienaar discovered a zone of pink chalcedony amygdales in O.F.S. goldfield in this stratigraphic position.
5. At base, Pienaar recognised a flow containing characteristic almond-shaped amygdales.
6. The "Porphyritic Marker" contains phenocrysts of felspar longer than 50 mm and thicker than 7 mm and consists of 1-4 flows of lava.

that the subdivision follows that presented in Winter's thesis (1965b), with the subsequent insertion of an agglomerate zone between the purple bands and the altered zones. Also mentioned is a new formation - the Vaal Bend Formation - proposed by Strydom for a succession of sediments, acid and basic lavas underlying the Klipriviersberg Group west and southwest of Klerksdorp, and resting on Witwatersrand rocks with a major unconformity. Winter could not fit this formation into his stratigraphic scheme due to "paucity of facts". The former Garfield Formation is now called the Garfield Member and refers to a group of pale green-grey porphyritic lavas, distinguished only from

some Makwassie quartz porphyries because they are not acid. The member may contain minor intercalations of sediments and non-porphyritic lavas and may be a member of either the Makwassie or Rietgat formations, as Winter states that "part or all of the Rietgat Formation can be contemporaneous with the Makwassie.". He noted that many occurrences previously allocated to the Zoetlief Series of the Ventersdorp System have subsequently been designated as Dominion Reef System. He therefore proposed that the Zoetlief Series, i.e. the Kameeldoorns, Makwassie and Rietgat Formations, be renamed the Platberg Group. The former Pniel Series was relegated to the informal term "Pniel Sequence", as Winter says there is no common lithology to bind the two formations covered by that term, the Bothaville and Allanridge. He advocated the abolition of the name "Ventersdorp Supergroup" because it obscures important structural relationships.

Potgieter and Lock (1978) described and named a new lithostratigraphic unit, the Ritchie Quartz Porphyry Formation, south of Kimberley. Originally mapped as equivalents of the Zoetlief Series by du Toit (1906), the quartz porphyry was found by these authors to grade upwards into a brecciated zone, the voids of which have been infilled by the overlying andesite of the Ventersdorp System. The authors correlated the andesite with the Allanridge Formation and suggested that the quartz porphyry was the equivalent of the Makwassie Quartz Porphyry Formation.

The Godwan Formation in the Eastern Transvaal had been correlated in the past with the Insuzi Series of the Pongola System, thought to be a Dominion Reef equivalent. Button (1978) reviewed its stratigraphical relationships and trends and suggested that it might be the easternmost exposure of the Upper Witwatersrand basin.

Van Niekerk and Burger (1978), with improved processing and analytical techniques, obtained a new age of 2643 ± 80 my for the Makwassie Quartz Porphyry Formation. This age supported Winter's (1976) correlation of the Makwassie with the acid lavas of the Zoetlief succession.

A petrological study of volcanic rocks at T'Kuip by Cornell (1978) showed that they had undergone burial metamorphism resulting in a mineral assemblage corresponding to the greenschist facies. This

metamorphic and metasomatic event was dated at about 2000 my ago, and found to occur in a number of other Ventersdorp rocks in other areas. Cornell concluded that "future petrological and geochronological studies on Transvaal and pre-Transvaal formations should seek to evaluate the effects of metamorphism and associated metasomatism before postulating unusual magma compositions or rearranging stratigraphic correlations."

Tyler (1979a) made a study of the Kanye Volcanic Group in the west-central Transvaal. It had originally been correlated with the Ventersdorp System and later with the Dominion Reef System. It has since been determined that rocks of this Group are older than 2750 my, which is the age of the intrusion of the Gaborone Granite into the volcanics. The group consists of a succession of felsites ranging compositionally from dacite to rhyolite. Variolitic structures similar to those seen in the basaltic komatiitic lavas of the Onverwacht Group have been observed. Evidence suggests that a distinct hiatus, during which the Kanye Volcanics weathered to form a palaeosoil unit, separated the consolidation of the felsites from the extrusion of the Ventersdorp lavas. Tyler suggested that the Kanye Volcanic Group constitutes the proto-basinal phase of deposition of the Ventersdorp Supergroup and therefore should be correlated with the Ventersdorp Supergroup.

East of the Botswana border, near Gaborone, is a sequence of sediments, basalts and rhyolites which have been correlated with the Ventersdorp Supergroup by Tyler (1979b). The succession is divided into three belts, the Tshwene-Tshwene, Seokongwana and Derdepoort belts, in which almost the full Ventersdorp succession is preserved. There is a basal, basaltic zone (correlated with the Klipriviersberg Group), a thick chaotic alluvial fan deposit (Kameeldoorns Formation), a zone of acid volcanics (Makwassie Quartz Porphyry Formation), and an upper meandering-stream-flood-plain sedimentary zone (Bothaville Formation). Tyler found that the basal volcanics have a major element geochemistry consistent with continental tholeiitic basalts, while the acid lavas correspond to rhyolites.

Watchorn (1980) published a review of sedimentation and volcanism in the Dominion Group. Although numerous attempts have been made to

correlate the Dominion Group with other volcanic and volcano-sedimentary units in South Africa, present evidence implies that its development was restricted to the Western Transvaal, where it represents a precursor to the main Witwatersrand basin (Watchorn, *ibid.*). Stratigraphically, the group is divided into three formations :

- (1) The Rhenosterspruit Formation at the base composed mainly of coarse arkosic sediments, lenticular grits and conglomerates;
- (2) The Rhenosterhoek Andesite Formation consisting of intermediate lavas and tuffs; and
- (3) The Syferfontein Porphyry Formation, consisting of porphyritic felsic lavas, with interbedded tuffs and cherty amygdaloidal volcanics.

Palaeocurrent data indicate that the basal sediments prograded in a westerly direction and a braided fluvial regime is postulated for their origin. The increased proportion of lenticular volcanic units towards the top of the Rhenosterspruit Formation indicates a period of coeval volcanism and sedimentation.

Because of their close spatial, structural and temporal association, the Dominion Reef, Witwatersrand and Ventersdorp Systems were grouped together conveniently as the "Witwatersrand Triad" (Hamilton and Cocke, 1960). The South African Committee for Stratigraphy (SACS)(1980) accepted the usefulness of this concept and termed the time span involved, thought to be some 240 Ma, the Randian Erathem. They recommended, however, that the word "triad" should not be capitalised, as the grouping of the three units is quite informal, because there is no overall unifying lithological association and, furthermore, such a large grouping of successions goes far beyond the scope of purely lithostratigraphic classification and nomenclature.

SACS (1980) changed the Dominion succession to the rank of group. From recent age determinations and more detailed geological mapping, it was concluded that development of the Dominion Group was restricted to the Western Transvaal between Ottosdal in the west and the Vredefort Dome in the east, as shown in Figure 1.3.

The following former correlates were assigned by SACS to other sequences:

- (1) The Insuzi Series (now Msuze Group) is part of the Pongola Sequence.
- (2) The Godwan Formation is part of the Transvaal Sequence.
- (3) The Uitkyk Formation is part of the Pietersburg Group.
- (4) The succession between Groblersdal and Dennilton has been named the Groblersdal Group, as it has been found to be appreciably younger than the Dominion Group.
- (5) The volcanics and associated rocks forming the strip running west from around Thabazimbi have been assigned provisionally to the Ventersdorp Supergroup.
- (6) The so-called pre-Ventersdorp lavas between Mafeking and the Botswana border are probably Ventersdorp correlates.

A lithostratigraphic column of the Dominion Group, amended from SACS (1980), is shown in Figure 1.6. The group comprises three formations: the Rhenosterspruit Quartzite Formation at the base, succeeded conformably by the Rhenosterhoek Andesite Formation, which in turn is overlain by the Syferfontein Porphyry Formation. SACS (1980) has mistakenly called the two lower formations the Rhenosterhoek Quartzite Formation and the Rhenosterspruit Andesite Formation. In the type area around the old Dominion Mine, the group is 2250m thick. A probable age of approximately 2 800 Ma has been determined for the Dominion Group (Van Niekerk and Burger, 1965).

The new name given by SACS to the former Jeppestown Amygdaloid or Jeppestown Lava is the Crown Formation. It occurs in both the Central Rand and Klerksdorp areas in about the middle of the Jeppestown Subgroup, which is the uppermost of three subgroups within the West Rand Group (formerly the Lower Division of the Witwatersrand Supergroup). The distribution of the Witwatersrand Supergroup is shown in Figure 1.4. There are no radiometric age dates for the lavas of the Witwatersrand Supergroup.

SACS maintained the designation of the Ventersdorp succession as a Supergroup, consisting of a lower Klipriviersberg Group, a middle Platberg Group, and two formations, the Bothaville and the Allanridge Formations. As these two formations have no unifying lithological features, the former being sedimentary and the latter volcanic, they cannot be united as a group.

The distribution of outcrops of the Ventersdorp Supergroup is shown in Figure 1.5. Drilling has shown, however, that Ventersdorp rocks occur under cover as far south as a line from Bethal, past Virginia towards Hopetown. The Ventersdorp repositories were thus larger than that of the Witwatersrand Supergroup, though superimposed upon the latter. Maximum thicknesses intersected in the Bothaville area combined amount to about 5 100m of volcanics and 2 900m of sediments. In the northern Cape Province the thickness is much less, as the basal Klipriviersberg Group is not known to occur outside the Witwatersrand basin.

The SACS (1980) subdivision of the Klipriviersberg Group follows that of Winter (1976) as shown in Table A.2. At the base of this group the Ventersdorp Conglomerate Formation, which contains the Ventersdorp contact reef, is not included in either contiguous Supergroup. The subdivision of the Platberg Group into the Kameeldoorns, the Makwassie Quartz Porphyry and the Rietgat Formations, basically also follows that of Winter (1965b, 1976), with the porphyritic Garfield Member being a part of either of the latter two formations. Overlying the Platberg Group, the Bothaville Formation is followed by the Allanridge Andesite Formation. The Vaal Bend formation of Strydom and Whiteside (Whiteside, 1970) was retained as an informal unit, as its relationship to contiguous units is still uncertain.

McIver et al. (1982) proposed that the Ventersdorp Supergroup represents the youngest komatiitic sequence in South Africa. Their proposal is based on the presence of high MgO rocks at the base of the Klipriviersberg Group south of Johannesburg, described by Wyatt (1976). Geochemical data compare favourably with known komatiites, but TiO₂ values in the Ventersdorp rocks are higher. To explain this McIver et al. suggested that, if two liquids were produced, one at higher pressure by smaller degrees of melting than the other, the MgO contents could be comparable but the TiO₂ content of the higher pressure melt would be higher.

Grobler et al. (1982) found they were able to distinguish between lavas of the Allanridge and Rietgat Formations in the Northern Cape and Western Transvaal using visual criteria if a relatively complete succession is present, or by using major element values if not.

APPENDIX II - Analytical Procedures

element	kV	mA	counter	vacuum	crystal	collimator tube	STANDARDS USED							
Mn	50	40	flow	on	LiF200	c Cr	AGV-1	BCR-1	G-2	GSP-1	JG-1	NIM-N		
Fe	"	"	"	"	"	f "	"	"	"	"	"	"		
Ti	"	"	"	"	"	" "	"	"	"	"	"	"		
Ca	"	"	"	"	"	" "	"	"	"	"	"	"		
K	"	"	"	"	"	" "	"	"	"	"	"	"		
Si	"	"	"	"	PET	c "	"	"	"	"	"	"		
Al	"	"	"	"	"	" "	"	"	"	"	"	"		
Mg	"	"	"	"	TLAP	f "	"	"	"	"	"	"		
P	"	"	"	"	Ge	c "	"	"	"	"	"	"		
<hr/>														
Ba	55	40	flow	on	LiF220	f Cr	AGV-1	BCR-1	CAR-08	GSP-1	KRF-13	NIM-P	PRO-1	S-12
Nb	"	"	scint	off	"	" W	AGV-1	CAR-08	G-2	KRF-13	PRO-1	S-9	S-12	S-15
Zr	}													
Y	}													
Sr	}													
Rb	}													
Zn	"	"	flow	on	"	" Mo	CAR-08	KRF-13	S-9	S-10	S-12	S-15		
Cu	}													
Ni	}													
Co	"	"	flow	"	"	" W	KRF-13	PCC-1	PRO-1	S-9	S-12	S-15		
Cr	}													
V	}													
La	"	"	flow	"	"	" "	AGV-1	BCR-1	G-2					
Ce	}													
Nd	}													
Na	"	"	"	"	TLAP	f Cr	AGV-1	BCR-1	G-2	GSP-1	JG-1	NIM-N		

Full corrections made for background, deadtime, instrumental drift, spectral and tube line interference

USGS standards: AGV-1, BCR-1, G-2, GSP-1, JG-1, PCC-1

NIM secondary reference standards: S-9, S-10, S-12, S-15

NIM primary standards: NIM-N, NIM-P

Rhodes University in-house standards: CAR-08, KRF-13, PRO-1

STANDARD CONCENTRATIONS - Trace elements

STD	Ba	Nb	Zr	Y	Sr	Rb	Zn	Cu	Ni	Co	Cr	V	La	Ce	Nd
AGV-1	1208												35.0	65.0	39.0
BCR-1	675												25.0	54.0	29.0
CAR-08	243	18.1	125.0	27.7	296.0	18.6	76.5	64.6	69.5						
G-2													96.0	165	55.0
GSP-1	1300														
JG-1															
KRF-13	205	5.4	100.5	30.7	205.0	10.1	75.0	59.5	51.4	45.5	293.0	259.0			
NIM-G															
NIM-P	40														
PCC-1										112.0	2730.0	30.0			
PRO-1	736									21.7	62.0	100.0			
S-09		8.0	137.0	17.2	54.3	56.7	73.0	43.2	111.0	16.9	114.0	137.0			
S-10							38.0	18.0	128.0						
S-12	29	113.0	406.0	89.0	4.0	11.2	267.0	6.1	15.0	8.5	385.0	427.0			
S-15										70.0	1995.0	17.0			

STANDARD CONCENTRATIONS - Major elements

STD	SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	P ₂ O ₅	LOI
AGV-1	59.00	1.04	17.25	6.76	0.10	1.53	4.90	4.26	2.89	0.49	1.01
BCR-1	54.50	2.20	13.61	13.40	0.18	3.46	6.92	3.27	1.70	0.36	0.39
G-2	69.11	0.50	15.40	2.65	0.03	0.76	1.94	4.07	4.51	0.14	0.55
GSP-1	67.38	0.66	15.25	4.33	0.04	0.96	2.02	2.80	5.53	0.28	0.63
JG-1	72.24	0.26	14.21	2.21	0.06	0.73	2.18	3.39	3.96	0.10	0.45
NIM-N	52.43	0.19	16.64	9.00	0.17	7.43	11.55	2.44	0.26	0.04	0.07

APPENDIX III

Table A.3 CIPW weight percent norms for Witwatersrand triad samples

$$CL = FO + FA$$

$$OPX = HYEN + HYFS$$

$$CPX = DIEN + DIFS + DIWC$$

$$NCI = GL + CPX + CPX + MT + IL + HM$$

$$NPC = 100 AN / (AN + AB + 5/3NE)$$

(after Irvine & Baragar, 1971)

$$DI = Q + OR + AB + NE + LC + KS$$

(after Thornton & Tuttle, 1960)

CIPW WEIGHT PERCENT NORM

SAMPLE	DR- 1	DR- 2	DR- 3	DP- 4	DP- 5	DP- 6	DB- 7	DR- 8	DR- 9	DR- 10	DR- 11	DR- 12
AP	.59	.62	.54	.64	.66	.57	.45	.45	.43	.99	.92	1.49
IL	2.01	1.93	1.86	1.58	1.52	1.56	1.63	1.65	1.58	2.18	2.46	3.14
OR	5.47	15.05	14.74	24.45	35.67	35.75	.97	1.39	1.80	5.52	4.13	19.21
AB	51.69	47.68	49.82	59.39	54.25	45.57	32.74	50.94	58.30	70.34	64.97	6.70
AN	42.84	37.28	35.44	16.16	10.09	18.68	66.29	47.67	39.90	24.14	30.90	74.09
C	0.00	0.00	0.00	1.31	1.40	1.51	0.00	0.00	0.00	0.00	0.00	0.00
MT	2.93	2.97	3.03	1.90	1.46	1.74	3.32	3.54	2.97	2.60	3.28	2.96
DIEN	1.72	1.82	1.83	0.00	0.00	0.00	5.04	2.09	3.63	1.96	1.03	1.97
DIFS	3.12	3.27	3.76	0.00	0.00	0.00	4.79	1.63	2.49	2.15	1.37	7.90
DIWO	4.74	4.98	5.42	0.00	0.00	0.00	10.05	3.85	6.39	4.16	2.40	9.23
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	6.95	6.92	5.99	6.20	4.28	4.23	13.99	21.68	19.30	10.24	11.06	1.66
HYFS	12.61	12.44	12.35	9.83	7.32	8.82	13.31	16.90	13.21	11.25	14.71	6.66
Q	11.34	11.63	11.89	26.82	30.19	31.83	4.77	0.00	0.00	3.51	0.00	29.63
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.74	.39	0.00	.53	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.64	.29	0.00	.78	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.02	100.01	100.00	100.00	100.03	100.01	100.00	100.01	100.01	100.03	100.01	100.01

OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.38	.68	0.00	1.31	0.00
OPX	19.56	19.37	18.34	16.03	11.60	13.05	27.30	38.58	32.51	21.49	25.77	8.33
CPX	9.57	10.06	11.01	0.00	0.00	0.00	19.88	7.57	12.51	8.28	4.80	19.10
NCI	34.1	34.3	34.2	19.5	14.6	16.4	52.1	52.7	50.2	34.5	37.6	33.5
NPC	45.3	43.9	41.6	21.4	15.7	29.1	66.9	48.3	40.6	25.5	32.2	91.7
DI	42.2	45.1	46.3	70.2	78.0	72.3	19.1	24.5	29.7	49.8	42.5	38.8

SAMPLE	DR- 13	DR- 14	DR- 15	DR- 16	DR- 17	DR- 18	DR- 19	DR- 20	DR- 21	DR- 22	DR- 23	DR- 24
AP	1.82	1.28	.45	.40	.40	.38	.36	.26	.38	.40	.45	1.09
IL	3.85	2.97	1.46	1.41	1.39	1.28	1.22	1.20	1.16	1.20	1.33	2.67
OR	16.12	8.25	11.66	7.82	2.48	7.77	6.13	6.82	10.38	8.42	9.91	14.30
AB	57.33	59.09	49.14	47.21	28.25	43.68	53.58	46.93	51.79	50.42	51.13	48.51
AN	26.54	32.66	39.20	44.98	69.27	48.55	40.29	46.24	37.83	41.16	38.95	37.19
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	3.81	2.83	2.93	2.99	3.60	2.71	2.54	2.70	2.45	2.38	2.58	3.18
DIEN	.10	1.37	3.69	5.42	4.83	4.92	3.38	5.12	3.36	2.59	2.34	.28
DIFS	.31	2.61	3.28	4.05	4.96	3.57	2.56	3.59	3.25	2.61	2.31	.95
DIWO	.38	3.89	7.16	9.84	9.96	8.84	6.17	9.08	6.75	5.29	4.74	1.15
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	5.88	6.00	14.17	13.90	14.47	15.08	6.31	11.86	10.46	10.24	11.80	4.50
HYFS	18.68	11.43	12.57	10.38	14.88	10.95	4.78	8.32	10.10	10.31	11.64	15.38
Q	10.60	16.11	0.00	0.00	2.70	0.00	0.00	0.00	3.47	6.77	1.41	21.66
FO	0.00	0.00	.03	1.78	0.00	.31	5.96	2.87	0.00	0.00	0.00	0.00
FA	0.00	0.00	.03	1.46	0.00	.25	4.98	2.21	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.04	100.03	100.00	100.02	100.02	100.01	100.01	100.02	100.02	100.02	100.00	100.04

OL	0.00	0.00	.05	3.24	0.00	.56	10.95	5.08	0.00	0.00	0.00	0.00
OPX	24.56	17.43	26.75	24.28	29.35	26.02	11.10	20.18	20.56	20.54	23.45	19.88
CPX	.79	7.87	14.12	19.31	19.75	17.34	12.11	17.78	13.36	10.49	9.40	2.38
NCI	33.0	31.1	45.3	51.2	54.1	47.9	37.9	46.9	37.5	34.6	36.8	28.1
NPC	31.6	35.6	44.4	48.8	71.0	52.6	42.9	49.6	42.2	44.9	43.2	43.4
DI	50.7	50.8	33.0	26.6	15.9	26.6	36.9	28.4	39.9	41.0	38.9	52.6

SAMPLE	DR- 25	DR- 26	DR- 27	DR- 28	DR- 29	DR- 30	DR- 31	DP- 32	DP- 33	DP- 34	DP- 35	DP- 36
AP	.59	.43	.69	.40	1.07	.76	.71	.69	1.21	.76	.59	.66
IL	1.48	1.16	2.33	2.63	2.85	2.12	2.01	1.67	2.78	1.67	1.45	1.52
OR	21.69	18.69	6.52	12.12	9.23	7.51	8.23	23.53	72.41	60.33	47.83	31.30
AB	47.65	46.81	47.12	60.75	50.13	48.22	51.17	57.58	18.86	28.14	47.19	61.19
AN	30.66	34.50	46.36	27.13	40.64	44.27	40.60	18.89	8.73	11.53	4.98	7.52
C	0.00	0.00	0.00	5.73	2.35	0.00	0.00	0.00	9.87	8.31	5.13	1.35
MT	1.64	2.06	3.51	4.23	3.91	3.67	3.44	1.71	1.65	3.23	2.38	1.41
DIEN	.05	2.93	2.40	0.00	0.00	.30	1.51	.82	0.00	0.00	0.00	0.00
DIFS	.17	2.36	4.91	0.00	0.00	.72	3.00	2.47	0.00	0.00	0.00	0.00
DIWO	.21	5.47	7.10	0.00	0.00	.98	4.39	3.13	0.00	0.00	0.00	0.00
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	2.59	10.87	6.64	6.87	7.02	7.80	7.78	2.04	3.46	5.35	3.34	.97
HYFS	8.09	8.75	13.59	22.40	20.33	18.86	15.40	6.14	7.22	17.44	12.63	6.93
Q	34.23	3.09	5.36	17.79	18.93	12.89	7.21	22.56	15.96	33.17	41.51	30.33
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.01	100.01	100.03	100.03	100.02	100.03	100.01	100.01	100.02	100.01	100.01	100.01

OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	10.68	19.62	20.23	29.28	27.36	26.65	23.17	8.18	10.68	22.79	15.97	7.90
CPX	.43	10.76	14.41	0.00	0.00	2.00	8.90	6.42	0.00	0.00	0.00	0.00
NCI	14.2	33.6	40.5	36.1	34.1	34.4	37.5	18.0	15.1	27.7	19.8	10.8
NPC	39.2	42.4	49.6	30.9	44.8	47.9	44.2	24.7	31.6	29.1	9.5	10.9
DI	69.6	44.3	34.1	46.9	44.8	41.8	39.6	70.2	68.8	59.8	72.9	82.9

CIPW WEIGHT PERCENT NORM

SAMPLE	DP- 37	DP- 38	DP- 39	DP- 40	DR- 41	KL- 42	KL- 43	KL -44	KL- 45	KL -46	DP- 47	DP- 48
AP	.64	.66	.59	.62	.36	.09	.09	.07	.09	.09	.50	.50
IL	1.54	1.50	1.46	1.41	2.46	.99	.83	.79	.81	.86	1.39	1.43
OR	29.83	34.02	17.11	44.90	6.58	20.87	21.59	7.95	8.43	9.57	26.99	24.57
AR	57.98	51.22	76.93	53.57	46.93	71.25	17.74	27.63	24.85	27.24	58.19	62.09
AN	12.19	14.76	5.97	1.53	46.49	7.88	60.68	64.42	66.71	63.19	14.81	13.34
C	.00	0.00	3.38	1.52	0.00	7.66	0.00	0.00	0.00	0.00	0.00	0.00
MT	1.62	1.51	1.86	.48	3.89	3.29	2.65	2.48	2.46	2.61	1.96	2.12
DIEN	0.00	.04	0.00	0.00	.57	0.00	6.29	8.01	6.91	6.80	.00	.22
DIFS	0.00	.28	0.00	0.00	1.58	0.00	3.58	4.40	3.98	4.44	.00	2.01
DIWO	0.00	.30	0.00	0.00	2.05	0.00	10.43	13.15	11.49	11.78	.00	2.02
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	1.47	1.15	1.82	.80	6.88	12.40	12.44	11.56	14.97	13.14	1.72	1.00
HYFS	8.17	7.32	9.51	1.55	19.00	18.37	7.08	6.35	8.63	8.57	10.22	9.15
Q	26.16	29.91	34.76	39.18	14.75	10.75	0.00	0.00	0.00	0.00	26.22	22.16
FO	0.00	0.00	0.00	0.00	0.00	0.00	5.18	3.95	1.42	1.71	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	3.25	2.39	.90	1.23	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00

TOTAL	100.01	100.01	100.01	100.03	100.02	100.00	100.01	100.00	100.00	100.00	100.01	100.01
OL	0.00	0.00	0.00	0.00	0.00	0.00	8.44	6.34	2.32	2.94	0.00	0.00
OPX	9.64	8.47	11.33	2.35	25.88	30.78	19.52	17.91	23.60	21.71	11.93	10.16
CPX	0.00	.63	0.00	0.00	4.20	0.00	20.30	25.56	22.37	23.02	.01	4.25
NCI	12.8	12.1	14.7	4.2	36.4	35.1	51.7	53.1	51.6	51.1	15.3	17.9
NPC	17.4	22.4	7.2	2.8	49.8	10.0	77.4	70.0	72.9	69.9	20.3	17.7
DI	79.2	78.8	78.6	92.8	40.7	53.5	18.9	16.7	16.1	18.0	75.6	73.6

SAMPLE	DP- 49	DR- 50	DP- 51	DR- 52	DR- 53	DR- 54	DR- 55	DR- 56	DR- 57	DR- 58	DP- 59	DR- 60
AP	.54	.59	.62	.43	.85	1.44	.40	.33	.31	.73	.62	.76
IL	1.48	1.95	1.46	1.50	2.14	3.10	1.41	1.24	1.20	1.80	1.45	1.84
OR	35.81	12.33	26.12	4.58	8.71	18.53	8.60	7.16	14.01	10.66	63.20	14.69
AR	50.07	47.49	62.28	46.64	58.63	55.10	44.32	53.79	42.12	45.49	22.74	34.29
AN	14.12	40.17	11.60	48.78	32.67	26.37	47.08	39.05	43.86	43.84	14.06	51.02
C	.72	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.08	0.00
MT	2.26	2.25	1.62	3.23	3.02	3.16	2.64	2.32	2.48	2.65	1.45	2.70
DIEN	0.00	1.01	.19	3.41	2.59	1.85	5.93	4.84	5.31	2.44	0.00	1.81
DIFS	0.00	2.43	.64	2.16	3.56	5.41	4.57	3.98	3.86	3.95	0.00	3.02
DIWO	0.00	3.31	.78	5.84	6.14	6.90	10.88	9.10	9.54	6.30	0.00	4.75
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	.65	3.79	2.23	20.36	8.91	3.53	12.69	10.43	11.14	6.25	3.01	6.71
HYFS	12.11	9.10	7.61	12.88	12.24	10.37	9.77	8.58	8.11	10.10	7.21	11.18
Q	26.48	19.99	25.36	0.00	3.56	10.37	2.27	2.23	0.00	12.59	32.37	19.38
FO	0.00	0.00	0.00	2.96	0.00	0.00	0.00	0.00	1.56	0.00	0.00	0.00
FA	0.00	0.00	0.00	2.06	0.00	0.00	0.00	0.00	1.25	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00

TOTAL	100.01	100.02	100.01	100.01	100.02	100.03	100.01	100.00	100.00	100.01	100.01	100.02
OL	0.00	0.00	0.00	5.03	0.00	0.00	0.00	0.00	2.81	0.00	0.00	0.00
OPX	12.76	12.90	9.84	33.24	21.15	13.90	22.46	19.01	19.25	16.34	10.22	17.89
CPX	0.00	6.76	1.61	11.41	12.29	14.16	21.38	17.92	18.71	12.69	0.00	9.59
NCI	16.5	23.9	14.5	54.4	38.6	34.3	47.9	40.5	44.5	33.5	13.1	32.0
NPC	22.0	45.8	15.7	51.1	35.8	32.4	51.5	42.1	51.0	49.1	38.2	59.8
DI	74.4	53.2	78.0	23.1	41.9	50.1	28.4	36.9	31.0	42.5	77.8	42.8

SAMPLE	DR- 61	DR- 62	DR- 63	DP- 64	DP- 65	DR- 66	DR- 67	DR- 68	DR- 69	DR- 72	DR- 73	DR- 75
AP	.38	.99	.83	.59	.64	.83	.88	.78	1.66	1.59	1.63	1.61
IL	1.13	2.67	2.25	1.41	1.46	2.14	2.18	2.12	3.64	3.44	3.51	3.38
OR	12.39	12.58	6.81	31.33	27.91	11.93	14.39	10.98	14.22	26.07	18.21	12.67
AR	43.45	41.60	56.18	56.23	62.02	58.35	55.58	58.44	51.53	46.61	55.53	49.84
AN	44.16	45.82	37.01	12.44	10.06	29.72	30.03	30.58	34.25	27.33	26.26	37.49
C	0.00	0.00	0.00	0.00	0.00	0.00	2.15	0.00	0.00	0.00	0.00	0.00
MT	2.20	3.61	3.38	1.45	1.45	2.99	3.02	3.04	3.36	3.45	3.58	3.28
DIEN	3.59	.92	1.49	.26	.30	1.89	0.00	2.04	.57	.88	.67	.54
DIFS	3.03	2.71	5.14	1.26	1.36	2.61	0.00	2.81	1.67	2.68	2.28	1.52
DIWO	6.82	3.46	6.25	1.42	1.55	4.49	0.00	4.83	2.13	3.37	2.79	1.96
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	10.63	5.48	3.72	1.25	1.29	9.49	14.67	9.54	5.04	4.75	4.61	5.19
HYFS	8.96	16.14	12.82	5.98	5.85	13.08	15.85	13.16	14.85	14.54	15.60	14.67
Q	6.00	17.06	5.11	29.07	27.04	4.07	9.55	4.65	20.16	17.97	15.37	19.86
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00

TOTAL	99.99	100.01	100.02	100.02	100.01	100.03	100.03	100.02	100.03	100.04	100.03	100.04
OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	19.59	21.62	16.54	7.23	7.14	22.57	30.51	22.71	19.88	19.29	20.21	19.86
CPX	13.44	7.09	12.87	2.94	3.21	9.00	0.00	9.67	4.37	6.93	5.75	4.01
NCI	36.4	35.0	35.0	13.0	13.3	36.7	35.7	37.5	31.3	33.1	33.0	30.5
NPC	50.4	52.4	39.7	18.1	14.0	33.7	35.1	34.4	39.9	37.0	32.1	42.9
DI	38.0	42.5	42.3	79.3	80.2	45.1	45.8	44.3	51.0	52.4	52.2	49.9

CIPW WEIGHT PERCENT NORM

SAMPLE	DR- 76	DR- 77	DR- 78	DR- 79	DP- 80	DP- 81	DB- 82	DR- 83	DR- 85	KL- 86	KL- 87	DR- 88
AP	1.40	1.40	.47	.36	.57	.57	.31	.33	.52	.07	.14	.69
IL	3.17	3.04	1.91	1.35	1.39	1.39	1.28	1.26	1.48	.77	1.15	2.16
OR	14.19	14.32	10.49	15.33	57.06	21.97	3.78	4.62	35.63	12.56	9.26	18.11
AR	46.86	47.95	57.53	44.19	31.55	38.29	40.45	47.32	35.16	27.89	38.78	43.85
AN	38.96	37.73	31.98	40.47	11.39	39.73	55.77	48.06	29.21	59.56	51.96	38.04
C	0.00	0.00	9.33	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.91
MT	3.36	2.75	3.77	2.58	1.44	1.55	2.97	2.84	2.39	2.36	2.90	3.58
DIEN	.34	1.28	0.00	1.84	.07	.21	5.79	5.74	1.03	7.42	4.95	0.00
DIFS	.72	2.58	0.00	1.25	.23	.81	4.22	4.33	1.10	4.37	4.18	0.00
DIWO	1.02	3.76	0.00	3.23	.28	.96	9.96	10.45	2.17	12.44	9.41	0.00
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	7.56	5.44	19.65	13.19	2.10	1.85	15.48	14.96	10.97	10.93	14.03	10.91
HYFS	16.24	10.96	20.43	8.96	6.96	7.12	12.13	11.30	11.67	6.44	11.86	19.14
Q	18.72	18.86	1.07	0.00	28.25	38.47	2.88	.18	7.81	0.00	.97	13.01
FO	0.00	0.00	0.00	3.86	0.00	0.00	0.00	0.00	0.00	2.93	0.00	0.00
FA	0.00	0.00	0.00	2.89	0.00	0.00	0.00	0.00	0.00	1.90	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.02	100.02	100.02	100.01	100.00	100.00	100.00	100.02	100.02	100.00	99.99	100.02

OL	0.00	0.00	0.00	6.76	0.00	0.00	0.00	0.00	0.00	4.83	0.00	0.00
OPX	23.80	16.40	40.08	22.15	9.06	8.97	27.60	26.26	22.64	17.37	25.89	30.05
CPX	2.08	7.63	0.00	6.31	.58	1.99	19.57	20.52	4.30	24.24	18.55	0.00
NCI	32.4	29.8	45.8	39.2	12.5	13.9	51.4	50.9	30.8	49.6	48.5	35.8
NPC	45.4	44.0	35.7	47.8	26.5	50.9	58.0	50.4	45.4	68.1	57.3	46.5
DI	47.7	50.0	30.6	36.0	80.3	66.8	23.0	25.4	50.9	20.4	25.2	43.1

SAMPLE	DP- 89	KL- 90	PR- 91	DR- 92	DR- 93	DR- 94	DB- 95	DP- 96	DR- 97	DP- 98	DP- 99	KA-100
AP	.62	.12	2.32	1.28	.40	.31	.59	.62	.33	.47	.43	.33
IL	1.56	.86	3.49	2.97	1.43	1.26	1.60	1.52	2.07	1.43	1.26	1.93
OR	31.57	19.18	1.49	10.13	7.19	8.83	12.96	41.66	2.58	52.22	32.37	11.99
AR	63.97	29.26	36.74	47.35	44.79	50.26	48.90	51.11	49.67	40.50	57.72	12.78
AN	4.46	51.55	61.78	42.52	48.02	40.91	38.14	7.23	47.74	7.27	9.92	75.24
C	2.75	0.00	0.00	0.00	0.00	0.00	0.00	1.11	0.00	2.73	3.02	0.00
MT	1.71	2.36	3.73	3.74	3.19	2.68	2.97	1.55	3.67	1.86	2.29	2.86
DIEN	0.00	6.63	.27	1.28	5.28	5.80	.85	0.00	1.23	0.00	0.00	2.71
DIFS	0.00	3.48	.26	2.95	4.10	4.06	1.06	0.00	2.22	0.00	0.00	2.94
DIWO	0.00	10.74	.54	4.08	9.72	10.28	1.92	0.00	3.38	0.00	0.00	5.73
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	2.89	8.75	18.81	7.11	13.83	11.66	12.05	2.49	9.67	1.79	3.21	11.13
HYFS	8.62	4.60	18.68	16.40	10.74	8.16	14.99	7.69	17.44	9.56	12.23	12.08
Q	33.22	0.00	8.21	11.81	0.00	0.00	5.14	29.59	11.97	33.27	37.39	15.92
FO	0.00	6.77	0.00	0.00	2.42	2.43	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	3.92	0.00	0.00	2.07	1.87	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.02	100.00	100.06	100.03	100.03	100.01	100.00	100.02	100.02	100.00	100.00	100.01

OL	0.00	10.69	0.00	0.00	4.50	4.30	0.00	0.00	0.00	0.00	0.00	0.00
OPX	11.50	13.35	37.49	23.51	24.57	19.82	27.04	10.18	27.12	11.36	15.45	23.21
CPX	0.00	20.86	1.07	8.31	19.10	20.13	3.84	0.00	6.84	0.00	0.00	11.39
NCI	14.8	48.1	45.8	38.5	52.8	48.2	35.4	13.3	39.7	14.6	19.0	39.4
NPC	6.5	63.8	62.7	47.3	51.7	44.9	43.8	12.4	49.0	15.2	14.7	85.5
DI	79.7	25.1	24.9	39.6	24.4	30.4	41.5	81.0	37.1	78.6	73.6	26.9

SAMPLE	KA-101	KA-102	KA-103	KA-104	KA-105	KA-106	KA-107	KO-108	KO-109	KO-110	KO-111	KO-112
AP	.33	.38	.36	.43	.40	.40	.38	.28	.31	.28	.28	.28
IL	2.01	2.03	1.93	2.01	2.05	1.90	2.07	1.91	2.01	1.93	2.03	1.86
OR	7.22	7.22	6.82	14.91	14.21	13.12	6.79	7.47	12.39	8.01	9.26	8.34
AR	55.81	56.73	59.94	55.38	46.68	49.15	56.24	57.07	50.25	56.90	54.30	42.33
AN	36.97	36.05	33.24	29.71	39.12	37.73	36.96	35.47	37.36	35.09	36.44	49.34
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	3.07	3.00	2.90	2.77	3.00	2.75	2.99	3.32	3.60	3.41	3.73	3.29
DIEN	.74	.70	1.50	3.11	3.29	3.21	.96	2.17	1.09	1.89	.89	3.45
DIFS	.86	.79	1.73	3.77	4.57	4.01	1.48	2.97	1.55	2.70	1.25	5.52
DIWO	1.62	1.51	3.26	6.92	7.82	7.25	2.42	5.13	2.63	4.56	2.13	8.85
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	13.30	13.34	11.70	8.87	8.07	8.42	9.25	10.81	12.53	10.91	9.35	7.61
HYFS	15.36	14.95	13.49	10.76	11.20	10.51	14.18	14.76	17.77	15.59	13.14	12.17
Q	7.07	2.76	5.98	.73	1.81	3.88	6.26	0.00	1.29	1.17	0.00	3.80
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.03	0.00	0.00	2.83	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.04	0.00	0.00	4.38	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.00	100.02	100.01	100.02	100.02	100.01	100.00	100.00	100.01	100.00	99.99	100.00

OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.07	0.00	0.00	7.21	0.00
OPX	28.66	28.29	25.19	19.63	19.27	18.93	23.43	25.57	30.30	26.51	22.49	19.77
CPX	3.22	3.00	6.48	13.80	15.68	14.48	4.86	10.27	5.26	9.15	4.26	17.83
NCI	37.0	36.3	36.5	38.2	40.0	38.1	33.3	41.1	41.2	41.0	39.7	42.7
NPC	39.8	38.9	35.7	34.9	45.6	43.4	39.7	38.3	42.6	38.1	40.2	53.8
DI	42.1	41.5	44.1	43.4	37.0	39.8	44.1	37.8	37.1	38.5	38.1	30.7

CIPW WEIGHT PERCENT NORM

SAMPLE	KO-113	KO-114	KO-115	KO-116	KO-117	KO-118	KO-119	KO-120	KO-121	KO-122	KO-123	KO-124
AP	.36	.31	.33	.28	.28	.31	.33	.33	.36	.28	.26	.26
IL	1.90	1.90	1.84	1.90	1.88	1.73	1.99	1.90	1.75	1.75	1.63	1.65
OR	9.43	13.13	24.92	7.98	11.89	3.78	13.74	8.65	4.62	15.12	8.50	21.41
AB	43.55	42.68	40.30	57.57	52.57	45.99	49.59	75.14	4.79	46.87	84.31	62.08
AN	47.02	44.19	34.77	34.45	35.53	50.23	37.07	16.20	91.10	38.01	7.19	16.52
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	3.20	3.65	2.97	3.42	3.39	3.04	3.58	3.44	2.60	3.19	3.06	3.00
DIEN	3.90	2.35	3.15	1.31	1.54	4.38	3.09	4.03	5.75	2.04	4.07	4.24
DIFS	5.19	3.83	3.81	1.69	2.09	6.03	4.22	4.88	13.61	2.64	5.94	6.16
DIWO	9.08	6.09	7.00	3.01	3.61	10.37	7.29	8.97	18.64	4.68	9.94	10.33
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	3.23	0.00	0.00	0.00
HYEN	8.98	9.06	9.90	12.91	11.89	7.48	3.46	0.00	0.00	11.26	0.00	0.00
HYFS	11.96	14.77	11.96	16.63	16.15	10.30	4.72	0.00	0.00	14.56	0.00	0.00
Q	2.80	0.00	4.04	1.12	2.54	3.23	0.00	0.00	6.71	2.93	0.00	0.00
FO	0.00	.49	0.00	0.00	0.00	0.00	5.32	7.89	0.00	0.00	5.07	4.81
FA	0.00	.87	0.00	0.00	0.00	0.00	7.99	10.52	0.00	0.00	8.17	7.72
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	8.05	0.00	0.00	8.54	4.94
TOTAL	100.01	100.01	100.01	100.02	100.02	100.03	100.01	100.01	100.00	100.01	100.03	100.01
OL	0.00	1.36	0.00	0.00	0.00	0.00	13.31	18.41	0.00	0.00	13.24	12.53
OPX	20.94	23.83	21.86	29.53	28.04	17.78	8.18	0.00	0.00	25.82	0.00	0.00
CPX	18.16	12.26	13.96	6.01	7.24	20.77	14.61	17.89	38.00	9.35	19.95	20.73
NCI	44.2	43.0	40.6	40.9	40.5	43.3	41.7	41.6	42.3	40.1	37.9	37.9
NPC	51.9	50.9	46.3	37.4	40.3	52.2	42.8	17.7	95.5	44.8	7.9	21.0
DI	30.7	31.6	39.9	39.0	39.1	29.7	36.5	48.6	10.9	38.1	57.4	51.6

SAMPLE	KO-125	KO-126	KO-127	KO-128	KO-129	KO-130	KL-131	KL-132	KL-133	KL-134	KL-135	KL-136
AP	.26	.29	.28	.26	.31	.31	.24	.21	.21	.19	.12	.17
IL	1.69	1.73	1.69	1.67	1.88	1.67	1.43	1.43	1.45	1.20	.90	1.22
OR	11.15	18.23	17.81	23.78	9.27	27.39	21.06	4.05	9.96	8.69	2.97	15.61
AB	55.38	45.50	52.72	43.62	46.90	34.56	53.75	58.29	55.40	54.17	43.28	35.87
AN	33.47	36.26	29.47	32.60	43.83	38.05	25.18	37.66	34.65	37.14	53.74	48.51
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	3.10	3.32	3.22	3.10	3.54	2.99	2.97	2.86	2.87	2.77	3.07	3.06
DIEN	1.99	3.25	2.36	2.64	3.18	2.56	4.05	2.87	3.02	5.58	5.51	4.79
DIFS	2.74	4.33	3.68	3.89	5.09	3.74	4.04	2.78	3.07	5.28	2.83	4.50
DIWO	4.71	7.57	5.97	6.48	8.16	6.26	8.24	5.77	6.20	11.11	8.86	9.50
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	10.11	4.14	8.79	8.69	5.27	8.47	5.53	13.14	12.24	10.56	28.09	13.17
HYFS	13.90	5.51	13.68	12.82	8.43	12.34	5.51	12.72	12.44	10.00	14.43	12.37
Q	2.68	0.00	1.59	2.07	0.00	5.36	0.00	4.24	6.26	5.05	.41	2.43
FO	0.00	4.28	0.00	0.00	2.45	0.00	4.74	0.00	0.00	0.00	0.00	0.00
FA	0.00	6.27	0.00	0.00	4.32	0.00	5.20	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.01	100.01	100.01	100.01	100.00	100.00	100.02	100.01	100.00	100.00	99.99	100.01
OL	0.00	10.55	0.00	0.00	6.77	0.00	9.94	0.00	0.00	0.00	0.00	0.00
OPX	24.01	9.66	22.48	21.51	13.70	20.81	11.03	25.86	24.68	20.56	42.51	25.54
CPX	9.44	15.16	12.01	13.01	16.43	12.56	16.34	11.41	12.29	21.97	17.20	18.78
NCI	38.2	40.4	39.4	39.3	42.3	38.0	41.7	41.6	41.3	46.5	63.7	48.6
NPC	37.7	44.3	35.9	42.8	48.3	52.4	31.9	39.2	38.5	40.7	55.4	57.5
DI	41.8	37.8	43.0	41.4	32.2	40.2	43.5	37.9	40.4	35.4	17.0	27.6

SAMPLE	KL-137	KL-138	KL-139	KL-140	KL-141	KL-142	KL-143	KL-144	KL-145	KL-146	KL-147	KL-148
AP	.24	.21	.24	.14	.19	.14	.17	.17	.17	.19	.19	.17
IL	1.28	1.30	1.30	.99	1.07	1.16	1.15	1.15	1.13	.96	1.11	1.07
OR	6.84	14.66	7.68	23.96	22.83	7.84	15.19	19.71	23.93	20.04	17.38	14.07
AB	46.49	48.34	47.89	44.00	30.54	51.13	45.73	47.45	2.95	35.36	54.34	49.64
AN	46.67	37.00	44.44	32.03	46.63	41.03	39.08	32.84	73.11	44.60	28.27	36.29
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	3.18	3.09	3.15	2.84	3.02	3.12	2.96	2.90	2.96	2.87	3.00	2.87
DIEN	4.95	3.95	1.97	7.61	5.97	6.77	5.83	3.69	3.62	3.62	4.61	4.60
DIFS	4.56	3.98	1.83	4.95	3.83	5.69	4.66	3.16	3.15	1.83	4.00	3.74
DIWO	9.74	8.08	3.90	13.16	10.28	12.84	10.85	7.05	6.95	5.80	8.85	8.61
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	10.28	3.43	16.65	10.39	18.50	2.44	11.66	13.86	15.26	17.24	8.63	13.55
HYFS	9.48	3.46	15.45	6.75	11.89	2.05	9.32	11.85	13.28	8.70	7.50	11.01
Q	0.00	0.00	1.64	0.00	0.00	0.00	0.00	0.00	10.53	0.00	0.00	0.00
FO	2.66	6.66	0.00	4.54	1.19	7.94	2.14	.82	0.00	7.65	4.30	1.00
FA	2.70	7.40	0.00	3.25	.84	7.36	1.89	.78	0.00	4.25	4.12	.89
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.03	100.02	100.02	100.01	100.01	100.00	100.01	99.99	100.01	100.01	100.02	100.01
OL	5.36	14.06	0.00	7.80	2.04	15.31	4.03	1.60	0.00	11.90	8.42	1.89
OPX	19.76	6.88	32.10	17.14	30.39	4.50	20.97	25.70	28.55	25.95	16.13	24.56
CPX	19.26	16.02	7.70	25.71	20.08	25.29	21.34	13.91	13.72	11.24	17.46	16.95
NCI	48.8	41.3	44.2	54.5	56.6	49.4	50.4	45.3	46.3	52.9	46.1	47.3
NPC	50.1	43.4	48.1	42.1	60.4	44.5	46.1	40.9	96.1	55.8	34.2	42.2
DI	27.2	36.8	31.6	30.8	23.1	29.8	30.1	36.6	22.1	26.0	38.5	33.4

CIPW WEIGHT PERCENT NORM

SAMPLE	KL-149	KL-150	KL-151	KL-152	KL-153	KL-154	KL-155	KL-156	KL-157	KL-158	KL-159	PG-160
AP	.17	.19	.31	.24	.24	.17	.19	.12	.09	.14	.17	1.23
IL	1.03	1.24	1.20	1.24	1.37	.88	.96	.77	.77	.84	.77	2.23
OR	19.31	14.88	22.72	12.37	23.57	18.82	17.50	4.67	2.00	6.13	4.74	17.43
AB	46.87	45.02	48.57	50.48	44.82	38.20	40.19	41.05	38.70	34.21	33.68	54.61
AN	33.81	40.11	28.71	37.15	31.61	42.98	42.32	54.28	59.29	59.66	61.58	27.96
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	4.13
MT	2.84	3.06	2.65	2.75	3.13	2.78	2.64	2.93	2.81	3.12	2.91	2.55
DIEN	4.81	4.02	3.18	3.60	2.21	4.30	2.88	5.34	4.43	4.85	4.53	0.00
DIFS	3.69	3.73	3.57	3.51	2.17	2.54	1.53	2.58	1.99	2.32	2.15	0.00
DIWO	8.82	7.93	6.82	7.26	4.46	7.21	4.67	8.45	6.89	7.65	7.13	0.00
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	8.54	7.38	9.75	11.86	4.08	20.77	24.84	27.80	30.80	22.99	30.18	14.42
HYFS	6.55	6.85	10.93	11.56	4.01	12.26	13.19	13.42	13.85	10.99	14.30	12.92
Q	0.00	0.00	3.08	1.15	0.00	0.00	.47	1.99	2.94	0.00	2.94	22.88
FO	5.09	4.73	0.00	0.00	7.88	1.02	0.00	0.00	0.00	6.24	0.00	0.00
FA	4.30	4.84	0.00	0.00	8.55	.66	0.00	0.00	0.00	3.29	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.00	99.98	100.02	100.00	100.01	100.00	100.00	99.99	100.00	100.00	100.01	100.03
OL	9.39	9.57	0.00	0.00	16.43	1.68	0.00	0.00	0.00	9.52	0.00	0.00
OPX	15.09	14.24	20.68	23.43	9.09	33.03	38.03	41.23	44.66	33.97	44.49	27.34
CPX	17.32	15.67	13.56	14.37	8.84	14.05	9.08	16.37	13.32	14.81	13.81	0.00
NCI	45.7	43.8	38.1	41.8	37.9	52.4	50.7	61.2	61.6	62.3	62.0	32.1
NPC	41.9	47.1	37.2	42.4	41.4	52.9	51.3	56.9	60.5	63.6	64.6	33.9
DI	35.8	33.5	44.8	36.9	42.3	27.0	28.5	18.8	17.4	15.2	16.4	51.5

SAMPLE	PG-161	PG-162	PG-163	PG-164	PG-165	PG-166	PG-167	PG-168	PG-169	PG-170	PG-171A	PG-171B
AP	1.28	1.26	1.87	1.92	1.78	1.75	1.87	1.87	1.73	1.89	1.87	1.80
IL	2.38	2.33	3.02	3.06	3.04	2.93	3.14	3.00	2.93	3.10	3.10	3.00
OR	27.64	29.64	11.56	16.97	13.56	6.67	2.02	12.23	5.73	10.31	21.55	15.23
AB	47.61	43.64	41.45	39.00	46.31	55.21	54.44	46.80	50.72	52.24	47.61	46.57
AN	24.75	26.72	47.00	44.03	40.12	38.12	43.54	40.97	43.55	37.46	30.84	38.21
C	0.00	2.75	0.00	0.00	0.00	0.00	.23	1.32	0.00	2.56	0.00	0.00
MT	2.33	2.51	2.97	3.00	3.31	3.02	2.80	3.18	2.73	2.68	2.51	2.60
DIEN	1.12	0.00	.09	3.02	1.62	1.59	0.00	0.00	2.68	0.00	2.00	2.50
DIFS	1.23	0.00	.08	2.63	1.55	1.57	0.00	0.00	3.94	0.00	2.63	3.54
DIWO	2.38	0.00	.17	5.81	3.23	3.23	0.00	0.00	6.57	0.00	4.63	6.02
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	9.39	13.27	16.70	14.16	15.87	13.70	13.05	17.58	6.48	13.85	7.16	6.41
HYFS	10.38	12.66	14.70	12.32	15.21	13.53	13.73	16.09	9.53	12.99	9.41	9.08
Q	11.61	15.99	10.84	1.20	3.47	1.99	12.70	6.76	11.89	13.09	3.12	10.77
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.04	100.01	100.04	100.05	100.03	100.02	100.05	100.05	100.03	100.05	100.05	100.05
OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	19.77	25.94	31.40	26.48	31.08	27.23	24.78	33.67	16.01	26.83	16.57	15.49
CPX	4.73	0.00	.33	11.46	6.40	6.39	0.00	0.00	13.20	0.00	9.26	12.07
NCI	29.2	30.8	37.7	44.0	43.8	39.6	32.7	39.9	34.9	32.6	31.4	33.2
NPC	34.2	38.0	53.1	53.0	46.4	40.8	44.4	46.7	46.2	41.8	39.3	45.1
DI	55.2	52.1	37.1	30.8	34.0	37.1	42.4	36.4	41.0	44.3	47.1	44.3

SAMPLE	PG-172	PG-173	PG-174	PG-175	PG-176	PG-177	PM-178	PM-179	PM-180	PR-181	PR-182	PR-183
AP	1.87	1.68	1.87	1.54	1.54	.09	1.09	1.28	1.33	1.61	1.42	1.73
IL	2.98	2.72	3.21	2.67	2.61	.47	2.08	2.20	2.42	2.67	2.38	2.68
OR	20.16	22.68	9.82	23.76	14.61	18.79	59.50	8.30	12.60	5.11	12.84	23.51
AB	47.23	47.59	42.99	47.87	45.65	5.91	23.56	59.83	45.41	51.05	53.10	44.49
AN	32.61	29.73	47.19	28.38	39.74	75.31	16.94	31.87	41.98	43.84	34.07	32.00
C	0.00	0.00	2.72	1.51	0.00	1.72	3.76	0.00	.59	1.87	0.00	0.00
MT	2.39	2.02	3.52	2.38	2.23	2.35	2.02	1.49	2.29	2.77	2.55	2.29
DIEN	2.13	1.65	0.00	0.00	2.28	0.00	0.00	.11	0.00	0.00	2.21	1.65
DIFS	2.59	2.18	0.00	0.00	3.51	0.00	0.00	.15	0.00	0.00	1.98	1.79
DIWO	4.75	3.83	0.00	0.00	5.73	0.00	0.00	.27	0.00	0.00	4.29	3.49
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	7.33	5.47	15.42	10.43	4.79	9.54	8.37	5.04	9.11	16.09	12.09	8.58
HYFS	8.88	7.26	17.85	11.62	7.36	13.34	9.99	6.72	11.27	13.95	10.83	9.33
Q	5.86	17.22	3.10	14.90	14.58	59.61	26.03	25.79	17.68	11.37	8.41	5.16
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.05	100.05	100.03	100.01	100.05	100.01	100.01	100.04	100.03	100.05	100.02	100.04
OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	16.21	12.74	33.26	22.05	12.15	22.88	18.36	11.76	20.39	30.04	22.92	17.91
CPX	9.47	7.66	0.00	0.00	11.52	0.00	0.00	.53	0.00	0.00	8.47	6.93
NCI	31.1	25.1	40.0	27.1	28.5	25.7	22.5	16.0	25.1	35.5	36.3	29.8
NPC	40.8	38.5	52.3	37.2	46.5	92.7	41.8	34.8	48.0	46.2	39.1	41.9
DI	47.1	56.6	30.7	54.3	48.0	62.8	64.8	64.6	49.8	39.3	43.9	48.2

CIPW WEIGHT PERCENT NORM

SAMPLE	PR-184	PR-185	PR-186	PR-187	PR-188	PR-189	PR-190	PR-191	PR-192	PR-193	PR-194	AR-195
AP	2.75	2.27	2.30	1.94	2.06	1.97	1.85	1.92	1.54	1.59	1.56	.47
IL	4.09	3.40	3.57	3.19	3.19	3.06	2.98	2.97	2.70	2.57	2.72	2.31
OR	2.05	6.88	7.04	21.71	5.94	1.59	23.39	9.76	14.73	8.08	17.14	3.17
AB	46.76	45.25	43.39	46.26	46.87	60.55	48.24	49.89	52.72	43.09	17.48	61.79
AN	51.19	47.87	49.57	32.03	47.18	37.86	28.38	40.35	32.55	48.83	65.38	35.04
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.14	0.00	.74	0.00	0.00
MT	3.68	3.33	3.23	2.87	2.90	2.81	2.61	3.38	1.64	3.07	2.77	3.12
DIEN	1.59	.17	1.37	.80	.27	1.15	1.73	0.00	2.97	0.00	.18	1.22
DIFS	2.14	.18	1.55	1.03	.35	1.18	2.11	0.00	3.33	0.00	.14	2.45
DIWO	3.72	.35	2.94	1.83	.62	2.37	3.86	0.00	6.37	0.00	.33	3.56
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	6.44	15.82	12.63	10.13	10.83	12.30	8.73	13.55	3.45	13.35	18.00	6.85
HYFS	8.70	16.48	14.30	13.02	13.86	12.70	10.63	17.22	3.87	15.82	13.75	13.79
Q	0.00	9.07	6.10	7.79	13.25	6.75	8.18	3.75	10.76	12.42	10.25	10.37
FO	3.71	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	5.52	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.06	100.05	100.05	100.03	100.06	100.03	100.03	100.03	100.02	100.04	100.03	100.01

OL	9.23	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	15.14	32.30	26.93	23.15	24.70	25.00	19.36	30.77	7.32	29.17	31.75	20.65
CPX	7.45	.70	5.86	3.65	1.25	4.70	7.70	0.00	12.67	0.00	.64	7.23
NCI	39.6	39.7	39.6	32.9	32.0	35.6	32.6	37.1	24.3	34.8	37.9	33.3
NPC	52.3	51.4	53.3	40.9	50.2	38.5	37.0	44.7	38.2	53.1	78.9	36.2
DI	28.2	34.6	32.4	46.8	41.1	41.4	49.3	36.6	53.5	38.2	27.7	46.7

SAMPLE	AR-196	AR-197	AR-198	AR-199	KA-200	KO-201	KO-202	KL-203	KL-204	KL-205	PG-206	PG-207
AP	.52	.47	.50	.54	.38	.31	.31	.17	.17	.14	1.33	1.85
IL	2.29	2.16	2.20	2.14	1.99	1.82	1.88	1.26	.94	.79	2.38	2.97
OR	6.23	5.68	13.67	14.41	15.80	24.23	10.23	15.04	18.22	8.04	23.69	8.59
AB	60.10	48.97	57.20	59.11	41.95	35.17	48.50	47.48	42.35	42.46	43.04	50.03
AN	33.67	45.35	29.13	26.48	42.26	40.61	41.27	37.47	39.43	49.50	33.27	41.38
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	3.06	3.90	2.60	2.96	3.00	3.20	3.44	2.99	2.64	2.70	2.62	3.06
DIEN	2.15	.06	1.30	2.53	1.32	1.80	1.85	3.52	4.90	6.38	.91	2.88
DIFS	6.29	.13	2.18	3.79	1.83	2.80	2.59	3.38	3.15	3.22	.95	2.80
DIWO	8.02	.19	3.42	6.26	3.13	4.55	4.42	7.05	8.44	10.22	1.89	5.80
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	3.31	9.92	6.62	7.78	10.07	9.25	11.35	13.59	17.89	23.82	11.82	12.96
HYFS	9.69	20.77	11.10	11.68	14.00	14.35	15.88	13.07	11.51	12.00	12.37	12.57
Q	6.82	8.34	14.34	4.93	10.43	9.26	2.83	1.16	3.03	3.26	10.12	4.26
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.01	100.03	100.00	100.01	100.00	100.03	100.02	99.99	100.02	100.00	100.02	100.03

OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	13.00	30.69	17.72	19.46	24.07	23.60	27.23	26.66	29.40	35.82	24.18	25.52
CPX	16.46	.39	6.89	12.58	6.28	9.15	8.86	13.95	16.49	19.82	3.75	11.47
NCI	34.8	37.1	29.4	37.1	35.3	37.8	41.4	44.9	49.5	59.1	32.9	43.0
NPC	35.9	48.1	33.7	30.9	50.2	53.6	46.0	44.1	48.2	53.8	43.6	45.3
DI	45.2	37.9	53.9	47.1	41.5	40.5	35.4	34.8	31.7	22.2	47.2	34.1

SAMPLE	PG-208	PM-209	PM-210	PM-212	PM-213	PM-217	PM-219	PM-222	PM-228	PM-231	PM-233	PR-234
AP	2.42	1.37	1.02	1.28	1.21	1.11	1.26	1.09	1.07	1.56	1.26	1.49
IL	3.53	2.55	1.99	2.29	2.27	2.14	2.33	2.03	1.99	2.87	2.20	2.52
OR	4.35	77.89	44.83	.66	55.57	23.38	28.09	11.21	7.02	30.67	18.90	8.08
AB	56.92	7.51	40.85	88.82	26.12	51.52	58.28	59.90	56.99	52.85	60.40	50.41
AN	38.74	14.60	14.31	10.52	18.31	25.10	13.62	28.89	35.99	16.49	20.71	41.51
C	1.07	3.52	0.00	1.90	2.97	0.00	5.59	.02	0.00	5.14	7.56	0.00
MT	3.20	2.23	1.57	1.65	1.52	1.93	2.55	1.57	1.67	2.00	3.39	2.57
DIEN	0.00	0.00	.14	0.00	0.00	.33	0.00	0.00	1.94	0.00	0.00	1.23
DIFS	0.00	0.00	.56	0.00	0.00	.74	0.00	0.00	6.00	0.00	0.00	1.36
DIWO	0.00	0.00	.66	0.00	0.00	1.03	0.00	0.00	7.53	0.00	0.00	2.62
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	14.15	9.31	1.77	4.11	4.26	3.90	8.09	4.21	.67	6.20	12.85	10.42
HYFS	15.75	10.73	6.89	7.75	6.86	8.74	12.93	7.49	2.07	9.20	17.91	11.51
Q	2.86	14.02	24.54	18.84	31.39	21.95	26.03	28.24	26.61	20.06	23.90	12.95
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.05	100.04	100.02	100.02	100.03	100.02	100.01	100.02	100.02	100.03	100.03	100.04

OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	29.90	20.05	8.66	11.85	11.12	12.64	21.02	11.70	2.74	15.40	30.76	21.94
CPX	0.00	0.00	1.36	0.00	0.00	2.11	0.00	0.00	15.47	0.00	0.00	5.21
NCI	36.6	24.8	13.6	15.8	14.9	18.8	25.9	15.3	21.9	20.3	36.3	32.2
NPC	40.5	66.0	25.9	10.6	41.2	32.8	18.9	32.5	38.7	23.8	25.5	45.2
DI	37.8	62.1	76.7	74.5	71.9	65.5	61.6	67.6	58.9	64.3	48.5	44.2

CIPW WEIGHT PERCENT NORM

SAMPLE	AR-236	AR-238	AR-241	AR-244	DR-247	DR-251	DR-254	DR-257	DR-259	DP-261	DP-263	DP-265
AP	.54	.50	.54	.59	1.33	.38	.33	.33	.85	.66	.36	.50
IL	2.48	2.22	2.33	2.27	3.08	1.24	1.28	1.16	2.44	1.58	2.01	1.24
OR	47.92	5.53	11.57	6.48	15.89	10.25	4.56	11.79	9.12	20.69	7.34	92.91
AB	2.40	61.97	53.22	47.98	32.77	53.49	52.18	38.38	47.28	67.30	60.99	4.99
AN	49.68	32.50	35.21	45.54	51.34	36.26	43.27	49.83	43.61	12.01	31.67	2.10
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	4.15	1.49	0.00	4.55
MT	1.32	2.62	3.22	3.73	3.31	2.45	3.19	2.44	4.41	1.75	3.04	.64
DIEN	5.03	2.64	1.71	3.28	.89	2.92	3.55	5.34	0.00	0.00	1.13	0.00
DIFS	5.96	6.31	3.31	4.95	1.91	2.33	3.23	3.52	0.00	0.00	1.38	0.00
DIWO	11.07	8.62	4.89	8.15	7.72	5.43	6.95	9.28	0.00	0.00	2.53	0.00
WO	6.86	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	0.00	2.98	7.03	9.97	6.95	11.28	9.51	14.93	13.57	3.88	12.04	2.12
HYFS	0.00	7.13	13.58	15.02	14.85	9.02	8.64	9.84	23.73	8.89	14.68	2.63
Q	10.41	9.23	5.68	.03	21.10	0.00	0.00	2.03	2.89	27.19	2.48	54.45
FO	0.00	0.00	0.00	0.00	0.00	1.74	4.43	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	1.54	4.43	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.03	100.01	100.00	100.01	100.02	100.01	100.01	100.01	100.01	100.02	100.00	100.02
OL	0.00	0.00	0.00	0.00	0.00	3.28	8.86	0.00	0.00	0.00	0.00	0.00
OPX	0.00	10.12	20.61	24.99	21.80	20.30	18.14	24.77	37.31	12.77	26.72	4.75
CPX	22.06	17.58	9.91	16.38	5.52	10.68	13.72	18.14	0.00	0.00	5.04	0.00
NCI	25.9	32.5	36.1	47.4	33.7	37.9	45.2	46.5	44.2	16.1	36.8	6.6
NPC	95.4	34.4	39.8	48.7	61.0	40.4	45.3	56.5	48.0	15.1	34.2	29.6
DI	38.8	48.2	43.1	28.4	42.5	39.3	30.9	27.7	29.9	75.2	43.7	87.6

SAMPLE	DP-269	DR-273	DP-276	DR-278	DR-281	DR-285	DR-286	DP-288	KA-291	KN-295	AR-300	AR-304
AP	.52	.40	.28	.90	.36	.57	.81	.62	.33	.28	.52	.54
IL	1.46	2.14	.98	2.33	1.26	1.52	2.63	1.52	1.95	1.99	2.27	2.37
OR	34.84	22.57	37.31	8.02	6.33	14.60	9.03	28.90	7.29	9.51	34.18	31.48
AB	47.50	36.89	54.30	55.34	41.80	45.49	46.36	50.99	55.47	50.48	43.21	47.04
AN	17.66	40.54	8.39	36.64	51.87	39.91	44.61	20.11	37.24	40.01	22.61	21.48
C	0.00	0.00	2.70	.16	0.00	0.00	0.00	0.00	0.00	0.00	1.71	0.00
MT	1.93	3.41	1.30	2.91	2.75	2.55	3.68	1.65	3.03	3.48	2.52	1.99
DIEN	.11	.08	0.00	0.00	4.90	2.68	.50	.41	.51	.72	0.00	2.18
DIFS	.95	.23	0.00	0.00	3.59	2.85	1.21	1.47	.65	1.04	0.00	2.95
DIWO	.97	.29	0.00	0.00	8.83	5.60	1.64	1.76	1.17	1.74	0.00	5.12
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	1.11	5.83	1.67	14.64	15.72	10.17	7.40	1.93	11.99	12.16	8.99	4.94
HYFS	9.15	17.88	6.85	15.15	11.49	10.81	18.07	6.96	15.38	17.55	12.83	6.67
Q	26.98	17.92	37.59	3.52	.72	3.74	11.72	28.27	5.80	5.73	16.53	8.69
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.00	100.01	99.99	100.00	100.01	100.01	100.03	100.00	100.00	100.00	100.01	100.02
OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	10.25	23.71	8.52	29.79	27.20	20.99	25.47	8.89	27.37	29.71	21.82	11.61
CPX	2.03	.60	0.00	0.00	17.32	11.13	3.35	3.64	2.33	3.50	0.00	10.25
NCI	15.7	29.9	10.8	35.0	48.5	36.2	35.1	15.7	34.7	38.7	26.6	26.2
NPC	27.1	52.4	13.4	39.8	55.4	46.7	49.0	28.3	40.2	44.2	34.4	31.3
DI	73.8	48.7	82.1	41.8	25.0	39.5	40.7	72.5	42.9	38.9	58.8	59.4

SAMPLE	AR-305	AR-307	PG-312	PG-314	PG-316	PG-319	PG-322	PG-325	PG-328	PM-329	PM-332	PM-335
AP	.40	.62	1.30	1.75	1.28	1.30	1.28	1.78	1.40	1.61	1.09	1.11
IL	2.40	2.23	2.40	3.21	2.40	2.38	2.53	2.85	2.37	3.21	2.03	2.08
OR	21.41	15.54	11.20	31.26	30.01	27.70	9.30	6.50	7.24	37.37	35.20	38.14
AB	53.68	51.21	50.35	51.12	42.22	38.44	67.88	52.52	62.61	43.13	46.69	45.75
AN	24.91	33.25	38.45	17.62	27.77	33.86	22.82	40.98	30.15	19.51	18.11	16.11
C	4.44	0.00	2.26	8.35	0.00	0.00	3.20	0.00	.80	5.63	0.00	0.00
MT	3.62	2.60	3.00	3.68	2.10	2.58	2.55	2.84	2.25	2.64	1.75	1.71
DIEN	0.00	2.00	0.00	0.00	.84	2.19	0.00	.41	0.00	0.00	.10	.00
DIFS	0.00	2.17	0.00	0.00	1.50	2.94	0.00	.35	0.00	0.00	.23	.02
DIWO	0.00	4.22	0.00	0.00	2.30	5.12	0.00	.78	0.00	0.00	.32	.02
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	11.78	10.23	13.17	11.33	4.96	7.63	13.87	16.03	11.03	10.04	3.83	1.89
HYFS	19.03	11.13	15.48	18.73	8.80	10.26	12.85	13.87	11.13	12.66	8.40	8.20
Q	14.98	8.16	17.07	3.83	19.74	10.18	12.90	10.22	12.27	7.17	22.83	26.74
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.00	100.03	100.05	100.04	100.03	100.04	100.04	100.04	100.02	100.04	100.02	100.04
OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	30.81	21.36	28.65	30.06	13.76	17.88	26.72	29.90	22.16	22.69	12.23	10.09
CPX	0.00	8.39	0.00	0.00	4.64	10.25	0.00	1.54	0.00	0.00	.65	.04
NCI	36.8	34.6	34.1	37.0	22.9	33.1	31.8	37.1	26.8	28.5	16.7	13.9
NPC	31.7	39.4	43.3	25.6	39.7	46.8	25.2	43.8	32.5	31.1	27.9	26.0
DI	47.5	46.0	45.0	44.3	60.3	46.9	52.2	40.3	53.3	53.1	71.5	75.6

CIPW WEIGHT PERCENT NORM

SAMPLE	PM-338	AR-339	AR-341	AR-343	AR-345	KA-346	KA-347	KA-348	KA-349	KA-350	KO-352	KO-355
AP	1.09	.57	.64	.57	.52	.36	.33	.38	.33	.40	.31	.24
IL	2.03	2.27	2.48	2.18	1.99	2.10	1.88	2.01	1.80	2.03	1.82	1.71
OR	23.63	27.82	9.88	6.70	5.86	12.20	5.95	8.71	7.94	25.68	13.52	16.78
AR	45.00	19.50	55.27	60.29	57.30	33.30	65.11	54.52	61.30	48.72	51.58	46.17
AN	31.37	52.68	34.85	33.01	36.84	54.50	28.94	36.78	30.77	25.60	34.90	37.55
C	.71	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	1.59	2.57	3.03	3.29	3.06	3.23	2.61	3.18	2.80	2.87	3.22	3.45
DIEN	0.00	1.87	2.36	4.44	.67	2.42	2.73	3.14	2.91	2.63	2.62	1.45
DIFS	0.00	3.00	3.70	6.38	.92	2.30	3.09	4.14	3.96	3.33	3.95	1.91
DIWO	0.00	4.81	5.99	10.75	1.58	4.82	5.88	7.28	6.85	5.97	6.51	3.36
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	6.00	6.40	7.57	3.33	11.21	15.56	9.35	8.39	7.92	7.05	8.83	12.82
HYFS	7.58	10.26	11.86	4.79	15.25	14.79	10.61	11.07	10.79	8.94	13.30	16.85
Q	28.78	17.61	4.98	0.00	7.37	3.12	5.97	0.00	0.00	0.00	2.20	.13
FO	0.00	0.00	0.00	3.04	0.00	0.00	0.00	.84	.04	1.56	0.00	0.00
FA	0.00	0.00	0.00	4.82	0.00	0.00	0.00	1.22	.06	2.18	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	100.04	100.00	100.03	100.00	100.01	100.02	100.02	100.00	100.00	100.01	100.01	100.00
OL	0.00	0.00	0.00	7.86	0.00	0.00	0.00	2.07	.11	3.74	0.00	0.00
OPX	13.58	16.66	19.44	8.12	26.46	30.36	19.97	19.46	18.71	15.99	22.13	29.67
CPX	0.00	9.68	12.05	21.57	3.17	9.53	11.70	14.56	13.72	11.93	13.08	6.72
NCI	17.2	31.1	37.0	43.0	34.7	45.2	36.2	41.3	37.1	36.6	40.3	41.5
NPC	41.1	73.0	38.7	35.4	39.1	62.1	30.8	40.3	33.4	34.4	40.4	44.9
DI	64.6	41.6	42.4	37.8	43.7	26.5	46.9	36.9	43.3	46.9	39.5	36.4

SAMPLE	KO-356	AR-359	AR-360	AR-361	AR-362	AR-363	AR-364	AR-365	JC-366	JC-367	JC-368	JC-369
AP	.26	.50	.52	.66	.62	.57	.59	.54	1.18	1.18	1.21	1.21
IL	1.73	2.31	2.31	2.46	2.20	2.05	2.12	2.03	1.86	1.95	1.93	1.91
OR	7.39	20.87	11.10	19.32	11.45	16.08	21.94	19.54	17.15	15.64	2.39	6.49
AR	51.61	49.23	57.23	43.05	50.43	52.20	55.84	53.12	58.31	65.16	66.07	65.09
AN	41.00	29.90	31.67	37.63	38.12	31.72	22.22	27.33	24.54	19.20	31.54	28.42
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.52	0.00	3.23	1.56
MT	2.86	2.57	3.39	3.33	3.04	3.12	2.51	2.71	2.96	3.02	3.26	3.07
DIEN	.07	3.00	1.48	1.67	3.20	4.51	3.69	2.39	0.00	.12	0.00	0.00
DIFS	.11	5.09	2.15	2.49	4.14	5.66	4.43	4.05	0.00	.95	0.00	0.00
DIWO	.18	7.96	3.60	4.13	7.34	10.20	8.17	6.33	0.00	.98	0.00	0.00
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	8.75	4.72	10.48	9.93	9.08	8.61	6.99	5.93	2.09	1.99	2.66	2.12
HYFS	15.11	7.99	15.24	14.82	11.75	10.81	8.40	10.07	15.73	15.11	17.49	16.42
Q	11.75	7.29	0.00	6.08	4.47	.85	2.50	6.37	23.31	20.70	35.65	27.32
FO	0.00	0.00	.26	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	.41	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
TOTAL	99.99	100.03	100.01	100.03	100.03	100.02	100.02	99.99	100.03	100.02	100.02	100.03
OL	0.00	0.00	.67	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	23.86	12.71	25.72	24.76	20.83	19.42	15.40	16.00	17.82	17.10	20.16	18.54
CPX	.35	16.05	7.23	8.29	14.67	20.37	16.30	12.77	0.00	2.05	0.00	0.00
NCI	28.8	33.6	39.3	38.8	40.7	45.0	36.3	33.5	22.6	24.1	25.4	23.5
NPC	44.3	37.8	35.6	46.6	43.1	37.8	28.5	34.0	29.6	22.8	32.3	30.4
DI	46.7	48.4	41.1	40.0	38.0	37.5	49.6	49.7	62.8	64.3	59.3	60.5

SAMPLE	JC-370	JC-371	KA-373	KA-376	KL-382	KL-392	PR-395	PR-398	AR-400	AR-403	PG-406	PG-417
AP	1.16	2.04	.33	.40	.17	.17	1.61	1.68	.64	.52	1.35	1.59
IL	1.90	3.10	1.90	2.18	.94	.94	3.08	2.85	2.18	2.16	2.65	2.80
OR	22.32	16.95	9.26	17.97	23.49	14.45	5.35	18.42	20.59	14.71	29.29	17.02
AB	64.77	64.14	52.39	44.33	36.93	46.29	46.13	39.17	31.69	25.98	36.70	45.79
AN	12.91	18.92	38.35	37.71	39.57	39.26	48.53	42.41	47.72	59.31	34.00	37.19
C	2.10	6.99	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.85	0.00
MT	2.91	4.32	2.74	3.16	2.81	2.87	2.36	2.89	2.99	3.80	2.61	2.86
DIEN	0.00	0.00	4.13	1.41	4.64	2.71	7.55	1.39	3.78	.54	0.00	1.41
DIFS	0.00	0.00	5.20	2.04	3.28	1.99	4.87	1.59	5.26	.80	0.00	1.37
DIWO	0.00	0.00	9.37	3.42	8.26	4.89	13.03	3.01	9.01	1.32	0.00	2.84
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	2.19	3.96	7.37	7.83	17.60	15.07	0.00	11.19	7.52	13.16	14.19	13.38
HYFS	15.48	22.69	9.28	11.33	12.42	11.07	0.00	12.88	10.47	19.50	12.99	13.02
Q	24.50	2.24	4.79	0.00	1.12	0.00	0.00	9.64	4.76	10.02	13.76	9.04
FO	0.00	0.00	0.00	1.59	0.00	2.83	7.17	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	2.53	0.00	2.29	5.09	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	3.37	0.00	0.00	0.00	0.00	0.00
TOTAL	100.02	100.08	100.02	100.02	100.03	100.02	100.03	100.04	100.01	100.01	100.04	100.04
OL	0.00	0.00	0.00	4.13	0.00	5.11	12.26	0.00	0.00	0.00	0.00	0.00
OPX	17.67	26.65	16.65	19.15	30.02	26.14	0.00	24.07	18.00	32.66	27.19	26.41
CPX	0.00	0.00	18.70	6.86	16.17	9.59	25.45	5.99	18.05	2.65	0.00	5.61
NCI	22.5	34.1	40.0	35.5	49.9	44.6	43.1	35.8	41.2	41.3	32.4	37.7
NPC	16.6	22.8	42.3	46.0	51.7	45.9	51.3	52.0	60.1	69.5	48.1	44.8
DI	67.8	46.6	38.6	40.0	30.6	33.5	28.4	40.1	32.7	29.6	47.8	41.5

CIPW WEIGHT PERCENT NORM

SAMPLE	IN-419	PG-421	PG-426	PG-433	AR-442	AR-446	AR-451	KA-452	KA-455	KO-457	KO-460	KO-462
AP	.69	1.42	1.85	1.49	.66	.57	.38	.33	.38	.28	.28	.28
IL	3.40	2.61	3.00	2.95	2.40	2.01	1.67	1.95	2.01	1.88	1.76	1.76
OR	4.97	24.18	28.20	6.49	4.82	16.30	22.78	4.77	19.50	10.31	15.09	28.09
AB	50.88	51.26	39.18	59.64	52.39	55.91	54.86	39.73	45.49	51.08	56.63	41.19
AN	44.15	24.55	32.61	33.87	42.79	27.79	22.36	55.50	35.00	38.62	28.29	30.72
C	0.00	0.00	0.00	.79	.43	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	4.19	2.60	2.75	2.71	3.25	2.75	2.10	3.18	3.06	3.45	3.13	3.35
DIEN	1.34	1.79	1.26	0.00	0.00	.74	3.99	3.96	2.20	1.45	2.61	1.70
DIFS	1.81	2.20	1.26	0.00	0.00	1.08	3.92	4.13	3.20	1.96	3.44	2.53
DIWO	3.15	4.00	2.56	0.00	0.00	1.81	8.07	8.22	5.36	3.40	6.04	4.19
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	8.53	8.75	12.32	13.87	12.75	9.19	7.12	12.25	8.91	12.23	6.98	9.96
HYFS	11.52	10.78	12.39	13.23	16.97	13.33	7.00	12.79	12.98	16.59	9.20	14.83
Q	0.00	8.96	6.30	2.16	3.47	8.01	3.75	6.14	2.21	6.60	0.00	0.00
FO	4.29	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.22	.32
FA	6.38	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	3.22	.52
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00

TOTAL	100.01	100.04	100.05	100.04	100.01	100.03	100.00	100.02	100.01	100.02	100.00	100.02
OL	10.67	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	5.44	.83
OPX	20.05	19.53	24.70	27.10	29.72	22.53	14.12	25.04	21.89	28.82	16.18	24.79
CPX	6.30	7.99	5.08	0.00	0.00	3.63	15.98	16.32	10.76	6.80	12.09	8.41
NCI	44.6	32.7	35.5	32.8	35.4	30.9	33.9	46.5	37.7	41.0	38.6	39.1
NPC	46.5	32.4	45.4	36.2	45.0	33.2	29.0	58.3	43.5	43.1	33.3	42.7
DI	30.6	51.9	44.3	43.7	37.8	51.7	51.9	27.1	41.0	38.6	43.8	42.0

SAMPLE	KL-465	KL-467	KL-468	KL-469	KL-471	PR-472	PR-474	PR-476	PR-477	PR-479	PR-480	PR-481
AP	.28	.19	.09	.19	.28	1.54	1.28	1.40	1.56	1.40	1.56	1.63
IL	1.63	1.37	.75	1.18	1.11	3.10	2.35	2.52	2.67	2.52	2.76	2.65
OR	17.41	8.91	2.43	15.20	26.13	20.49	26.94	24.73	7.36	19.42	19.19	2.49
AB	50.01	48.21	6.62	51.28	55.09	55.38	41.82	51.06	50.01	40.16	46.11	60.84
AN	32.58	42.88	90.94	33.52	18.78	24.13	31.24	24.21	42.63	40.42	34.70	36.67
C	0.00	0.00	0.00	0.00	2.94	1.90	0.00	0.00	0.00	0.00	0.00	4.22
MT	3.10	2.81	3.15	2.87	2.73	2.83	2.62	1.91	2.62	2.67	2.58	3.00
DIEN	3.04	2.98	4.74	2.46	0.00	0.00	2.82	1.28	.02	1.15	.85	0.00
DIFS	4.93	3.34	1.96	2.37	0.00	0.00	2.76	1.53	.02	1.21	1.00	0.00
DIWO	7.87	6.39	7.20	4.94	0.00	0.00	5.69	2.83	.04	2.39	1.86	0.00
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	7.27	10.67	38.32	13.03	10.81	7.32	10.88	6.24	11.91	11.68	9.98	19.50
HYFS	11.77	11.99	15.84	12.57	14.98	13.91	10.63	7.47	13.00	12.29	11.71	15.19
Q	3.09	6.55	3.67	0.00	1.59	5.39	6.17	13.21	14.40	12.71	7.35	13.53
FO	0.00	0.00	0.00	.64	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
FA	0.00	0.00	0.00	.68	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00

TOTAL	100.01	100.00	100.00	100.02	100.01	100.02	100.02	100.03	100.04	100.03	100.03	100.05
OL	0.00	0.00	0.00	1.32	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
OPX	19.04	22.66	54.17	25.60	25.79	21.24	21.51	13.71	24.91	23.96	21.70	34.69
CPX	15.84	12.70	13.90	9.77	0.00	0.00	11.27	5.64	.08	4.75	3.71	0.00
NCI	39.6	39.5	72.0	40.8	29.6	27.2	37.7	23.8	30.3	33.9	30.7	40.3
NPC	39.5	47.1	93.2	39.5	25.4	30.4	42.8	32.2	46.0	50.2	42.9	37.6
DI	41.5	37.2	5.9	39.3	54.9	54.0	43.9	59.9	45.3	43.7	46.8	39.1

SAMPLE	PR-483	PR-484	PR-485	PR-486	PR-487	PR-489	PR-491	PR-492	PR-494	AR-495	AR-496	AR-497
AP	1.28	1.54	1.68	1.47	1.78	1.42	1.47	1.87	2.06	.62	.62	.54
IL	2.31	2.74	2.82	2.55	2.68	2.57	2.57	3.25	3.30	2.37	2.23	2.25
OR	33.02	24.87	17.23	16.51	11.20	22.55	24.28	20.90	26.19	17.85	14.67	3.44
AB	31.27	31.43	45.37	37.01	54.57	40.56	43.91	43.23	37.50	32.24	37.26	49.45
AN	35.71	43.69	37.40	46.48	34.23	36.89	31.80	35.87	36.31	49.90	48.07	47.11
C	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MT	2.62	2.86	3.04	3.10	2.16	2.48	2.57	2.64	3.45	3.29	3.16	3.52
DIEN	1.82	2.36	1.83	1.88	2.72	1.22	.86	.88	2.11	1.06	2.70	1.79
DIFS	2.43	2.28	2.08	1.49	3.04	1.55	1.14	1.58	2.81	1.45	4.45	2.35
DIWO	4.24	4.74	3.95	3.49	5.82	2.77	2.00	2.42	4.92	2.50	7.05	4.14
WO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
HYEN	8.17	12.55	11.72	18.34	6.57	8.45	8.78	6.19	10.96	11.42	7.33	12.43
HYFS	10.92	12.09	13.34	14.48	7.35	10.74	11.70	11.06	14.56	15.67	12.09	16.32
Q	11.32	8.27	6.42	9.55	14.12	14.61	12.49	16.71	.82	3.39	5.20	0.00
FO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.02
FA	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.03
NE	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00

TOTAL	100.03	100.04	100.04	100.01	100.04	100.04	100.03	100.03	100.05	100.01	100.02	100.03
OL	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	.04
OPX	19.09	24.65	25.06	32.82	13.92	19.19	20.48	17.25	25.52	27.09	19.42	28.75
CPX	8.48	9.38	7.87	6.85	11.58	5.53	4.00	4.88	9.84	5.01	14.20	8.28
NCI	32.5	39.6	38.8	45.3	30.3	29.8	29.6	28.0	42.1	37.8	39.0	42.9
NPC	53.3	58.2	45.2	55.7	38.5	47.6	42.0	45.3	49.2	60.7	56.3	48.8
DI	46.6	36.8	39.7	32.9	49.5	48.8	51.0	51.0	35.9	32.6	33.9	30.0

CIPW WEIGHT PERCENT NORM

SAMPLE	AR-498	AR-500
AP	.54	.52
IL	2.14	1.90
OR	2.19	13.41
AB	29.53	27.58
AN	24.26	17.30
C	0.00	0.00
MT	2.94	2.89
DIEN	.39	1.86
DIFS	.43	2.34
DIWO	.82	4.21
WO	0.00	0.00
HYFN	13.46	10.37
HYFS	14.99	13.00
Q	8.33	4.64
FO	0.00	0.00
FA	0.00	0.00
NE	0.00	0.00
TOTAL	100.02	100.01
OL	0.00	0.00
OPX	28.45	23.37
CPX	1.64	8.41
NCI	35.2	36.6
NPC	45.1	38.5
DI	40.0	45.6